



OPEN

Carbon dioxide reduction by photosynthesis undetectable even during phytoplankton blooms in two lakes

Karla Münzner[✉], Silke Langenheder, Gesa A. Weyhenmeyer, Bianka Csitári & Eva S. Lindström

Lakes located in the boreal region are generally supersaturated with carbon dioxide (CO₂), which emerges from inflowing inorganic carbon from the surrounding watershed and from mineralization of allochthonous organic carbon. While these CO₂ sources gained a lot of attention, processes that reduce the amount of CO₂ have been less studied. We therefore examined the CO₂ reduction capacity during times of phytoplankton blooms. We investigated partial pressure of CO₂ (pCO₂) in two lakes at times of blooms dominated by the cyanobacterium *Gloeotrichia echinulata* (Erken, Sweden) or by the nuisance alga *Gonyostomum semen* (Erssjön, Sweden) during two years. Our results showed that pCO₂ and phytoplankton densities remained unrelated in the two lakes even during blooms. We suggest that physical factors, such as wind-induced water column mixing and import of inorganic carbon via inflowing waters suppressed the phytoplankton signal on pCO₂. These results advance our understanding of carbon cycling in lakes and highlight the importance of detailed lake studies for more precise estimates of local, regional and global carbon budgets.

Lakes in the boreal region are generally supersaturated with carbon dioxide (CO₂) and, thus, net sources of CO₂ to the atmosphere^{1,2}. We know that much of the supersaturation emerges from import of inorganic carbon from the surrounding catchment³, including groundwater inflow^{4,5}, and from mineralization of imported organic carbon^{6–9}. On the other hand, processes that reduce the amount of CO₂ in the water, like photosynthesis, are much less studied in this context. However, recent studies suggest that CO₂ uptake by algae can have a stronger influence on CO₂ emissions from lakes, even boreal ones, than previously thought^{10,11}. Therefore, determining the scale of algal CO₂ fixation on lake CO₂ concentrations is a crucial step towards understanding the carbon cycling in those lakes and predicting their contribution to climate change.

We aimed here to get a better understanding of the role of phytoplankton in lake carbon cycling in the boreal region by assessing relations between phytoplankton densities and the partial pressure of CO₂ (pCO₂) in two lakes in which two different types of phytoplankton blooms occurred, i.e. blooms of cyanobacteria and blooms of raphidophytes. Phytoplankton communities in boreal lakes are commonly dominated by cryptophytes, cyanobacteria, chrysophytes, diatoms and raphidophytes^{12,13}. Cyanobacteria blooms mostly occur in eutrophic non-colored lakes¹⁴, while humic lakes are more susceptible to blooms of flagellates, especially from the nuisance alga *Gonyostomum semen*^{15–17}. Both cyanobacteria and *G. semen* blooms are expected to increase as a result of higher temperatures and nutrient concentrations due to increased runoff in a changing climate^{15,18–22}. That makes them perfect model organisms to assess the impact of algal blooms on lake carbon cycling.

To fulfill our aim, we used pCO₂ monitoring data from two Swedish lakes within the Swedish Infrastructure for Ecosystem Science (SITES): a humic lake (Erssjön) in Västergötland (spring – autumn 2017 and 2018), and a mesotrophic non-humic lake (Erken) in Uppland (spring – autumn 2018). The former one is subject to blooms of the raphidophyte *G. semen*, and the latter to blooms of the cyanobacterium *Gloeotrichia echinulata* in summer and diatoms in spring and autumn. We identified phytoplankton blooms based on chlorophyll *a* measurements and *G. semen* and *G. echinulata* densities. This was done with the help of water chemistry and phytoplankton monitoring data, as well as quantitative PCR (qPCR) for *G. semen* specifically. We hypothesized that a) pCO₂ in Erssjön decreases with increasing *G. semen* densities, and b) pCO₂ in Erken decreases with increasing *G. echinulata* densities.

Division of Limnology, Department of Ecology and Genetics, Uppsala University, Uppsala, Sweden. ✉email: karla.munzner@ebc.uu.se

Results

Algal blooms and phytoplankton community composition. We confirmed the presence of the bloom-forming raphidophyte *G. semen* in Erssjön in all DNA samples (analysis: qPCR) taken in 2017 and 2018. Densities of *G. semen* (Fig. 1) varied from being almost absent (110 and 20 cells L⁻¹ in 2017 and 2018 respectively) to bloom conditions (102,760–119,660 cells L⁻¹ in May, June and August 2017, Fig. 1a), based on our phytoplankton bloom identification of cell abundances > 100,000 cells L⁻¹²³. In 2018, no *G. semen* bloom was observed (maximum abundance 20,360 cells L⁻¹ in 2018–05-31, Fig. 1b). In Erken, the bloom-forming cyanobacterium *G. echinulata* was present in phytoplankton samples from 2018–07-17 to 2018–08-14. Chlorophyll *a* concentration in 2018 ranged from 1.9 – 11.95 µg L⁻¹ (Fig. 2) and was only once high enough to be considered as a potential phytoplankton bloom according to EU-monitoring standards²³ (11.95 µg L⁻¹, 2018–08-08). At this date, *G. echinulata* densities and biomass was high enough to be considered a bloom (81,588 cells L⁻¹, biomass: 0.37 mm³ L⁻¹) based on WHO guidelines²⁴ (2000 cells mL⁻¹ or 0.2 mm³ L⁻³).

In Erssjön we found a typical summer phytoplankton community for a humic, mesotrophic lake (based on four sampling occasions in 2018). Dominant phytoplankton groups were green algae, cryptophytes and colonial cyanobacteria (20–45% of the total phytoplankton biomass). Phytoplankton biomass was high (8.77–13.54 mm³ L⁻¹), with *G. semen* making up for 0.5–10% of the total phytoplankton community biomass (Supplementary material, Table S1). The number of species ranged from 24 to 30 and Shannon diversity (H') ranged from 0.5 to 1.1. No single species dominated the phytoplankton community during this time. The C_{phyto}:TOC ratio (= phytoplankton share in total organic carbon) in summer 2018 ranged from 5–9%.

The phytoplankton community in Erken from June to October in 2018 was mainly comprised of diatoms, cryptophytes, dinoflagellates and cyanobacteria. Cryptophytes and dinoflagellates dominated the community in June and July, while cyanobacteria took over in August. Once autumn mixing started in mid-September, diatoms dominated the phytoplankton community up to 79% (2018–10-23, Supplementary material, Table S2). Total phytoplankton biomass (maximum: 2.9 mm³ L⁻¹) was 3–6 times lower compared to Erssjön. The number of species ranged from 20 to 43 and Shannon diversity was (H') ranged from 0.47 to 0.92. The C_{phyto}:TOC ratio remained < 5% through all seasons.

pCO₂ and its relation to phytoplankton blooms. Lake pCO₂ in Erssjön ranged from 334 to 1737 µatm in 2017 (Fig. 1a) and from 527 to 1290 µatm in 2018 (Fig. 1b). Thus, Erssjön was supersaturated in CO₂ (pCO₂ > 400 µatm) during the sampling period, except for 2017–05-16 (Fig. 1a). In 2017, pCO₂ decreased during spring, increased at the beginning of summer and then once at the beginning of autumn. In 2018, there was a different seasonal pattern: pCO₂ remained constant during late spring and early summer, and started to increase from the middle of summer onwards. In early autumn, pCO₂ decreased before increasing again in the middle of autumn. There was no significant relationship between *G. semen* densities and pCO₂ in 2017 (linear regression, $p=0.52$, $R^2=0$, $n=29$). In 2018, pCO₂ did decrease slightly with increasing *G. semen* abundance (linear regression, $R^2=0.15$, $p=0.02$, $n=32$).

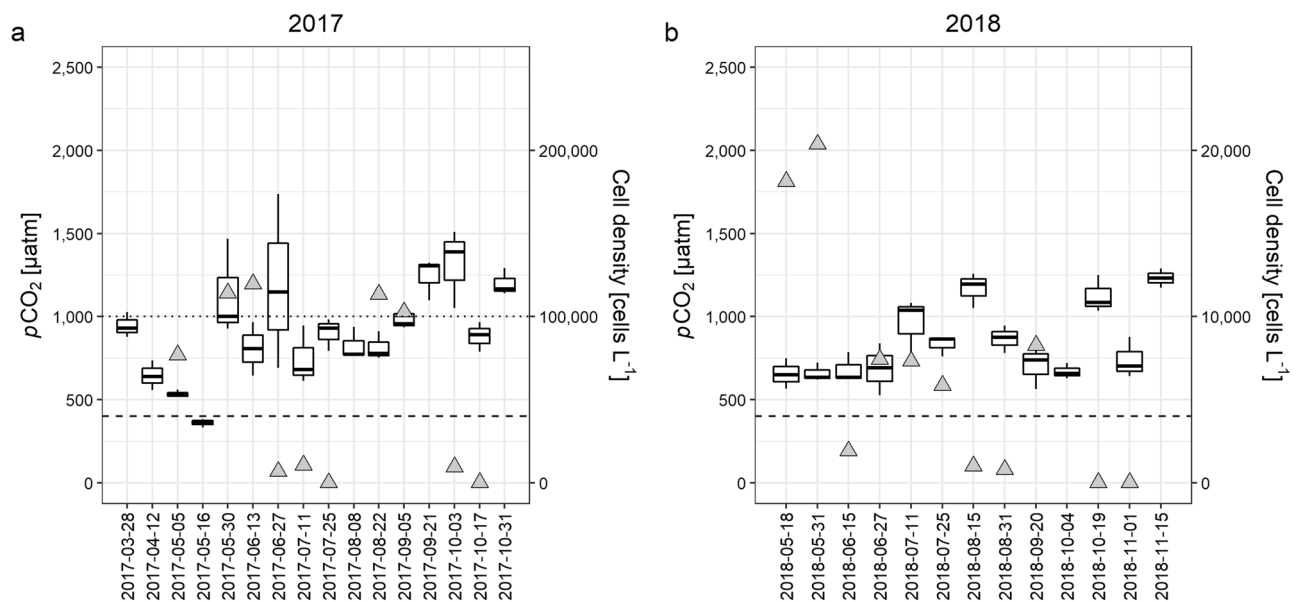


Figure 1. Lake pCO₂ and *G. semen* densities based on qPCR results in Erssjön in 2017 (a) and 2018 (b). For each date, results from 3 sampling locations (pCO₂ measurements, 4 replicates per sampling spot) or one sampling location (*G. semen* abundance) are shown. The boxes symbolize 50% of the data and the line is the median. Boxplots = pCO₂, grey triangles = *G. semen* abundance, dashed line = pCO₂ equilibrium (400 µatm), dotted line = cell number threshold indicating a phytoplankton bloom (100,000 cells L⁻¹). Note the different scales in cell densities in (a) and (b).

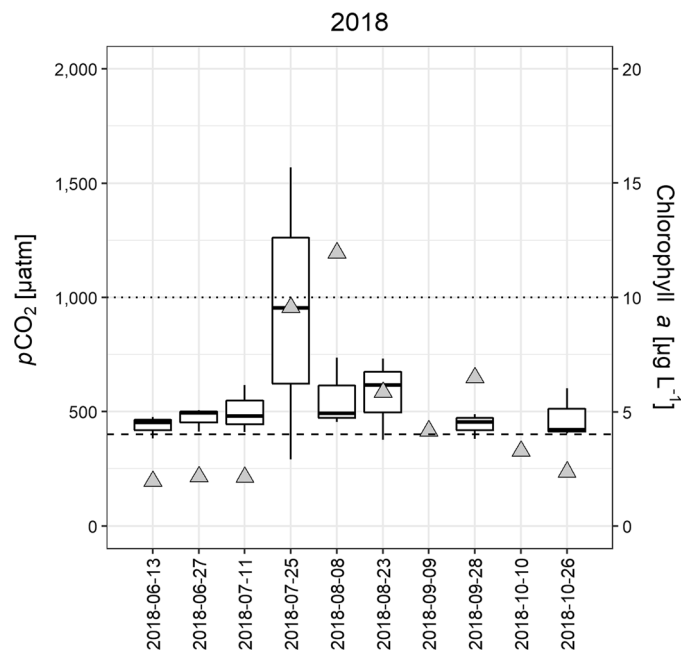


Figure 2. Lake $p\text{CO}_2$ and chlorophyll a concentration in Erken in 2018. For each date in the figure, results from 3 sampling locations ($p\text{CO}_2$ measurements, 4 replicates per sampling spot) or one sampling location (chlorophyll a) are shown. The boxes symbolize 50% of the data and the line is the median. Chlorophyll a concentration was measured from an integrated water sample of either the epilimnion (2018–06–13 until 2018–09–09) or the whole water column after autumn stratification started (2018–09–28 – 2018–10–26). Boxplots = $p\text{CO}_2$, grey triangles = chlorophyll a , dashed line = $p\text{CO}_2$ equilibrium (400 μatm), dotted line = chlorophyll a threshold indicating an algal bloom (10 $\mu\text{g L}^{-1}$).

Lake $p\text{CO}_2$ in Erken ranged from 291 to 1569 μatm in summer and autumn of 2018, with most occasions showing a slight supersaturation (450 to 550 μatm , Fig. 2). The highest $p\text{CO}_2$ was observed in the end of July (2018–07–25), though $p\text{CO}_2$ varied considerably between replicates at those dates. Overall, there was no significant fluctuation in $p\text{CO}_2$ between samplings (Kruskal–Wallis test, chi-squared = 3.51, $df=7$, $p=0.84$). While the second highest chlorophyll a concentration (9.56 $\mu\text{g L}^{-1}$) was observed during the $p\text{CO}_2$ maximum (2018–07–25), there was no significant relationship between chlorophyll a concentration and $p\text{CO}_2$ (linear regression, $R^2=0.1$, $p=0.13$, $n=23$) over the year. Phytoplankton total biomass and $p\text{CO}_2$ (linear regression, $R^2<0.01$, $p=0.7$, $n=23$) and total phytoplankton density and $p\text{CO}_2$ (linear regression, $R^2<0.01$, $p=0.4$, $n=23$) were not significantly correlated with each other either.

Since a relationship between phytoplankton and $p\text{CO}_2$ might be non-linear, we also used an additional approach where we took a closer look at the three periods when algal blooms occurred: May/June and August 2017 (Erssjön) and August 2018 (Erken). Using the Kruskal–Wallis test, we compared $p\text{CO}_2$ during bloom periods (*G. semen* abundance and chlorophyll a maximum respectively) to $p\text{CO}_2$ sampled two weeks before (= pre-bloom) and after (= post-bloom) the blooms were detected. Applying this approach for Erssjön, we found that $p\text{CO}_2$ was similar before, during and after the *G. semen* bloom in May/June 2017 (Kruskal–Wallis test, chi-squared = 7.62, $df=3$, p value = 0.06, Fig. 3a). During the bloom in August 2017, $p\text{CO}_2$ appeared to be significantly higher after the *G. semen* bloom compared to before and during the bloom was observed (Kruskal–Wallis test, chi-squared = 8.74, $df=3$, p value = 0.03, Fig. 3b). However, the post-hoc test revealed that $p\text{CO}_2$ during this bloom was no different from before and after the bloom (Dunn-test with Bonferroni correction for p values, all p values > 0.05, Fig. 3b), indicating that the significant result from the Kruskal–Wallis test was due to a type-I error.

Also, in Erken, $p\text{CO}_2$ remained the same before, during and after the *G. echinulata* bloom in August 2018 (Kruskal–Wallis test, chi-squared = 0.62, $df=2$, p value = 0.72, Fig. 3c). Since chlorophyll a concentration was quite high two weeks before the *G. echinulata* bloom (2018–07–25, “pre-bloom” in Fig. 3c), we also ran a Kruskal–Wallis test including 2018–07–11 as the pre-bloom period, to account for either a longer bloom period or two separate blooms happening on 2018–07–11 and 2018–07–25. Still, neither the pre-bloom nor the post-bloom $p\text{CO}_2$ was higher or lower than $p\text{CO}_2$ during the blooms (Kruskal–Wallis test, chi-squared = 1.05, $df=3$, p value = 0.79).

$p\text{CO}_2$ and its relation physical lake variables. During sampling, water temperatures at 1-m depth in Erssjön ranged from 5 to 20 °C in 2017 and from 5 to 24 °C in 2018, with summer temperatures (June – August) > 16 °C (Supplementary material, Table S3). Air temperatures ranged from –1 to 21 °C (2017) and –1 to 28 °C (2018). Wind speed ranged from 0.2 to 8.4 $\text{m}^1 \text{s}^{-1}$ (2017) and 0.2 to 6.4 $\text{m}^1 \text{s}^{-1}$ (2018), the average wind speeds being 2.3 and 2.4 $\text{m}^1 \text{s}^{-1}$ for 2017 and 2018 respectively. In Erken, water temperatures ranged from 6 to

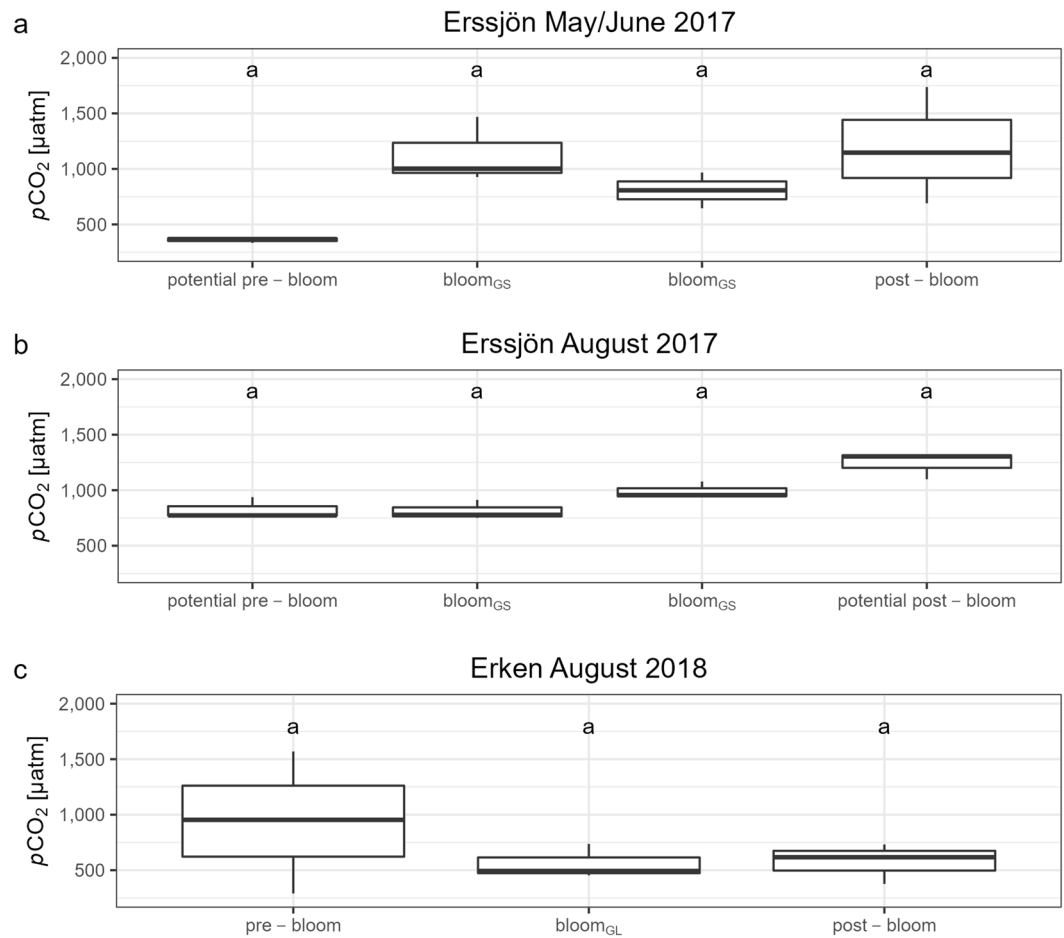


Figure 3. Lake $p\text{CO}_2$ before, during and after observed and potential blooms of *G. semen* (Erssjön) and *G. echinulata* (Erken). In May 2017 in Erssjön, no *G. semen* density data is available for the potential pre- and post-bloom periods. Letters signify the results of the post-hoc Test (Dunn-test, p values adjusted with the Bonferroni method) for the Kruskal–Wallis test results. bloom^{GS} = *G. semen* bloom, bloom^{GL} = *G. echinulata* bloom.

24 °C (18 – 24 °C between June–August), air temperatures from 2 to 27 °C and wind speed from 0.4 to 8.3 $\text{m}^1 \text{s}^{-1}$ (average: 3.7 $\text{m}^1 \text{s}^{-1}$).

We observed no significant correlations between $p\text{CO}_2$ and any of the described above physical variables in the two lakes (Supplementary material, Table S4). In 2018, there was a trend towards decreasing $p\text{CO}_2$ with increasing wind speed in Erssjön, but very little of the overall variation in $p\text{CO}_2$ was explained by this ($R^2 = 0.06$). Furthermore, water column stability (expressed as the Schmidt number) and gas transfer velocity for water–air gas exchange (expressed as k_{600}) were not significantly correlated with $p\text{CO}_2$ either.

Discussion

In this study we investigated whether the occurrence of phytoplankton blooms can suppress $p\text{CO}_2$ in lakes. We studied two different lakes, one in which cyanobacterial (*G. echinulata*) blooms occurred (Erken), and another one in which we observed blooms of the raphidophyte *G. semen* (Erssjön). Contrary to what we expected the phytoplankton signal was not detectable as short-term fluctuations in $p\text{CO}_2$ during those times, which indicates that other factors are more important for $p\text{CO}_2$ in those lakes.

One criterium for phytoplankton to impact lake $p\text{CO}_2$ is a high enough biomass so that photosynthesis dominates over bacterial respiration and mineralization of carbon¹⁰. Available $C_{\text{phyto}}:\text{TOC}$ ratios were > 5% in Erssjön during summer of 2018. Since this is a threshold value for when phytoplankton biomass potentially can have a detectable effect on CO_2 ¹⁰, we would have expected that a phytoplankton bloom in this lake caused a decrease in $p\text{CO}_2$. Since our hypothesis could not be supported, an interesting question in this context is the identity of the bloom forming phytoplankton species, especially the proportion of mixotrophic species. Erssjön's phytoplankton community is typical for a boreal brown water lake with a low pH, with cryptophytes, chrysophytes, euglenids and chlorophytes dominating in the summer of 2018. Some of these are mixotrophic species (*Uroglena* sp., *Dinobryon* sp., *Gymnodium* sp., *Perdinium* sp. and *Ceratium* sp.²⁵, see Supplementary material, Table S5) and could potentially significantly contribute to heterotrophic respiration instead of autotrophic carbon fixation in boreal lakes²⁶, especially in a warming climate²⁷. This could be one contributing factor as to why CO_2 in many

boreal brown water lakes is seemingly not impacted by photosynthetic processes and remains supersaturated throughout the year, even though these lakes demonstrate a high phytoplankton biomass with bloom conditions.

In Erken, the total phytoplankton biomass remained relatively low ($< 2 \text{ mm}^3 \text{ L}^{-1}$) during most of the year, which is below the $C_{\text{phyto}}:\text{TOC}$ ratio of 5% that Engel et al.¹⁰ identified as a threshold for a detectable phytoplankton influence on lake CO_2 . However, a phytoplankton bloom still occurred (*G. echinulata* in August), where also chlorophyll *a* concentrations were higher than $10 \mu\text{g L}^{-1}$ (total phytoplankton biomass: $1.12 \text{ mm}^3 \text{ L}^{-1}$). However, again, these bloom conditions did not leave a detectable signal on the $p\text{CO}_2$. The lack of a direct relation between phytoplankton and $p\text{CO}_2$ in Erken might be a result of measuring phytoplankton and $p\text{CO}_2$ at different sites. Erken is a relatively big lake that is strongly wind exposed with frequent lake water mixing, causing a patchy phytoplankton distribution, especially regarding large floating *G. echinulata* colonies.

Wind exposure is not only affecting the phytoplankton distribution in a lake but also $p\text{CO}_2$. Previous studies have shown that wind-induced upwelling and vertical mixing affect the transport of CO_2 through the water column^{28,29}, as well as the CO_2 flux to the atmosphere³⁰. It is possible that wind suppresses the phytoplankton signal on lake $p\text{CO}_2$, because CO_2 at the lake surface, where phytoplankton are photosynthesizing and, thus, consuming CO_2 , is mixed with CO_2 from waters from below. This may be especially true at Erken, which is highly susceptible to wind mixing, even causing perturbations in the thermal stratification during summer³¹. Furthermore, wind exposure may explain why lake $p\text{CO}_2$ does not increase in autumn after mixing starts in this lake, even though this is typically observed in eutrophic lakes when deep-water, CO_2 water mixes with surface oxygen-rich water³².

However, the *G. echinulata* bloom in Erken occurred mostly at low wind speeds ($1.2\text{--}5.4 \text{ m}^1 \text{ s}^{-1}$, mean: $2.9 \text{ m}^1 \text{ s}^{-1}$) when water mixing is low, as did the *G. semen* blooms in Erssjön ($0.5\text{--}5.1$ and $0.4\text{--}2.8 \text{ m}^1 \text{ s}^{-1}$ in May and June respectively). Consequently, CO_2 import from deeper waters should be low. On the other hand, the gas transfer velocity of CO_2 from water to the atmosphere is lower at low wind speeds, which could lead to higher observed $p\text{CO}_2$ than would be expected during periods of high photosynthetic activity. Thus, the interaction between wind-controlled processes in the water column may be why $p\text{CO}_2$ and physical lake variables do not correlate in our study.

Besides mixing, inorganic carbon import via groundwater and river inflow could have affected $p\text{CO}_2$ in both lakes, which is common for lakes in the studied geographical region³. This could for example explain the increase of $p\text{CO}_2$ in Erssjön in autumn in 2017 and 2018, as this is when autumn mixing starts, precipitation events become more common and water in- and outflow from the lake increase. Furthermore, Erssjön is a rather small lake with a potentially shorter water retention time than Erken (7 years), so hydrological factors contributing to CO_2 supersaturation might be dominating over biological processes.

Even if blooms caused by *G. semen* or other phytoplankton species lead to a decrease in lake $p\text{CO}_2$, it is unclear whether this would have a short-term or long-term effect on lake CO_2 emissions. Autochthonous carbon is preferably used by bacteria and often gets re-mineralized before it can sink to the lake bottom and contribute to the fixed carbon in the sediment^{33,34}. Erssjön's $p\text{CO}_2$ increased in autumn in 2017 and 2018, indicating that phytoplankton biomass may contribute to internal lake carbon respiration and thus CO_2 emissions from the lake when being degraded. In order to understand the ecosystem effects, we would therefore need to monitor *G. semen* blooms and phytoplankton in general more closely, both regarding their biomass and the sedimentation of organic material before, during and after a bloom.

Sampling frequency is also critical for capturing times when *G. semen* migrates vertically through the water column during the day³⁵. During migration, it is easy to underestimate *G. semen* densities by sampling too early or too late during the day. It is also known that the carbon flux has high diel and seasonal variability^{32,36}, with diel variations being especially pronounced during mixing events³². Even though the floating chambers were deployed for around 24 h, deployment and measurement times were not consistent between samplings (ranging from $\sim 8:00$ to $20:00$, see Supplementary material, Table S6). This means that changes in $p\text{CO}_2$ during the day may not always have been captured completely. Thus, it is likely that the $p\text{CO}_2$ measurements at the surface of the lakes caused an under- or overestimation of $p\text{CO}_2$ for our sampling dates due to the time of sampling.

In summary, our results show that neither *G. semen* nor *G. echinulata* blooms were associated with decreases in $p\text{CO}_2$ in two Swedish boreal lakes. It is likely that physical factors, such as wind induced water column mixing and import of inorganic carbon via groundwater inflow and runoff, suppress the phytoplankton signal on $p\text{CO}_2$. A higher sampling frequency is necessary to further study the dynamics between phytoplankton blooms and lake CO_2 concentrations, as phytoplankton blooms can occur rapidly and disappear within days. Further studies should focus on monitoring phytoplankton biomass and $p\text{CO}_2$ at higher frequency to capture diel variations, and include sedimentation rates of autochthonous organic matter and daily CO_2 fluctuations to determine if there is a long-term effect of phytoplankton on lake carbon fluxes. Another interesting aspect to study is the species composition of phytoplankton blooms, since the degree of mixotrophy could affect the relationship between phytoplankton and $p\text{CO}_2$.

Methods

Sampling sites. Sampling was performed in a mesotrophic brown water lake (Erssjön) in the Båveån catchment in Västergötland, Sweden ($58^\circ 22' 16.5'' \text{N}$, $12^\circ 09' 40.3'' \text{E}$) and in a mesotrophic non-humic lake (Erken) in the Broströmmen catchment in Uppland, Sweden ($59^\circ 50' 26.35'' \text{N}$, $18^\circ 37' 39.75'' \text{E}$). Background information about the two lakes are publicly available at the SITES website (<https://www.fieldsites.se/en-GB/sites-thematic-programs/lakes-in-sites-44856978>). Erssjön is situated 73 m above sea level, has an area of 0.061 km^2 and a maximum depth of 4.5–5 m. The surrounding area is mostly forest and some agricultural land. The lake is dimictic, with anoxic periods during summer stratification. Samples were taken from March to October in 2017 (16 CO_2

samplings, 11 DNA samplings) and from May to November in 2018 (13 CO₂ samplings, 11 DNA samplings) every 2–4 weeks.

Erken is 13 m above sea level, has an area of 24 km² and a maximum depth of 21 m. The surrounding area is mostly forest, water and some agricultural land and pasture. The lake is dimictic. Samples were taken from June to October in 2018 (8 CO₂ samplings, every 2–4 weeks). Phytoplankton densities and chlorophyll *a* are monitored through weekly samplings conducted by the Erken Laboratory.

CO₂ measurements. Carbon dioxide concentrations were measured with floating chambers designed according to Bastviken et al.³⁷. In each lake, 4 floating chambers were set along 3 transects from the lake shore towards the center (Fig. 4, check Supplementary material, Table S7 for coordinates). The lake water depths below the floating chambers ranged from 0.3 to 1.9 m (Erssjön) and 0.6–4.7 m (Erken). After a deployment of 18–28 h (Supplementary material, Table S6) a subsample of the headspace was taken with a syringe (3 × 60 mL) and transferred into a gas-tight, sealed glass vial through a needle (0.5 mm; 25 gauge). Samples were measured on a gas chromatograph autosampler. These headspace CO₂ concentrations should be close to equilibrium and pCO₂ was calculated according to Henry's Law. This data is publicly available upon request on the SITES website (www.fieldsites.se).

Phytoplankton bloom identification. Phytoplankton blooms in lake surface waters were identified by analyzing chlorophyll *a* concentration, phytoplankton biomass and phytoplankton densities^{38,39}. Sampling for those variables took place between 9:00 and 12:00 at the two lakes. A chlorophyll *a* threshold of 10 µg L⁻¹ is often used to indicate a phytoplankton bloom^{23,40}. In addition, phytoplankton biomass is often monitored to identify blooms of harmful algal species (HABs), e.g. toxic cyanobacteria^{40,41}. To identify phytoplankton blooms in the two study lakes, we decided to use a chlorophyll *a* threshold of 10 µg L⁻¹ to screen for phytoplankton blooms, as is recommended for moderately deep lakes in Europe²³. For Erken, chlorophyll *a* data was available from the weekly monitoring program upon request (Erken Laboratory, Uppsala University, Sweden, <https://www.ieg.uu.se/erken-laboratory/lake-monitoring-programme/>), and is now also available on the SITES Data portal:

(<https://meta.fieldsites.se/objects/NT59BPT70Ce7h1huzA0xSc1c>). For chlorophyll analysis, water was filtered through a Whatman GF/C filter. Chlorophyll was extracted with 90% acetone and absorbance measured with a spectrophotometer at 750, 664, 647 and 630 nm according to the Swedish standard SS028146.

Since reliable chlorophyll *a* data from Erssjön was not available, we also identified blooms by determining densities of *G. semen* (qPCR) and *G. echinulata* (phytoplankton monitoring data). For densities the following thresholds for a bloom identification were used: 2,000 cells mL⁻¹ or 0.2 mm³ L⁻¹ for *G. echinulata* which is the

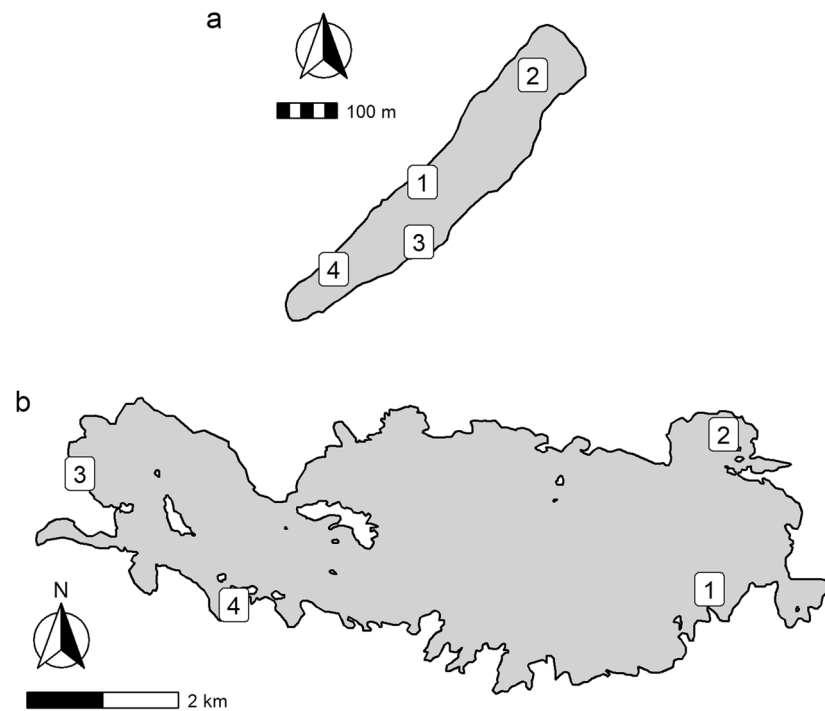


Figure 4. Floating chamber and phytoplankton sampling locations in Erssjön (a) and Erken (b). Numbers in Erssjön (a) refer to 1 = phytoplankton and DNA sampling location, 2 = chambers A1 – A4, 3 = chambers B5 – B8, 4 = chambers C9 – C12, numbers in Erken (b) to 1 = phytoplankton and chlorophyll sampling location, 2 = NJÄ_01 – NJÄ_04, 3 = KHM_01 – KHM_04, 4 = SBA_01 – SBA_04. For individual chamber coordinates, refer to Supplementary material, Table S7. This map was generated with the R software (R 4.1.2, package: ggmap, ggspatial) with shapefiles created in Google Earth.

lowest WHO threshold for dangerous cyanobacteria blooms⁴⁰, based on the fact that these species are potentially toxic^{42,43}, and 100,000 cells L⁻¹ for *G. semen* since this species has an unusual big cell size⁴⁴.

DNA analysis for the identification of *G. semen* densities. In Erssjön, a DNA sample was taken every 2–4 weeks from the Southern side of the mesocosm platform (N 58.371322°, E 12.161074°, see Fig. 4a) during 2017–05–15 – 2017–10–16 and 2018–05–15 – 2018–11–01. An integrated water sample (0–1 m) of 10 L was taken and a subsample of 2 L filtered through a membrane filter (0.2 µm pore size, 47 mm diameter). The filter was then folded, put into a 2 mL cryotube and stored at –80 °C. DNA was extracted with the Qiagen DNeasy PowerSoil Kit according to the kit's instruction manual. The DNA concentrations were quantified using PicoGreen, measured on a Tecan Spark plate reader (Tecan, Switzerland) and found to be high enough for further analysis (~3–20 µg DNA µL⁻¹).

For determining the *G. semen* abundance, we used qPCR with the specific primers developed by Johansson et al.⁴⁵. The DNA was diluted 20× and the master mix done with SYBR Green. The qPCR was run on a CFX96 Touch Real-Time PCR Detection System (Bio-Rad, USA) with the following cycle conditions: 2 min at 90 °C, 40 cycles at 95 °C for 15 s, 40 cycles at 60 °C for 15 s and 40 cycles at 72 °C for 15 s. For each sample, three technical replicates were measured and the average copy number per vegetative cell was used for quantification of *G. semen* cells. To determine the copy number per vegetative cell, we measured the extracted and 10× diluted DNA of three samples of isolated and washed *G. semen* vegetative cells (40 cells per sample) from a *G. semen* culture established from a lake in Northern Uppland (Siggefora, culture SF1A8, established by Ingrid Sassenhagen). More details on this procedure is available in Münzner et al.⁴⁶. The copy number per cell ranged from 171,500 to 304,770, with an average copy number of 243,596 (standard deviation: 54,605) per *G. semen* vegetative cell.

Phytoplankton community composition. To get a better understanding on phytoplankton dynamics we also determined the phytoplankton biomass and community composition. In Erssjön we sampled during four dates in 2018 (July 10, July 24, August 14, September 1). Phytoplankton samples were taken from surface water samples (~0.5 m depth) and a subsample 100 mL was fixed with Lugol. Samples were stored at 4 °C in the dark before analysis. Phytoplankton samples were acclimatized to room temperature (24 h) and were counted with the Utermöhl method with an inverted microscope (Leica DMIL LED fluo) after settling for at least 12 h in a 10 mL chamber. We did a whole chamber count at 40× magnification for phytoplankton > 100 µm (e.g. colonial algae), counts of diagonals (1 or 2) at 200× magnification until a cell count of at least 200 cells (> 20 µm) was reached, and 10 random field counts at 400× magnification. For each identified species, cell length and width were measured for one individual and used for biomass calculations. Phytoplankton shapes and biovolume formulas were determined according to Hillebrand et al.⁴⁷. Phytoplankton diversity (H') and evenness (e_n) were calculated using the Shannon–Wiener Index.

For Erken, we used publicly available plankton data from the Erken Laboratory (Uppsala University, Sweden, <https://www.ieg.uu.se/erken-laboratory/lake-monitoring-programme/>, data available upon request). Phytoplankton samples used in this data set were collected the following way: phytoplankton samples were collected at the deepest point of the lake with an integrated water sampler (“Ramberg tube”) in 2 m intervals. The samples were mixed together according to their proportion to the total lake volume (all samples if the lake was fully mixed, epilimnion and hypolimnion samples mixed separately if the lake was stratified). Then, a 100 mL subsample of the integrated sample (only epilimnion sample if the lake was stratified) was taken, fixated with a few drops of Lugol and stored at 4 °C before further analysis. Phytoplankton were identified and quantified with an inverted light microscope according to the quality standards (SS-EN 15,204:2006) approved by the Swedish Environmental Protection Agency (EPA).

From the phytoplankton data we calculated the C_{phyto}:TOC ratio to identify a previously described threshold when phytoplankton dynamics become detectable in lake waters¹⁰. For C_{phyto} a conversion factor of 0.15 was chosen according to Engel et al.¹⁰. For TOC (total organic carbon) we used data on dissolved organic carbon (DOC). DOC has been shown to stand for 97% of the TOC in boreal lakes⁴⁸, thus we assumed DOC to be equal to TOC in our study lakes. Since carbon monitoring data were not available for Erssjön, we used an average of previously reported values (22.33 mg DOC L⁻¹) by Groeneveld et al.⁴⁹.

Lake physical variables. Publicly available data collected by SITES (SITES Data portal: <https://data.field.sites.se/portal/>) of air temperature, water temperature at 1-m depth, wind direction and wind speed were used. For analyses, average values of those variables over the total deployment time of the chambers at each date were calculated. Schmidt number Sc was calculated according to Wanninkhof et al.⁵⁰, and for piston velocity k₆₀₀, the formula for low wind speeds by Cole and Caraco⁵¹ was used.

Statistics. All data were analyzed with R 4.1.2. Group comparisons of CO₂ measurements within transects with the Kruskal–Wallis test (post-hoc test: Dunn-test with Bonferroni correction for *p* values) showed that *p*CO₂ was significantly higher in the chambers closest to the shore compared to the chambers further away from the shore, both in Erssjön (chi-square = 9.18, df = 3, *p* = 0.03) and Erken (chi-square = 8.43, df = 3, *p* = 0.04). However, even though *p*CO₂ decreased with increasing distance to the shore (linear regression analysis) in Erssjön (*p* < 0.001, R² = 0.02, *n* = 316) and Erken (*p* < 0.05, R² = 0.05, *n* = 104), almost none of the overall variation in *p*CO₂ was explained by that. Following this, we used the average *p*CO₂ of each transect for further analyses (*n* = 3 per lake per sampling).

Shown regressions and correlations are based on linear relationships. Since the distribution of all variables was skewed, they were log-transformed. Changes in *p*CO₂ over sampling dates in each lake for each year were assessed with the Kruskal–Wallis test. The Kruskal–Wallis test was also used for comparisons of *p*CO₂ before,

during and after phytoplankton blooms. For post-hoc analysis of significant results, the Dunn test was used and p values were adjusted with the Bonferroni method (package: FSA).

Data availability

The datasets generated during and/or analysed during the current study are available in the DiVA (Digitala Vetenskapliga Arkivet) portal, [urn:nbn:se:uu:diva-479668].

Received: 22 July 2022; Accepted: 14 August 2023

Published online: 19 August 2023

References

- Cole, J. J., Caraco, N. F., Kling, G. W. & Kratz, T. K. Carbon-dioxide supersaturation in the surface waters of lakes. *Science* **265**, 1568–1570 (1994).
- Raymond, P. A. *et al.* Global carbon dioxide emissions from inland waters. *Nature* **503**, 355–359 (2013).
- Weyhenmeyer, G. A. *et al.* Significant fraction of CO₂ emissions from boreal lakes derived from hydrologic inorganic carbon inputs. *Nat. Geosci.* **8**, 933–936 (2015).
- Einarsdottir, K., Wallin, M. B. & Sobek, S. High terrestrial carbon load via groundwater to a boreal lake dominated by surface water inflow. *J. Geophys. Res. Biogeosci.* **122**, 15–29 (2016).
- Nydahl, A. C., Wallin, M. B., Laudon, H. & Weyhenmeyer, G. A. Groundwater carbon within a boreal catchment: Spatiotemporal variability of a hidden aquatic carbon pool. *J. Geophys. Res. Biogeosci.* **125**, e2019JG005244. <https://doi.org/10.1029/2019JG005244> (2019).
- del Giorgio, P. A., Cole, J. J., Caraco, N. F. & Peters, R. H. Linking planktonic biomass and metabolism to net gas fluxes in Northern temperature lakes. *Ecology* **80**, 1422–1431 (1999).
- Jonsson, A., Karlsson, J. & Jansson, M. Sources of carbon dioxide supersaturation in clearwater and humic lakes in northern Sweden. *Ecosystems* **6**, 224–235 (2003).
- Sobek, S., Algesten, G., Bergström, A.-K., Jansson, M. & Tranvik, L. J. The catchment and climate regulation of pCO₂ in boreal lakes. *Glob. Chang. Biol.* **9**, 630–641 (2003).
- Cole, J. J. *et al.* Plumbing the global carbon cycle: Integrating inland waters into the terrestrial carbon budget. *Ecosystems* **10**, 172–185 (2007).
- Engel, F., Drakare, S. & Weyhenmeyer, G. A. Environmental conditions for phytoplankton influenced carbon dynamics in boreal lakes. *Aquat. Sci.* **81**, 35. <https://doi.org/10.1007/s00027-019-0631-6> (2019).
- Rohrlack, T., Frostad, P., Riise, G. & Hagman, C. H. C. Motile phytoplankton species such as *Gonyostomum semen* can significantly reduce CO₂ emissions from boreal lakes. *Limnologia* **84**, 125810. <https://doi.org/10.1016/j.limno.2020.125810> (2020).
- Lepistö, L. & Rosenström, U. The most typical phytoplankton taxa in four types of boreal lakes. *Hydrobiologia* **369**, 89–97 (1998).
- Maileht, K. *et al.* Water colour, phosphorus and alkalinity are the major determinants of the dominant phytoplankton species in European lakes. *Hydrobiologia* **704**, 115–126 (2013).
- Ptácnik, R. *et al.* Quantitative responses of lake phytoplankton to eutrophication in Northern Europe. *Aquat. Ecol.* **42**, 227–236 (2008).
- Cronberg, G., Lindmark, G. & Björk, S. Mass development of the flagellate *Gonyostomum semen* (Raphidophyta) in Swedish forest lakes—An effect of acidification?. *Hydrobiologia* **161**, 217–236 (1988).
- Lepistö, L., Antikainen, S. & Kivinen, J. The occurrence of *Gonyostomum semen* (Ehr) diesing in finnish lakes. *Hydrobiologia* **273**, 1–8 (1994).
- Hagman, C. H. C. *et al.* The occurrence and spread of *Gonyostomum semen* (Her.) Diesing (Raphidophyceae) in Norwegian lakes. *Hydrobiologia* **744**, 1–14 (2015).
- Paerl, H. W. & Huisman, J. Blooms like it hot. *Science* **320**, 57–58 (2008).
- Jeppesen, E. *et al.* Climate change effects on runoff, catchment phosphorus loading and lake ecological state, and potential adaptations. *J. Environ. Qual.* **38**, 1930–1941 (2009).
- Heilmann, A., Krienitz, L. & Koschel, R. Long-term phytoplankton changes in an artificially divided, top-down manipulated humic lake. *Hydrobiologia* **448**, 83–96 (2001).
- Findlay, D. L., Paterson, M. J., Hendzell, L. L. & Kling, H. J. Factors influencing *Gonyostomum semen* blooms in a small boreal reservoir lake. *Hydrobiologia* **533**, 243–252 (2005).
- Rengefors, K., Weyhenmeyer, G. A. & Bloch, I. Temperature as a driver for the expansion of the microalga *Gonyostomum semen* in Swedish lakes. *Harmful Algae* **18**, 65–73 (2012).
- Poikane, S. *et al.* Defining ecologically relevant water quality targets for lakes in Europe. *J. Appl. Ecol.* **51**, 592–602 (2014).
- World Health Organization (WHO). Cyanobacterial toxins: microcystins. Background document for development of WHO guidelines for drinking-water quality and guidelines for safe recreational water environments. *World Health Organization*. (2020) <https://apps.who.int/iris/handle/10665/338066>.
- Sanders, R. W. & Porter, K. G. Phagotrophic phytoflagellates In: *Advances in Microbial Ecology* (ed. Nelson, K. E.), 167–192 (Springer, 1988).
- Hansson, T. H., Grossart, H.-P., Beisner, B. E., del Giorgio, P. A. & St-Gelais, N. F. Environmental drivers of mixotrophs in boreal lakes. *Limnol. Oceanogr.* **64**, 1688–1705 (2019).
- Wilken, S., Huisman, J., Naus-Wiezer, S. & van Donk, E. Mixotrophic organisms become more heterotrophic with rising temperature. *Ecol. Lett.* **16**, 225–233 (2012).
- MacIntyre, S., Flynn, K. M., Jellison, R. & Romero, J. R. Boundary mixing and nutrient fluxes in Mono lake. *California. Limnol. Oceanogr.* **44**, 512–529 (1999).
- Polsenaere, P. *et al.* Thermal enhancement of gas transfer velocity of CO₂ in an Amazon floodplain lake revealed by eddy covariance measurements. *Geophys. Res. Lett.* **40**, 1734–1740 (2013).
- Czikowsky, M. J., MacIntyre, S., Tedford, E. W., Vidal, J. & Miller, S. D. Effects of wind and buoyancy on carbon dioxide distribution and air-water flux of a stratified temperate lake. *J. Geophys. Res. Biogeosci.* **123**, 2305–2322 (2018).
- Yang, Y., Colom, W., Pierson, D. & Petterson, K. Water column stability and summer phytoplankton dynamics in a temperate lake (Lake Erken, Sweden). *Inland Waters* **6**, 499–508 (2016).
- Rudberg, D. *et al.* Diel variability of CO₂ emissions from northern lakes. *J. Geophys. Res. Biogeosci.* **126**, 2021006246. <https://doi.org/10.1029/2021JG006246> (2021).
- Kritzberg, E. S. *et al.* Autochthonous versus allochthonous carbon sources of bacteria: Results from whole-lake ¹³C addition experiments. *Limnol. Oceanogr.* **49**, 588–596 (2004).
- Guillemette, F., McCallister, S. L. & Del Giorgio, P. A. Selective consumption and metabolic allocation of terrestrial and algal carbon determine allochthony in lake bacteria. *ISME J.* **10**, 1373–1382 (2015).

35. Rohrlack, T. The diel vertical migration of the nuisance alga *Gonyostomum semen* is controlled by temperature and by a circadian clock. *Limnologia* **80**, 125746. <https://doi.org/10.1016/j.limno.2019.125746> (2020).
36. Spafford, L. & Risk, D. Spatiotemporal variability in lake-atmosphere net CO₂ exchange in the littoral zone of an oligotrophic lake. *J. Geophys. Res. Biogeosci.* **123**, 1260–1276 (2018).
37. Bastviken, D., Sundgren, I., Natchimuthu, S., Reyier, H. & Gålfalk, M. Technical note: cost-efficient approaches to measure carbon dioxide (CO₂) fluxes and concentrations in terrestrial and aquatic environments using mini loggers. *Biogeosciences* **12**, 3849–3859 (2015).
38. Armon, R. H. & Starosvetsky, J. Algal bloom indicators in *Environmental indicators* (eds. Armon, R. H. & Hänninen, O.), 635–640 (Springer, 2015).
39. Zhang, F., Hu, C., Shum, C. K., Liang, S. & Lee, J. Satellite remote sensing of drinking water intakes in Lake Erie for cyanobacteria populations using two MODIS-based indicators as a potential tool for toxin tracking. *Front. Mar. Sci.* **4**, 124. <https://doi.org/10.3389/fmars.2017.00124> (2017).
40. Binding, C. E., Greenberg, T. A., McCullough, G., Watson, S. B. & Page, E. An analysis of satellite-derived chlorophyll and algal bloom indices on Lake Winnipeg. *J. Great Lakes Res.* **44**, 436–446 (2018).
41. Sellner, K. G., Doucette, G. J. & Kirkpatrick, G. J. Harmful algal blooms: causes, impacts and detection. *J. Ind. Microbiol. Biotechnol.* **30**, 383–406 (2003).
42. Carey, C. C., Haney, J. F. & Cottingham, K. L. First report of microcystin-LR in the cyanobacterium *Gloeotrichia echinulata*. *Environ. Toxicol.* **22**, 337–339 (2007).
43. Carey, C. C. *et al.* Occurrence and toxicity of the cyanobacterium *Gloeotrichia echinulata* in low-nutrient lakes in the northeastern United States. *Aquat. Ecol.* **46**, 395–409 (2012).
44. Carletti, A. & Heiskanen, A.-S. Water framework directive intercalibration technical report – part 3: coastal and transitional waters. <https://publications.jrc.ec.europa.eu/repository/handle/JRC51341> (2009).
45. Johansson, K. S. L., Lühhig, K., Klaminder, J. & Rengefors, K. Development of a quantitative PCR method to explore the historical occurrence of a nuisance microalga under expansion. *Harmful Algae* **56**, 67–76 (2016).
46. Münzner, K., Gauthier, V. A., Sassenhagen, I., Tranvik, L. J. & Lindström, E. S. (2023) Effects of a dominant algal species (*Gonyostomum semen*) on the carbon balance of brown water lakes. *In preparation*.
47. Hillebrand, H., Dürselen, C.-D., Kirschtel, D., Pollinger, U. & Zohary, T. Biovolume calculation for pelagic and benthic microalgae. *J. Phycol.* **35**, 403–424 (1999).
48. von Wachenfeldt, E. & Tranvik, L. J. Sedimentation in boreal lakes—The role of flocculation of allochthonous dissolved organic matter in the water column. *Ecosystems* **11**, 803–814 (2008).
49. Groeneveld, M., Tranvik, L. J. & Koehler, B. Photochemical mineralisation in a humic boreal lake: Temporal variability and contribution to carbon dioxide production. *Biogeosci. Discuss.* **12**, 17125–17152 (2016).
50. Wanninkhof, R. Relationship between wind speed and gas exchange over the ocean. *J. Geophys. Res.* **97**(C5), 7373–7382 (1992).
51. Cole, J. J. & Caraco, N. F. Atmospheric exchange of carbon dioxide in a low-wind oligotrophic lake measured by the addition of SF₆. *Limnol. Oceanogr.* **43**(4), 647–656 (1998).

Acknowledgements

This study has been made possible by the Swedish Infrastructure for Ecosystem Science (SITES), in this case at the Skogaryd Research Catchment and the Erken Laboratory. SITES receives funding through the Swedish Research Council under the grant no 2017-00635. We want to especially thank David Bastviken from the Department of Thematic Studies (TEMA) at Linköping University for sharing the SITES greenhouse gas data and his advice on how to use it. Many thanks also go to Pia Larsson at the Erken Laboratory for providing the phytoplankton community data for Erken, and Stina Drakare from the Section for Ecology and Biodiversity at the Department of Aquatic Sciences and Assessment at the Swedish University of Agricultural Sciences (SLU) for her help with phytoplankton species identification. Thank you to Ingrid Sassenhagen for providing samples for determining the copy number of vegetative *G. semen* cells, as well as her advice on how to interpret the qPCR results. This study was financed by grants from Stiftelsen Oscar och Lili Lamms minne, the Åforsk foundation, and Carl Tryggers Stiftelse för Vetenskaplig forskning to ESL, and the Malméns Stiftelse for KM. GAW received financial support from the Swedish Research Council (Grant No. 2020-03222) and the Swedish Research Council for Environment, Agricultural Sciences and Spatial Planning (FORMAS; Grant No. 2020-01091).

Author contributions

S.L., K.M., G.W. and E.S.L. designed the study. K.M. and B.C. conducted the laboratory analyses. K.M. analyzed the data and prepared the figures. K.M. and S.L. wrote the initial manuscript draft. All authors reviewed the manuscript. E.S.L., S.L. and G.W. acted as supervisors. E.S.L. and S.L. provided the resources for the study.

Funding

Open access funding provided by Uppsala University.

Competing interests

The authors declare no competing interests.

Additional information

Supplementary Information The online version contains supplementary material available at <https://doi.org/10.1038/s41598-023-40596-6>.

Correspondence and requests for materials should be addressed to K.M.

Reprints and permissions information is available at www.nature.com/reprints.

Publisher's note Springer Nature remains neutral with regard to jurisdictional claims in published maps and institutional affiliations.



Open Access This article is licensed under a Creative Commons Attribution 4.0 International License, which permits use, sharing, adaptation, distribution and reproduction in any medium or format, as long as you give appropriate credit to the original author(s) and the source, provide a link to the Creative Commons licence, and indicate if changes were made. The images or other third party material in this article are included in the article's Creative Commons licence, unless indicated otherwise in a credit line to the material. If material is not included in the article's Creative Commons licence and your intended use is not permitted by statutory regulation or exceeds the permitted use, you will need to obtain permission directly from the copyright holder. To view a copy of this licence, visit <http://creativecommons.org/licenses/by/4.0/>.

© The Author(s) 2023