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# Interactive effects of MnO<sub>2</sub>, organic matter and pH on abiotic formation of N<sub>2</sub>O from hydroxylamine in artificial soil mixtures

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Abiotic conversion of the reactive nitrification intermediate hydroxylamine (NH $_2$ OH) to nitrous oxide (N $_2$ O) is a possible mechanism of N $_2$ O formation during nitrification. Previous research has demonstrated that manganese dioxide (MnO $_2$ ) and organic matter (OM) content of soil as well as soil pH are important control variables of N $_2$ O formation in the soil. But until now, their combined effect on abiotic N $_2$ O formation from NH $_2$ OH has not been quantified. Here, we present results from a full-factorial experiment with artificial soil mixtures at five different levels of pH, MnO $_2$  and OM, respectively, and quantified the interactive effects of the three variables on the NH $_2$ OH-to-N $_2$ O conversion ratio (R $_{NH2OH-to-N2O}$ ). Furthermore, the effect of OM quality on R $_{NH2OH-to-N2O}$  was determined by the addition of four different organic materials with different C/N ratios to the artificial soil mixtures. The experiments revealed a strong interactive effect of soil pH, MnO $_2$  and OM on R $_{NH2OH-to-N2O}$ . In general, increasing MnO $_2$  and decreasing pH increased R $_{NH2OH-to-N2O}$ , while increasing OM content was associated with a decrease in R $_{NH2OH-to-N2O}$ . Organic matter quality also affected R $_{NH2OH-to-N2O}$ . However, this effect was not a function of C/N ratio, but was rather related to differences in the dominating functional groups between the different organic materials.

Nitrous oxide  $(N_2O)$  is a potent greenhouse gas that can be formed by several soil processes, such as microbial nitrification and denitrification. The  $N_2O$  production from nitrification, especially from its reactive intermediate hydroxylamine  $(NH_2OH)$ , has received increasing attention in the recent past, fostered by the development of analytical techniques for the determination of the  $^{15}N$  site preference in the  $N_2O$  molecule that allows for constraining the contribution of different source processes to total  $N_2O$  formation  $^{1-4}$ . Also, increasing knowledge from molecular biological and genetic studies has contributed to elucidating the different  $N_2O$  formation mechanisms during nitrification<sup>3</sup>. Still, the role of  $N_2OH$  in  $N_2O$  formation in the soil is insufficiently understood. While there is evidence, e.g., from measurements in wastewater treatment systems that  $NH_2OH$  can contribute about 65% of total  $N_2O$  formation<sup>2</sup>, the formation of  $N_2O$  from  $NH_2OH$  in soil and its controlling factors have rarely been studied  $^{5,6}$ .

Hydroxylamine was first identified by Lees  $(1952)^7$  as an intermediate of the first step of nitrification by ammonia oxidizing bacteria (AOB), in which ammonia is oxidized to nitrite. Understanding the nitrification process in ammonia-oxidizing archaea (AOA), however, is much more fragmentary, but NH<sub>2</sub>OH has been identified as an intermediate of ammonia oxidation also in AOA<sup>8</sup>. In most circumstances, NH<sub>2</sub>OH is quickly oxidized to nitrite in the periplasm of the AOB, and N<sub>2</sub>O may be produced as a side product during this process<sup>3</sup>. However, also a leakage of NH<sub>2</sub>OH from the periplasm across the outer membrane of the AOB into the soil matrix, followed by a chemical reaction with soil constituents yielding N<sub>2</sub>O, could be a potential mechanism of N<sub>2</sub>O formation during nitrification. This assumption is supported by the fact that AOB can take up NH<sub>2</sub>OH from the surrounding medium<sup>9</sup> as well as by the observation that the medium of AOB cultures contains measurable amounts of NH<sub>2</sub>OH. The latter was found for *Nitrosomonas europaea* under oxic conditions, both for wild-type *N. europaea* and even more so for NirK and NorB-deficient mutants<sup>10</sup>. In accordance with this assumption, a positive relationship between NH<sub>2</sub>OH content of the soil and soil N<sub>2</sub>O emissions under oxic conditions has been detected in

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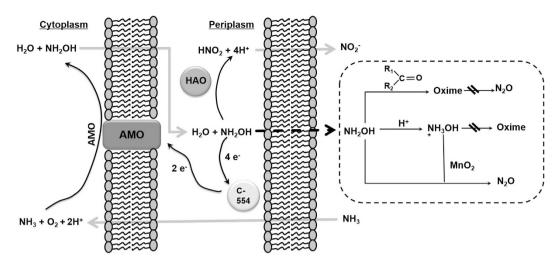


Figure 1. Hypothetical model of  $NH_2OH$  release by ammonia-oxidizing bacteria to the soil environment and potential abiotic reactions of  $NH_2OH$  with  $MnO_2$  and organic matter in the soil at different pH conditions ( $R_1R_2C=O$  represents carbonyl groups of SOM). AMO is ammonia monooxygenase; HAO is hydroxylamine oxidoreductase.

natural forest soil samples  $^{11}$ . In addition, abiotic formation of  $N_2O$  from  $NH_2OH$  has been observed in sterilized soil samples from different ecosystems  $^6$ .

In soil,  $N_2O$  can be formed chemically, among a range of possible reactions, according to the following equations<sup>12</sup>:

$$NH_2OH + NO_2^- \rightarrow N_2O + H_2O + OH^-$$
 (1)

$$2MnO_2 + 2NH_2OH \rightarrow 2MnO + N_2O + 3H_2O.$$
 (2)

Owing to its high oxidization potential, manganese dioxide  $(MnO_2)$  acts as a strong oxidant in soil that plays an important role not only in the turnover of organic substances<sup>13,14</sup>, but also in the N cycle<sup>15</sup>, even under anoxic conditions<sup>16,17</sup>. Soil organic matter (SOM) plays a crucial role in the storage and release of N as well as in the emission of N<sub>2</sub>O from soils. Quick disappearance of nitrite and nitrate within a few hours after addition has been observed in forest soils<sup>18–20</sup>, whereas NH<sub>2</sub>OH disappeared completely in soil several minutes after addition<sup>5,11</sup>. Abiotic reactions of SOM and inorganic N may contribute to the quick disappearance, as nitrite and nitrate can react with SOM or dissolved organic carbon (DOC), leading to the formation of organic N, such as nitroso and nitro compounds<sup>21,22</sup>, while NH<sub>2</sub>OH can also react with carbonyl groups to form oximes<sup>23,24</sup>:

$$R_1(R_2)C = O + NH_2OH \rightarrow R_1(R_2)C = NOH + H_2O$$
 (3)

The quality of SOM, or more specifically the C/N ratio and the type and abundance of functional groups, influence the bonding of inorganic N to  $SOM^{22}$ . Phenolic lignin derivatives, an important constituent of SOM, can covalently bind reactive N compounds and thereby stabilize N in soil<sup>25,26</sup>. The N binding form can be affected by the plant species from which the SOM is derived due to the different characteristics of phenolic compounds, e.g. condensed or hydrozable tannin<sup>27</sup>.

Soil pH is another key factor influencing most nitrogen transformations in soil. High pH can lead to an increase of chemical  $N_2O$  production involving nitrite by favoring nitrite accumulation, either directly through increasing nitrite stability, or indirectly by inhibiting biological nitrite oxidation due to a higher concentration of free NH<sub>3</sub> (an inhibitor of nitrite oxidizers) in the soil<sup>28</sup>. In contrast, high soil  $N_2O$  emissions have also been observed in acid forest soils<sup>29,30</sup>. In this case, the effect of pH on enzyme activities during denitrification and nitrification was suggested as the main reason<sup>31</sup>. However, also chemical reactions that produce  $N_2O$  in the soil, such as the reaction of nitrite with SOM and the reaction of NH<sub>2</sub>OH with MnO<sub>2</sub>, are subject to a strong pH dependence and can contribute substantially to  $N_2O$  emissions under acidic conditions<sup>32–34</sup>.

The aim of this study was to quantify the interactive effects of the major control factors of abiotic  $N_2O$  formation from  $NH_2OH$  in soil, i.e.  $MnO_2$  content, pH and OM quantity and quality, by means of experiments with artificial soil mixtures. We hypothesized that the control factors interact with each other in the following way: At higher pH, unprotonated  $NH_2OH$  would react more readily with carbonyl groups of OM, leading to oxime formation and making  $NH_2OH$  less available for oxidation to  $N_2O$  by  $MnO_2$ . Lower soil pH would lead to increased protonation of  $NH_2OH$ , making  $NH_2OH$  more stable against the reaction with carbonyl groups of OM and more prone to the reaction with  $MnO_2$ , leading to higher  $N_2O$  formation from the same amount of  $NH_2OH$  (Fig. 1). To test these hypotheses, we performed two laboratory experiments with artificial soil mixtures, which were produced from pure quartz sand, quartz powder, kaolin clay,  $MnO_2$  powder and different plant-derived organic materials, resembling SOM of different quality, at different mixing ratios. In these experiments,  $N_2O$  formation was determined after  $NH_2OH$  addition to the different mixtures at different pH levels and related to the different control factors.

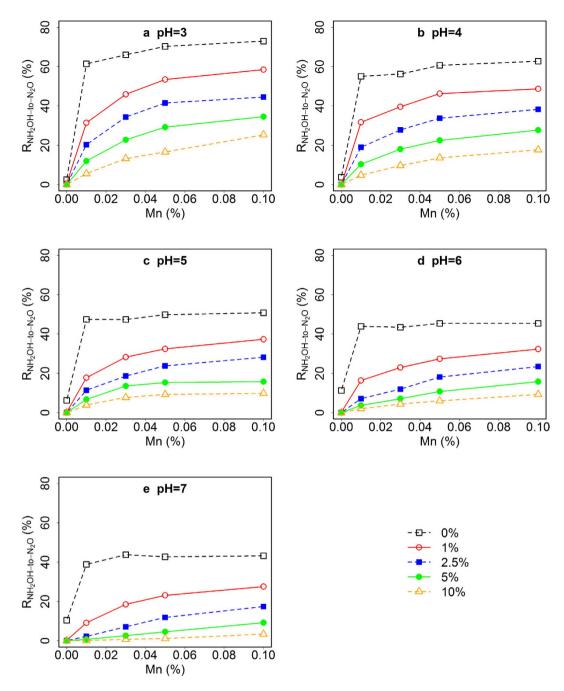


Figure 2. Mean NH<sub>2</sub>OH-to-N<sub>2</sub>O conversion ratios ( $R_{\rm NH2OH-to-N2O}$ ) in artificial soil mixtures at different pH as well as MnO<sub>2</sub> and organic matter (OM, peat moss) contents. The total amount of NH<sub>2</sub>OH added was 5 nmol (equivalent to 70 µg N per kg dry material). Different symbols represent  $R_{\rm NH2OH-to-N2O}$  at different OM content (n = 3, SD < 5%, not shown).

### **Results and Discussion**

 $R_{NH2OH-to-N2O}$  at different pH, MnO<sub>2</sub> and OM contents (%). In the present study all three factors, i.e. pH, MnO<sub>2</sub> and OM content, affected  $R_{NH2OH-to-N2O}$  from peat moss significantly (Fig. 2, S1 and S2). The  $R_{NH2OH-to-N2O}$  increased greatly with an increase in MnO<sub>2</sub> content from 0% to 0.1% (Fig. 2). This finding is consistent with Bremner *et al.*<sup>5</sup>, who studied 19 soils with a wide range of properties and found that the formation of  $N_2O$  by decomposition of  $N_2OH$  was highly correlated with oxidized Mn content of the soils. The fact that  $NH_2OH$  was used in the past for the selective extraction of Mn oxides from soil samples<sup>35</sup> indicates that  $NH_2OH$  can efficiently reduce Mn(IV) to Mn(II) or Mn(III) (and in turn is oxidized to  $N_2O$ ) in natural soil samples. With increasing OM content,  $R_{NH2OH-to-N2O}$  decreased remarkably, especially at high pH (Fig. 2c,d,e). For example, an increase in OM by only 1% at 0.01% MnO<sub>2</sub> led to about 50% and 80% decrease in  $N_2O$  emissions at pH 3 and pH 7, respectively (Fig. 2e, S2). This could be caused by the oxime-forming reaction between  $NH_2OH$  and carbonyl groups of OM, such as in quinones. The oximes may undergo a tautomeric equilibrium with their corresponding nitrosophenol

forms  $^{23}$ . In fact, NH $_2$ OH has been used in a number of previous studies to determine the carbonyl content of humic substances  $^{36}$ , indicating a high affinity of NH $_2$ OH to OM that contains carbonyl groups. In the absence of OM and MnO $_2$ , increasing pH led to a slight increase in R $_{\rm NH2OH-to-N2O}$  due to the self-decomposition of NH $_2$ OH at higher pH, whereas in the presence of OM and absence of MnO $_2$  nearly no NH $_2$ OH was converted to N $_2$ O (Fig. S2a). In contrast, the effect of increasing pH on R $_{\rm NH2OH-to-N2O}$  became negative already in the presence of 0.01% MnO $_2$  (Fig. S2b). This finding suggests that acidic conditions are favorable for the redox reaction between NH $_2$ OH and MnO $_2$ .

Also strong interactive effects of pH and MnO<sub>2</sub>, pH and OM, and OM and MnO<sub>2</sub> were observed for the conversion of NH<sub>2</sub>OH to N<sub>2</sub>O. The largest  $R_{\rm NH2OH-to-N2O}$  found in the present experiment was 81.5% in the absence of SOM at pH 3, and with a MnO<sub>2</sub> content of 0.1% (Fig. 2a), while the lowest  $R_{\rm NH2OH-to-N2O}$  was about 9%, when SOM content was 10% in the presence of 0.1% MnO<sub>2</sub> at pH 7 (Fig. 2e). This suggests that even at the highest MnO<sub>2</sub> level and in all other respects optimal conditions a small fraction of NH<sub>2</sub>OH had not been converted to N<sub>2</sub>O, but to some other unidentified product.

In the treatments without OM,  $MnO_2$  had only a small effect on  $R_{NH2OH-to-N2O}$  at all pH conditions, while it had a larger effect especially at higher OM content (Fig. 2, S1), suggesting a strong competition between OM and  $MnO_2$  for  $NH_2OH$ . The competition was biased by pH, with lower pH favouring the reaction of  $NH_2OH$  and  $MnO_2$ , while higher pH favoured the reaction of  $NH_2OH$  with OM. These findings confirmed our hypothesis that at low pH  $NH_2OH$  is more protected against reaction with OM and more available for the oxidation by  $MnO_2$  due to the higher degree of  $NH_2OH$  protonation at lower pH.

 $R_{NH2OH-to-N2O}$  as a function of pH, MnO<sub>2</sub> content and OM quality. Organic matter quality had a clear influence on  $R_{NH2OH-to-N2O}$  in this study (Fig. 3, S3, and S4). Most of the OM types were associated with a significantly lower  $R_{NH2OH-to-N2O}$  compared to the mixtures without OM within the pH range of the experiment. In general, the inhibitory effect of the organic materials on the conversion of  $NH_2OH$  to  $N_2O$  showed a clear pH dependency, but was not a function of C/N ratio (Fig. 3, S3). At acidic conditions (pH 3–4), peat moss and watermilfoil with their relatively large C/N ratio inhibited  $R_{NH2OH-to-N2O}$  the least, while the cyanobacterium material and clover had a stronger inhibitory effect on  $R_{NH2OH-to-N2O}$  despite their smaller C/N ratio (Fig. 3a,b). The differences between peat moss, cyanobacterium and watermilfoil material as OM became smaller at higher pH, and were no longer significant at pH 7 in the presence of 0.01% MnO<sub>2</sub> (Fig. 3e), while clover showed always the smallest  $R_{NH2OH-to-N2O}$  at all pH levels. In the absence of MnO<sub>2</sub>, all OM forms showed a  $R_{NH2OH-to-N2O}$  close to zero, except for the watermilfoil material that was associated with a  $R_{NH2OH-to-N2O}$  significantly above zero within the pH range 3–6 (Fig. S4a). A possible explanation could be the fact that, in contrast to the other OM sources, the watermilfoil material contained about 0.03% Mn (Table 1), which could have caused the  $N_2O$  emission after  $NH_2OH$  addition even without external MnO<sub>2</sub> addition.

We assumed that  $R_{\rm NH2OH-to-N2O}$  would be a function of the C/N ratio of the different SOM types, as larger C/N ratios would be indicative of a lower degree of N-containing functional groups, i.e. leaving a higher chance for NH<sub>2</sub>OH to react with SOM and not to be converted to N<sub>2</sub>O. However, as stated above we did not observe any clear relationship between C/N ratio and  $R_{\rm NH2OH-to-N2O}$ , e.g. peat moss had the largest C/N ratio, but did not lead to the lowest  $R_{\rm NH2OH-to-N2O}$ . Instead, clover with a much lower C/N ratio had the largest inhibitory effect on  $R_{\rm NH2OH-to-N2O}$ . The addition of 2.5% dry clover powder (C/N ratio = 11.3) to the artificial soil mixture decreased  $R_{\rm NH2OH-to-N2O}$  by 48% at pH 3 (Fig. 3a), which was similar to the effect of 10% peat moss (C/N ratio = 67.2) at the same pH (Fig. 2a). The reason for this observation could lie in the differences in functional groups between the different organic materials used in this study.

A better insight into the effects of C and N functional groups of the different organic materials was obtained from NMR analysis. The peat moss OM had the lowest proportion of ester or amide carbonyl at around 170 ppm of all materials (Fig. 4, Table 2). This is in accordance with the observation that – despite having the largest C/N ratio – peat moss OM had a lower inhibitory effect on  $R_{\rm NH2OH-to-N2O}$  compared to clover and watermilfoil OM (if the background MnO<sub>2</sub> effect was subtracted), i.e. the lack of almost any carbonyl groups in peat moss was clearly visible in its chemical behaviour toward NH<sub>2</sub>OH. In addition, peat moss OM exhibited the largest proportion of O-substituted aliphatic compounds, which might have also contributed to the relatively low inhibitory effect on  $R_{\rm NH2OH-to-N2O}$  in comparison to clover and watermilfoil OM. In contrast, cyanobacterium OM had the highest proportion of acid/amide carbonyl of all four organic materials, suggesting the highest inhibitory effect on  $R_{\rm NH2OH-to-N2O}$  due to the competitive reaction of carbonyl groups with NH<sub>2</sub>OH. The clover material, however, contained lower amounts of O-substituted aliphatics and di-O-substituted C in comparison to peat moss and watermilfoil OM, which may have increased its affinity for NH<sub>2</sub>OH. For the proportion of unsaturated C no clear trend emerged across the different materials, suggesting that the effect of unsaturated C on  $R_{\rm NH2OH-to-N2O}$  is of minor importance.

**Development of a stepwise multiple regression model from the artificial soil mixtures and application to natural soils.** The multiple regression model obtained from the first experiments was  $R_{\rm NH2OH-to-N2O}=45.9-3.1~{\rm SOM}+241.1~{\rm MnO}_2-4.5~{\rm pH},~R^2=0.62~(P<0.01),$  which could explain about 62% variation of  $R_{\rm NH2OH-to-N2O}$ , and the contributions of pH, Mn and SOM content to the model's performance were all significant (P<0.01). It could well explain the observations (Fig. 3) for peat moss, watermilfoil and clover OM ( $R^2$  close to 0.8, P<0.01, Fig. 5). This demonstrated the general applicability of the model for the OM derived from the different plant and cyanobacterium materials, with different N content, aliphatic C content and C/N ratios. In contrast, the model proved to be not appropriate for the artificial soil mixture without any MnO<sub>2</sub>, indicated by the decreased goodness of the simulation.

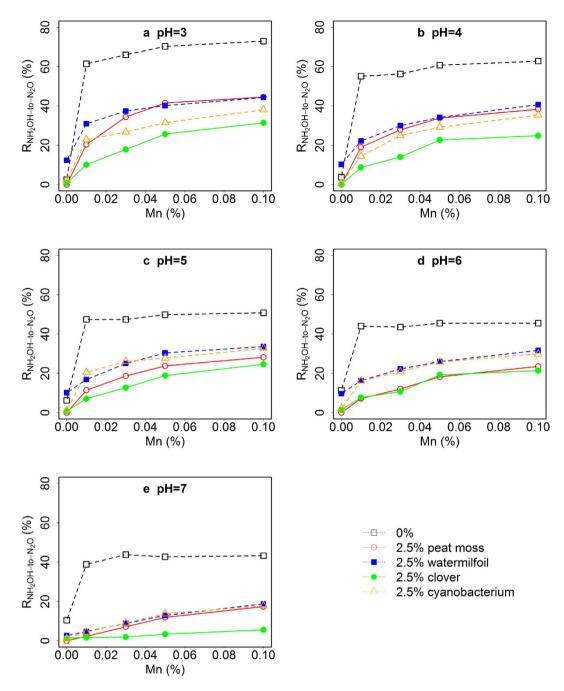


Figure 3. Mean NH<sub>2</sub>OH-to-N<sub>2</sub>O conversion ratios ( $R_{\rm NH2OH-to-N2O}$ ) in artificial soils at different pH and MnO<sub>2</sub> content, and for organic matter of different origins at a fixed content of 2.5% (w/w). The total amount of NH<sub>2</sub>OH added was 5 nmol. Different symbols represent  $R_{\rm NH2OH-to-N2O}$  for the artificial soil mixtures with the different organic materials (n = 3, SD < 5%, not shown).

	С	N	C/N	Al	Ca	Fe	K	Mg	Mn	Na	P	Si
Peat moss	41.3!	0.6	67.2	0.03	0.13	0.06	0.06	0.07	< 0.01	0.01	0.03	0.08
Watermilfoil	35.4	2.1	17.0	0.12	2.26	0.11	1.21	0.25	0.031	0.67	0.12	0.21
Clover	41.4	3.7	11.3	< 0.01	1.10	0.01	2.68	0.20	< 0.01	< 0.01	0.34	0.03
Cyanobacterium	44.9	9.9	4.5	0.02	0.31	0.09	1.22	0.31	< 0.01	1.36	0.92	0.07

**Table 1.** Element contents (%) and C/N ratios of the organic materials used in this study. All elements are reported as % of dry weight (mean of three replicates). The standard deviation is 3% for the values larger than 1%, 20% for the values smaller than 0.1%, and 10% for the values in the range of 0.1% to 1%.

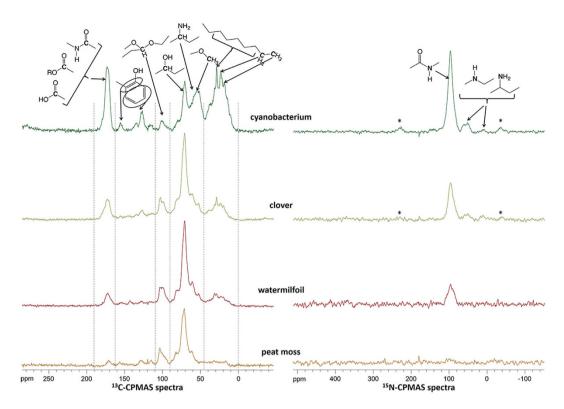


Figure 4. <sup>13</sup>C- and <sup>15</sup>N-CPMAS-NMR spectra of the different organic materials (cyanobacterium, clover, watermilfoil, peat moss) used in the experiment.

Spectral range (ppm)	Chemical features	Found in	Cyanobacterium (%)	Clover (%)	Watermilfoil (%)	Peat moss (%)
45-0	Aliphatic compounds	waxes, suberin, cutin, cyanophycin, chlorophyll (a,b,d)	41	17	15	11
64.5–45	N- and O-substituted aliphats	amino acids, amino sugars, lignin, cyanophycin	19	14	14	12
90-64.5	O-substituted aliphats	polysaccharides, cellulose, hemi-cellulose, starch, pectin, lignin	14	38	42	49
109-90	di-O-substituted C	polysaccharides, cellulose, hemi-cellulose, starch, pectin	3	11	12	14
162-109	unsaturated C, aromatic C	suberin, lignin, chlorophyll	7	11	10	11
190-162	acid, ester, amide	cutin, proteins, cyanophycin, chlorophyll	17	10	7	4

Table 2. Relative proportions of chemical features of the different plant materials derived from <sup>13</sup>C CPMAS NMR spectra. Sums within columns greater than 100 are due to rounding errors.

Finally,  $R_{\rm NH2OH-to-N2O}$  was simulated with the same regression model for the natural soils described in Heil et al.<sup>6</sup>. The results showed that the application of the model to natural soils was promising, no matter if it was applied to fumigated or fresh soils (Fig. 6). The simulated  $R_{\rm NH2OH-to-N2O}$  explained more than 90% of the observed rates, especially for cropland, grassland and deciduous forest soils. However, the model failed at correctly predicting  $R_{\rm NH2OH-to-N2O}$  for the spruce forest soil of Heil et al.<sup>6</sup>, which could be related to the high SOM and relatively low  $MnO_2$  content of the spruce soil as compared to the other soils. This finding suggests that there is a threshold value for the SOM content of 10% above which – and a  $MnO_2$  content of 0.01% below which – the model fails to predict the correct  $R_{\rm NH2OH-to-N2O}$  values.

Soil pH, MnO $_2$  and SOM content were identified as crucial control variables of  $R_{\rm NH2OH-to-N2O}$ , i.e. the conversion ratio of NH $_2$ OH to N $_2$ O in the artificial soil experiments of this study. Organic matter derived from different plant species and a cyanobacterium also affected  $R_{\rm NH2OH-to-N2O}$  due to the differences in composition, type and abundance of functional groups, as more carbonyl C leads to higher reactivity of NH $_2$ OH with organic matter, thereby lowering its availability for the oxidation to N $_2$ O by MnO $_2$ . The multiple regression model of pH, MnO $_2$  and OM developed here could explain about 60% of the variance of  $R_{\rm NH2OH-to-N2O}$  in the artificial soil mixtures, and proved also to be promising for the prediction of  $R_{\rm NH2OH-to-N2O}$  of chemical N $_2$ O production from NH $_2$ OH in natural soils, when SOM content was below 10% and Mn content was larger than 0.01%. If these findings can

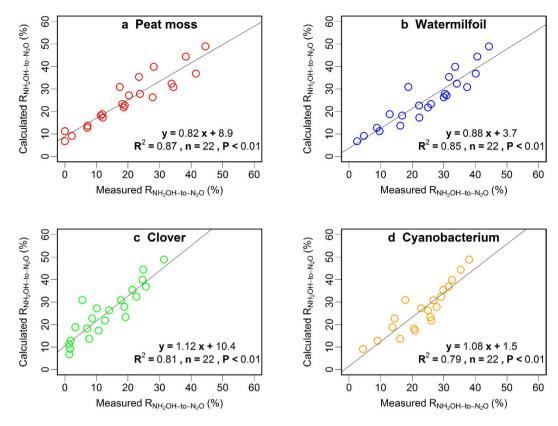


Figure 5. Results of the application of the artificial soil regression model for the calculation of NH<sub>2</sub>OH-to-N<sub>2</sub>O conversion ratios ( $R_{\rm NH2OH\text{-}to\text{-}N2O}$ ) to artificial soil mixtures amended with the different organic materials (n=22). The three points for which  $R_{\rm NH2OH\text{-}to\text{-}N2O}$  was determined at pH 3, 4, and 5 without MnO<sub>2</sub> addition were excluded from the simulation.

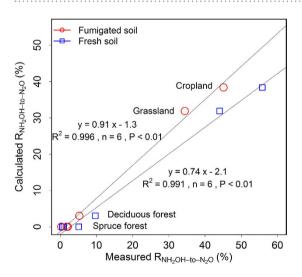


Figure 6. Results of the application of the artificial soil regression model for the calculation of NH<sub>2</sub>OH-to-N<sub>2</sub>O conversion ratios ( $R_{\rm NH2OH-to-N2O}$ ) to six natural fresh and chloroform-fumigated soils as reported in Heil *et al.*<sup>6</sup>.

be confirmed for other soils from different ecosystems, this improved understanding of the controls of  $N_2O$  formation from the reactive nitrification intermediate  $NH_2OH$  in soils can have large implications for developing appropriate management options, such as adding organic amendments with suitable chemical characteristics, for mitigating  $N_2O$  emissions from agricultural land, the largest anthropogenic source of  $N_2O$  to the atmosphere.

### Methods

**Experimental setup.** Two full-factorial artificial soil experiments were conducted. The first experiment comprised three factors (pH, MnO<sub>2</sub> and OM content) and five levels of each factor. The scond experiment comprised also three factors (pH, MnO<sub>2</sub> and OM quality) with five levels of pH and MnO<sub>2</sub>, and four different organic materials at the same concentration level (2.5% w/w on a dry weight basis), but of different quality. Each experiment was conducted in triplicate.

**Preparation of the artificial soil mixtures.** The artificial soil mixtures consisted of 15% (expressed as percentage of dry weight) fine quartz sand (50% of the particles  $0.05-0.2\,\mathrm{mm}$ ), representing the sand fraction, 65% quartz powder ( $0.002-0.063\,\mathrm{mm}$ ), representing the silt fraction, and 20% kaolin clay ( $\leq 0.002\,\mathrm{mm}$ ), representing the clay fraction, mimicking the soil texture of the agricultural Terrestrial Environmental Observatories (TERENO) field site Selhausen<sup>37</sup>. Freeze-dried, finely ground and sieved ( $< 0.75\,\mathrm{mm}$ ) peat moss (Sphagnum magellanicum, collected from Dürres Maar, Eifel, Germany) was amended as SOM to the artificial soil mixtures at levels of 0%, 1%, 2.5%, 5%, 10% dry weight, while the relative amount of sand, clay and silt was reduced according to the amount of peat moss added. The water holding capacity (WHC) was determined for each of the artificial soil mixtures. The WHC increased with increasing organic matter (OM) content, and amounted to 29%, 44%, 55%, 76%, and 132% for the five OM contents, respectively. Each of those artificial soil mixtures was amended with MnO<sub>2</sub> (Merck, Darmstadt, Germany) at five different levels (0%, 0.01%, 0.025%, 0.05%, 0.1% Mn), then the ingredients were thoroughly homogenized.

**Preparation of artificial soil mixtures with different OM qualities.** Organic materials with different C/N ratios (Table 1) were derived from two different plant species, i.e. watermilfoil (*Myriophyllum* spec.) and clover (*Trifolium repens*), and from a cyanobacterium (*Spirulina platensis*). Watermilfoil and clover had been collected previously on the campus of Forschungszentrum Jülich (2004 and 2014, respectively), while the cyanobacterium material had been purchased in 2006 (Concept Vitalprodukte, Schwerte, Nordrhein-Westfalen, Germany). The finely ground, freeze-dried and sieved ( $<0.75\,\mathrm{mm}$ ) organic material was amended to the inorganic quartz-kaolin mixture as described above at a rate of 2.5% dry weight, while the relative amount of sand, clay and silt was reduced accordingly. Also for this experiment, each of the artificial soil mixtures was amended with MnO<sub>2</sub> at five different levels (0%, 0.01%, 0.025%, 0.05%, 0.1% Mn), and again mixed thoroughly to obtain a homogeneous composition.

Addition of  $NH_2OH$  to the artificial soil mixtures and analysis of the  $N_2O$  formed. One gram of each artificial soil mixture was weighed into individual 22-mL gas chromatograph (GC) vials. Subsequently,  $NH_2OH$  in different buffer solutions was added to each vial to obtain a soil water content of 50% WHC, which required addition of varying volumes of buffer solution to the different soil mixtures depending on the OM content, and adaptation of the  $NH_2OH$  concentration of each of the buffer solutions accordingly. The total amount of  $NH_2OH$  added to each of the soil mixtures was always 5 nmol (equivalent to  $70\,\mu g\,N$  per kg dry material). The pH buffer solutions at pH 3, 4, 5 and 6 were prepared with citric acid (0.1 M) and sodium citrate (0.1 M) according to Gomori<sup>38</sup>, whereas the buffer at pH 7 was prepared with tris(hydroxymethyl)aminomethane and maleate (Tris-maleate buffer). The vials were closed immediately after  $NH_2OH$  addition. After 10 hours of incubation, the  $N_2O$  concentration in the headspace of the vials was measured with a GC equipped with an electron capture detector (Clarus 580, PerkinElmer, Rodgau, Germany). Details of the GC setup and analytical conditions have been described previously<sup>11</sup>.

 $\begin{tabular}{ll} \textbf{Calculation of the NH$_2$OH-to-N$_2$O conversion ratio.} & The $NH$_2$OH-to-N$_2$O conversion ratio $(R_{NH2OH-to-N2O}$, moles $N$_2$O-N per mole $NH$_2$OH-N, %) was determined according to the following equation: $(R_{NH2OH-to-N2O}$, moles $N_2$O-N per mole $NH$_2$OH-N, %) and $(R_{NH2OH-to-N2O}$, moles $N_2$O-N per mole $NH_2$OH-N, %) and $(R_{NH2OH-to-N2O}$, moles $N_2$O-N per mole $NH_2$O-N per mole $NH_2$O-N$ 

$$R_{NH_2OH-to-N_2O} = (c_1 - c_0) \cdot V/V_m \cdot 2/n \cdot 100$$
(4)

where  $c_0$  is the background N<sub>2</sub>O mixing ratio in the headspace of the control with the same amount of water instead of NH<sub>2</sub>OH solution (nL L<sup>-1</sup>);  $c_1$  is the N<sub>2</sub>O mixing ratio in the headspace of the sample with NH<sub>2</sub>OH addition (nL L<sup>-1</sup>); the factor 2 represents the molar N ratio of N<sub>2</sub>O and NH<sub>2</sub>OH; V is the volume of the vial headspace (0.022 L);  $V_m$  is the molar volume of N<sub>2</sub>O at standard pressure and room temperature (24.465 L mol<sup>-1</sup>); n is the amount of NH<sub>2</sub>OH added to the sample vials (5 nmol).

Determination of the basic properties of the organic materials. Three replicates of each organic material were analyzed to determine its basic properties. The C and N content of the different organic materials was analyzed by weighing  $200-300\,\mu g$  dry material into tin capsules, followed by combustion at  $1080\,^{\circ}$ C in an elemental analyzer (EuroEA, EuroVector, Milan, Italy) interfaced to an isotope-ratio mass spectrometer (Isoprime, Isoprime Ltd, Stockport, United Kingdom). The C and N content was determined through peak integration of m/z 44 (CO<sub>2</sub>) and 28 (N<sub>2</sub>), respectively, and calibrated against elemental standards.

The elemental composition of the organic materials was analyzed by using inductively coupled plasma optical emission spectrometry (ICP-OES) in the central analytical laboratory (ZEA-3) of Forschungszentrum Jülich. Briefly,  $100\,\mathrm{mg}$  of sample material were mixed with  $3\,\mathrm{mL}$  HNO $_3$  and  $2\,\mathrm{mL}$  H $_2\mathrm{O}_2$ , heated in the microwave at  $800\,\mathrm{W}$  for  $30\,\mathrm{min}$ . The mixtures were subsequently filled up to  $14\,\mathrm{mL}$  and diluted 10-fold with deionized water followed by the ICP-OES measurement.

For the determination of characteristic molecule structures and functional groups of the different organic materials used in the experiments, <sup>13</sup>C and <sup>15</sup>N cross-polarisation magic-angle spinning (CPMAS) nuclear magnetic resonance (NMR) spectra were obtained. <sup>13</sup>C CPMAS spectra were obtained on a 7.05 T Varian INOVA<sup>TM</sup> Unity (Varian Inc., Palo Alto, CA, USA) at a <sup>13</sup>C resonance frequency of 75.4 MHz. <sup>15</sup>N CPMAS spectra were

obtained on a 14.09 T Varian NMR system (Varian Inc., Palo Alto, CA, USA) at a  $^{15}$ N resonance frequency of 60.8 MHz. Samples were packed into 6 mm diameter cylindrical zirconia rotors with Vespel® drive tips and spun at 8000  $\pm$  3 Hz in an HX Apex probe. The spectra were collected with a sweep width of 25 kHz and an acquisition time of 20 ms. In preliminary experiments, the optimal contact time and recycle delay for the cross-polarization experiment were determined. A contact time of 1 ms and a 5 s recycle delay time were used for  $^{13}$ C, whereas a contact time of 1 ms and a 1 s recycle delay time were used for  $^{15}$ N. During cross-polarization the  $^{1}$ H radio frequency (RF) field strength was set to 47 kHz for  $^{13}$ C and to 33.7 kHz for  $^{15}$ N, respectively. The  $^{13}$ C and  $^{15}$ N RF field strength was set to 41 and 41.7 kHz, respectively. An ascending ramp of 15 and 12.2 kHz on the  $^{1}$ H-RF field was used for  $^{13}$ C and  $^{15}$ N during contact time to account for inhomogeneities of the Hartmann-Hahn condition, respectively. Proton decoupling was done using a spinal sequence with a  $^{1}$ H field strength of 50 and 35.6 kHz, a phase of 4.5° and 5.5°, and a pulse length of 12 and 9.5  $\mu$ s, respectively.

The free induction decays (FID) were recorded with VnmrJ (Version 1.1 RevisionD, Varian Inc., Palo Alto, CA, USA) and processed with Mestre-C (Version 4.9.9.9, Mestrelab Research, Santiago de Compostela, Spain). All FIDs were fourier-transformed with an exponential filter function with a line broadening (LB) of 20 to 50 Hz. Baseline correction was done using the manual baseline correction function of Mestre-C.

The  $^{13}$ C chemical shifts are reported relative to tetramethylsilane (=0 ppm) using adamantane as an external reference. The relative intensities of the regions were determined using the integration routine of the MestRe-C software. The  $^{15}$ N chemical shifts are reported relative to ammonium nitrate (NH<sub>4</sub> $^+$ =0 ppm).

**Data analysis.** The homogeneity of variance was tested with the Bartlett test. One-way analysis of variance (one-way ANOVA) of the main controlling factors in the two experiments was performed, followed by a Tukey Honest Significant Difference (HSD) test. A stepwise multiple regression model for the NH<sub>2</sub>OH-to-N<sub>2</sub>O conversion ratio was developed on the basis of the co-variables pH, MnO<sub>2</sub> and SOM content by using the data from the first experiment. In this case, significance was tested with the F test. Linear regression was performed for simulated and measured  $R_{\rm NH2OH-to-N2O}$  in artificial and natural soils described in Heil *et al.*<sup>6</sup> and tested for significance. All analyses were performed with the R software package (version 3.1.0, R Development Core Team, 2013)<sup>40</sup>.

# References

- 1. Sutka, R. L. et al. Distinguishing nitrous oxide production from nitrification and denitrification on the basis of isotopomer abundances. Appl. Environ. Microb. 72, 638–644 (2006).
- 2. Rathnayake, R. M. L. D. et al. Source identification of nitrous oxide on autotrophic partial nitrification in a granular sludge reactor. Water Res. 47, 7078–7086 (2013).
- 3. Stein, L. Y. 6 Surveying  $N_2$ O-Producing Pathways in Bacteria In *Methods in enzymology* (ed Klotz, M. G.) 131–152 (Elsevier, 2011).
- 4. Wunderlin, P. et al. Isotope Signatures of N<sub>2</sub>O in a Mixed Microbial Population System: Constraints on N<sub>2</sub>O Producing Pathways in Wastewater Treatment. *Environ. Sci. Technol.* 47, 1339–1348 (2012).
- 5. Bremner, J. M., Blackmer, A. M. & Waring, S. A. Formation of nitrous oxide and dinitrogen by chemical decomposition of hydroxylamine in soils. *Soil Biol. Biochem.* 12, 263–269 (1980).
- Heil, J., Liu, S., Vereecken, H. & Brüggemann, N. Abiotic nitrous oxide production from hydroxylamine in soils and their dependence on soil properties. Soil Biol. Biochem. 84, 107–115 (2015).
- 7. Lees, H. Hydroxylamine as an Intermediate in Nitrification. Nature 169, 156-157 (1952).
- 8. Vajrala, N. et al. Hydroxylamine as an intermediate in ammonia oxidation by globally abundant marine archaea. *Proc. Natl. Acad. Sci. USA* 110, 1006–1011 (2013).
- 9. Schmidt, I., Look, C., Bock, E. & Jetten, M. S. M. Ammonium and hydroxylamine uptake and accumulation in Nitrosomonas. *Microbiology* 150, 1405–1412 (2004).
- 10. Schmidt, I., van Spanning, R. J. M. & Jetten, M. S. M. Denitrification and ammonia oxidation by Nitrosomonas europaea wild-type, and NirK- and NorB-deficient mutants. *Microbiology* **150**, 4107–4114 (2004).
- 11. Liu, S., Vereecken, H. & Brüggemann, N. A highly sensitive method for the determination of hydroxylamine in soils. *Geoderma* 232–234, 117–122 (2014).
- 12. Bremner, J. Sources of nitrous oxide in soils. Nutr. Cycl. Agroecosys. 49, 7-16 (1997).
- 13. Lehmann, R. G., Cheng, H. H. & Harsh, J. B. Oxidation of Phenolic Acids by Soil Iron and Manganese Oxides. *Soil Sci. Soc. Am. J.* 51, 352–356 (1987).
- 14. Li, C., Zhang, B., Ertunc, T., Schaeffer, A. & Ji, R. Birnessite-Induced Binding of Phenolic Monomers to Soil Humic Substances and Nature of the Bound Residues. *Environ. Sci. Technol.* **46**, 8843–8850 (2012).
- Luther, G. III & Popp, J. Kinetics of the Abiotic Reduction of Polymeric Manganese Dioxide by Nitrite: An Anaerobic Nitrification Reaction. Aquat. Geochem. 8, 15–36 (2002).
- 16. Hulth, S. et al. Nitrogen removal in marine environments: Recent findings and future research challenges. Mar. Chem. 94, 125–145 (2005).
- 17. Hulth, S., Aller, R. C. & Gilbert, F. Coupled anoxic nitrification/manganese reduction in marine sediments. *Geochim. Cosmochim. Acta* 63, 49–66 (1999).
- 18. Dail, D., Davidson, E. & Chorover, J. Rapid abiotic transformation of nitrate in an acid forest soil. *Biogeochemistry* **54**, 131–146 (2001).
- Davidson, E. A., Chorover, J. & Dail, D. B. A mechanism of abiotic immobilization of nitrate in forest ecosystems: the ferrous wheel hypothesis. Global Change Biol. 9, 228–236 (2003).
- 20. Schmidt, B. M. & Matzner, E. Abiotic reaction of nitrite with dissolved organic carbon? Testing the Ferrous Wheel Hypothesis. *Biogeochemistry* **93**, 291–296 (2009).
- El Azhar, S., Verhe, R., Proot, M., Sandra, P. & Verstraete, W. Binding of nitrite-N on polyphenols during nitrification. *Plant Soil* 94, 369–382 (1986).
- 22. Thorn, K. A. & Mikita, M. A. Nitrite Fixation by Humic Substances Nitrogen-15 Nuclear Magnetic Resonance Evidence for Potential Intermediates in Chemodenitrification. Soil Sci. Soc. Am. J. 64, 568–582 (2000).
- 23. Thorn, K. A., Arterburn, J. B. & Mikita, M. A. Nitrogen-15 and carbon-13 NMR investigation of hydroxylamine-derivatized humic substances. *Environ. Sci. Technol.* 26, 107–116 (1992).
- 24. Nelson, D. Transformations of hydroxylamine in soils. Proc. Indiana Acad. Sci. 87, 409-413 (1977).
- Halvorson, J. J. & Gonzalez, J. M. Tannic acid reduces recovery of water-soluble carbon and nitrogen from soil and affects the composition of Bradford-reactive soil protein. Soil Biol. Biochem. 40, 186–197 (2008).
- 26. Olk, D. C. *et al.* Chemical stabilization of soil organic nitrogen by phenolic lignin residues in anaerobic agroecosystems. *Soil Biol. Biochem.* **38**, 3303–3312 (2006).

- 27. Kraus, T. E. C., Zasoski, R. J., Dahlgren, R. A., Horwath, W. R. & Preston, C. M. Carbon and nitrogen dynamics in a forest soil amended with purified tannins from different plant species. *Soil Biol. Biochem.* 36, 309–321 (2004).
- 28. Venterea, R. T., Clough, T. J., Coulter, J. A. & Breuillin-Sessoms, F. Ammonium sorption and ammonia inhibition of nitrite-oxidizing bacteria explain contrasting soil N<sub>2</sub>O production. Sci. Rep. 5 (2015).
- 29. ŠImek, M. & Cooper, J. E. The influence of soil pH on denitrification: progress towards the understanding of this interaction over the last 50 years. *Eur. J. Soil Sci.* **53**, 345–354 (2002).
- Martikainen, P. J., Lehtonen, M., Lång, K., De Boer, W. & Ferm, A. Nitrification and nitrous oxide production potentials in aerobic soil samples from the soil profile of a Finnish coniferous site receiving high ammonium deposition. FEMS Microbiology Ecology 13, 113–121 (1993).
- 31. Liu, B., Mørkved, P. T., Frostegård, Å. & Bakken, L. R. Denitrification gene pools, transcription and kinetics of NO, N<sub>2</sub>O and N<sub>2</sub> production as affected by soil pH. FEMS Microbiology Ecology **72**, 407–417 (2010).
- 32. Samarkin, V. A. et al. Abiotic nitrous oxide emission from the hypersaline Don Juan Pond in Antarctica. Nature Geosci. 3, 341–344 (2010).
- 33. van Cleemput, O. Subsoils: chemo-and biological denitrification, N2O and N2 emissions. Nutr. Cycl. Agroecosys. 52, 187-194 (1998).
- 34. Venterea, R. T. Nitrite-driven nitrous oxide production under aerobic soil conditions: kinetics and biochemical controls. *Global Change Biol.* **13**, 1798–1809 (2007).
- Chao, T. T. Selective Dissolution of Manganese Oxides from Soils and Sediments with Acidified Hydroxylamine Hydrochloride1. Soil Sci. Soc. Am. J. 36, 764–768 (1972).
- 36. Gierer, J. & Söderberg, S. Über die Carbonylgruppen des Lignins. Acta Chem. Scand. 13, 1 (1959).
- 37. Bornemann, L., Herbst, M., Welp, G., Vereecken, H. & Amelung, W. Rock Fragments Control Size and Saturation of Organic Carbon Pools in Agricultural Topsoil. Soil Sci. Soc. Am. J. 75, 1898–1907 (2011).
- 38. Gomori, G. Preparation of buffers for use in enzyme studies In *Handbook of Biochemistry and Molecular Biology 4th edn* (eds Lundblad, R. L. & F. M. MacDonald) 138–146 (CRC Press, 2010).
- Berns, A. E. & Conte, P. Effect of ramp size and sample spinning speed on CPMAS <sup>13</sup>C NMR spectra of soil organic matter. Org. Geochem. 42, 926–935 (2011).
- 40. R Development Core Team. R: A language and environment for statistical computing. R Foundation for Statistical Computing. Vienna, Austria. ISBN 3-900051-07-0, URL http://www.R-project.org/ (2013).

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# **Author Contributions**

N.B. and S.L. conceived the experiments. S.L. conducted the experiments, analysed the data and drafted the manuscript. A.E.B. performed the NMR measurements. D.W. conducted the work related to the exclusion of biological nitrification activity in the artificial soil samples. All authors interpreted the data and contributed to writing the manuscript.

### **Additional Information**

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