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Modelling of the tsunami from the December 22, 2018 lateral collapse of Anak Krakatau volcano in the Sunda Straits, Indonesia

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On Dec. 22, 2018, at approximately 20:55–57 local time, Anak Krakatau volcano, located in the Sunda Straits of Indonesia, experienced a major lateral collapse during a period of eruptive activity that began in June. The collapse discharged volcanoclastic material into the 250 m deep caldera southwest of the volcano, which generated a tsunami with runups of up to 13 m on the adjacent coasts of Sumatra and Java. The tsunami caused at least 437 fatalities, the greatest number from a volcanically-induced tsunami since the catastrophic explosive eruption of Krakatau in 1883 and the sector collapse of Ritter Island in 1888. For the first time in over 100 years, the 2018 Anak Krakatau event provides an opportunity to study a major volcanically-generated tsunami that caused widespread loss of life and significant damage. Here, we present numerical simulations of the tsunami, with state-of-the-art numerical models, based on a combined landslide-source and bathymetric dataset. We constrain the geometry and magnitude of the landslide source through analyses of pre- and post-event satellite images and aerial photography, which demonstrate that the primary landslide scar bisected the Anak Krakatau volcano, cutting behind the central vent and removing 50% of its subaerial extent. Estimated submarine collapse geometries result in a primary landslide volume range of 0.22–0.30 km³, which is used to initialize a tsunami generation and propagation model with two different landslide rheologies (granular and fluid). Observations of a single tsunami, with no subsequent waves, are consistent with our interpretation of landslide failure in a rapid, single phase of movement rather than a more piecemeal process, generating a tsunami which reached nearby coastlines within ~30 minutes. Both modelled rheologies successfully reproduce observed tsunami characteristics from post-event field survey results, tide gauge records, and eyewitness reports, suggesting our estimated landslide volume range is appropriate. This event highlights the significant hazard posed by relatively small-scale lateral volcanic collapses, which can occur *en-masse*, without any precursory signals, and are an efficient and unpredictable tsunami source. Our successful simulations demonstrate that current numerical models can accurately forecast tsunami hazards from these events. In cases such as Anak Krakatau's, the absence of precursory warning signals together with the short travel time following tsunami initiation present a major challenge for mitigating tsunami coastal impact.

In the evening of December 22, 2018, a lateral collapse on the southwest flank of Anak Krakatau (AK) volcano, Indonesia, generated a tsunami along adjacent coastlines in which at least 437 people died¹. This was the most damaging volcanically-generated tsunami since the 1883 eruption of Krakatau, which killed 36,000 people^{2,3}, and the lateral-collapse generated tsunami at Ritter Island, Papua New Guinea, in 1888⁴. Over the past 20 years,

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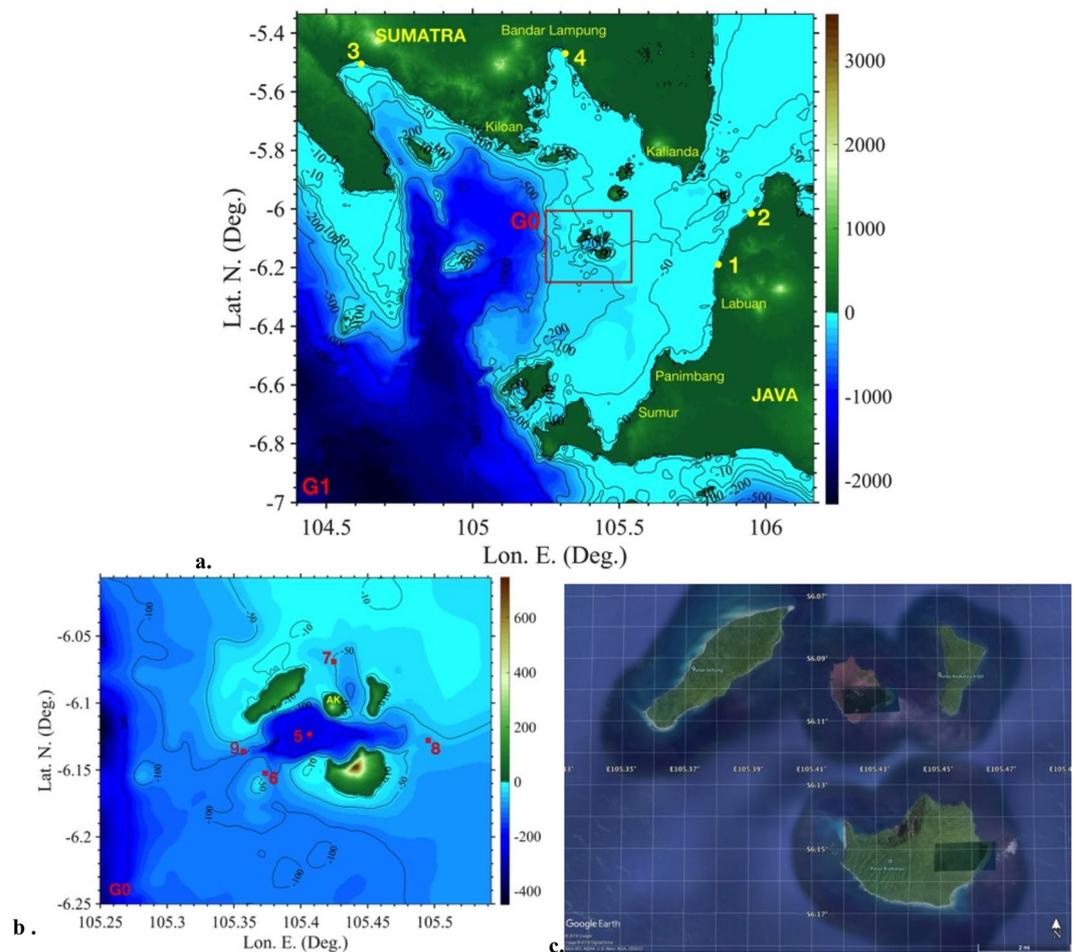


Figure 1. Sunda Straits area, for AK collapse and tsunami simulations. (a,b) Bathymetry/topography (colour scales and contours in meter; Mean Sea Level bathymetry/topography dataset obtained from¹², +1.5 m for tide), tide (1–4) and wave gauges (5–9) (Table 1): (a) FUNWAVE G1 (100 m horizontal resolution grid); (b) NHWAVE G0 grid (90 m horizontal resolution and 5 vertical σ -layers), with footprint marked by red box in (a). (c) Pre-collapse image showing AK and Sertung (W), Panjang/Kitjil (E), and Rakata/Krakatau. Maps in (a,b) were produced by the authors based on the bathymetry/topography dataset, using MATLAB version 2017b (c), map data is from: GOOGLE EARTH, 2019 DigitalGlobe (Data: SIO, NOAA, U.S. Navy, NGA, GEBCO).

catastrophic tsunamis in Papua New Guinea (1998), the Indian Ocean (2004), and Japan (2011) have driven major advances in understanding of earthquakes and submarine landslides as tsunami sources, from developments in constraining source mechanisms, their geographical distribution, and tsunami numerical modelling capability. Tsunamis from volcanic sources, including both eruptions and landslides, have resulted in significant losses of life and property^{5,6}, accounting for 20% of all volcanic fatalities over the past 400 years⁷, but there are no direct observations of a large-scale volcanic tsunami since that at Ritter Island⁸. Tsunamis from lateral collapses have been widely modelled, including very large landslide scenarios on ocean islands e.g.,^{9–11}, as well as smaller scale landslides more comparable to the AK collapse^{12–14}. However, because of the paucity of historical examples the results of these models are not fully validated, and aspects of both landslide source mechanisms e.g.,^{15–17} and tsunami behaviour e.g.,^{10,11} remain poorly understood and challenging to model.

Anak Krakatau (Fig. 1), a small composite volcanic cone that developed within the 250 m deep caldera of Krakatau, emerged subaerially in 1928^{18–20}. It is situated on the northeast margin of the caldera wall and is aligned with the feeder vents of the 1883 Krakatau eruption². During the past 90 years, it has grown from a submarine volcano to a subaerial edifice with a pre-collapse height of about 335 m. Retreats of the coastline to the SW by several hundred metres in 1934, 1935 and 1950²⁰ align with the 2018 collapse and imply instability of the edifice on the SW side, as a result of its position on the submerged scarp of the 1883 caldera. Early activity at Anak Krakatau was dominated by phreatomagmatic explosions. The first lava flows were erupted between 1960 and 1963 as the vent site became fully subaerial²¹ and numerous subsequent eruptions have mostly featured expansion of the island by the growth of lava deltas and a steep sided central pyroclastic cone²². A small tsunami ~2 m high was recorded on Rakata (the southernmost and largest island of the Krakatau archipelago) during a subaerial eruption in 1981, and inferred to originate from a small flank landslide, highlighting the potential instability of the southwest flank of the volcano¹⁸. Apart from this event, no other tsunamis from the volcano have been reported. The most recent

Tide/wave Gauges (Fig. 1a,b)	Lon E. (Deg.)	Lat N. (Deg.)	Measured Arrival (s)	Simulated Arrival (s)	Depth (m)
1. TG: Serang, Jambu	105° 50' 27.6"	−6° 11' 21.5"	1920	1870	12.1
2. TG: Ciwandan	105° 57' 10.8"	−6° 01' 02.5"	2570	2500	3.7
3. TG: Kota-Angung	104° 37' 08.5"	−5° 30' 01.2"	2340	2240	3.7
4. TG: Panjang	105° 19' 06.1"	−5° 28' 08.7"	3420	3390	3.9
5. WG	105.4066°	−6.1234°	N/A	28	256.9
6. WG	105.3733°	−6.1524°	N/A	130	67.4
7. WG	105.4246°	−6.0691°	N/A	70	49.8
8. WG	105.4954°	−6.1279°	N/A	205	63.3
9. WG	105.3571°	−6.1361°	N/A	145	95.5

Table 1. Lat-Lon/depth and measured⁴⁵/simulated (0.27 km³ volume; granular slide) arrival times at Tide Gauges 1–4 (TG; Figs 1a, 5a–d; see Supplementary #3), and location/depth at numerical Wave Gauges 5–9 (WG; Fig. 1b). In simulations, AK collapse is assumed to take place at 20:57' local time (UTC + 7). Listed depth is that in model grid G1 assuming a +1.5 m tide elevation. Difference in arrival time at tide gauges are 30–100 s, compared to the 1 min data interval. Note, as shown in Fig. 5a–d, the simulated arrival time (first crest) is closely identical for both rheologies (viscous or granular collapse); hence a single number is listed in the table. N/A: Not Applicable.

period of volcanic activity of AK started in June 2018 and continued into December, producing Strombolian explosions, lava flows, and ash plumes reaching altitudes of up to 5 km²³. On Dec. 22, 2018, a major lateral collapse occurred on AK's southwest flank, which discharged volcanoclastic material into the sea and triggered a destructive tsunami^{24,25}. Based on seismic records²⁶, eyewitness reports e.g.²⁴, and the agreement of modelled waves with tsunami arrival times at tide gauges (Ina-COAP, 2019; see below; Fig. 1, Table 1), the collapse is estimated to have taken place at 20:55'–57' (UTC + 7). Within 30 minutes, the tsunami flooded the coasts of west Java and southeast Sumatra, causing up to 13 m runups. 13,000 people were injured, 33,000 displaced, and thousands of buildings destroyed^{1,27,28}. The tsunami struck near high tide (about 1.5 m over Mean Sea Level (MSL)) thereby increasing its impact.

In this paper, we use high-resolution satellite imagery and aerial photography to develop a model of the AK lateral collapse scenario as a basis for numerical tsunami simulation. Tsunami simulation results are tested against time series of sea surface elevation recorded at tide-gauge, field observations of tsunami flow depth and inundation along the coasts of Java and Sumatra, and eyewitness accounts. Here, for the first time since 1883, we have an opportunity to test state-of-the-art tsunami modelling methodologies against direct observations constraining both volcanic-tsunami source parameters and observations of the generated waves. With remarkable prescience, a tsunami from a hypothetical AK collapse was modelled by¹²; the collapse in their model was to the SW, with a volume of 0.28 km³, and they predicted tsunami wave heights and arrival times along surrounding coastlines. Our work provides an opportunity to test the validity of this hypothetical scenario from an actual event. The results presented here constrain the style of the AK lateral collapse, providing an analogy for future events at other volcanic islands, and test current landslide-tsunami models, thus forming an important contribution towards assessing tsunami hazard from similar future events and developing improved volcanic-tsunami mitigation strategies.

Results

The 22nd December AK collapse volume and geometry. Satellite photographs and radar images provide the key evidence from which we interpret the AK 2018 collapse geometry and volume. Oblique aerial photographs from December 23rd are particularly important (Supplementary #1)^{29,30}, because they unambiguously identify breaks in the pre-collapse coastline that mark the edge of the failure scarp. The photographs show a sharp, steep-sided cut of the coastal lava deltas to the NW and SE of the pre-collapse vent site. These breaks in the coast constrain the margins of the primary collapse scar on the Sentinel-1A satellite radar images from December 23rd (Western Indonesian Time) and ALOS-2 images on December 24th^{31,32}, which confirm the form of the landslide headwall and provide critical observations of the post-collapse coastline. Our interpretations of these radar images show that the opening angle of the headwall scar was very wide, defining a broadly linear collapse that cut behind the vent. Based on the coastlines defined in Fig. 2, the collapse reduced the subaerial area of the island by 49%. Although the photographs and radar images provide evidence on the subaerial part of the collapse, to produce such a dramatic change in the volume and shape of the island, the collapse plane evidently extended below sea level. This resulted in the submergence of the vent site, as can be seen from the intense Surtseyan explosive eruptions observed on 23rd December^{29,30}. In addition to the observations of the subaerial headwall shape from the satellite data, the Surtseyan explosions provide an approximate constraint on the depth of the landslide basal plane at the vent site. Historical observations suggest Surtseyan activity is characteristic of vent depths ≤ 50 m^{33–35}. To estimate the geometry of the landslide basal plane (and hence the collapse volume), we have combined our constraints on the subaerial headwall with an assumed depth of 25 m at the vent site. We have taken a more conservative value than the 50 m depth noted above, based on the observation that debris rapidly accumulated around the vent site in the days following the collapse, reconstructing the subaerial island in this region. Finally, we have restricted the extent of the basal plane to lie within the submarine bulge of Anak Krakatau evident in bathymetry³⁶ (i.e., we assume that the collapse was confined to the AK edifice, rather than cutting into older rocks

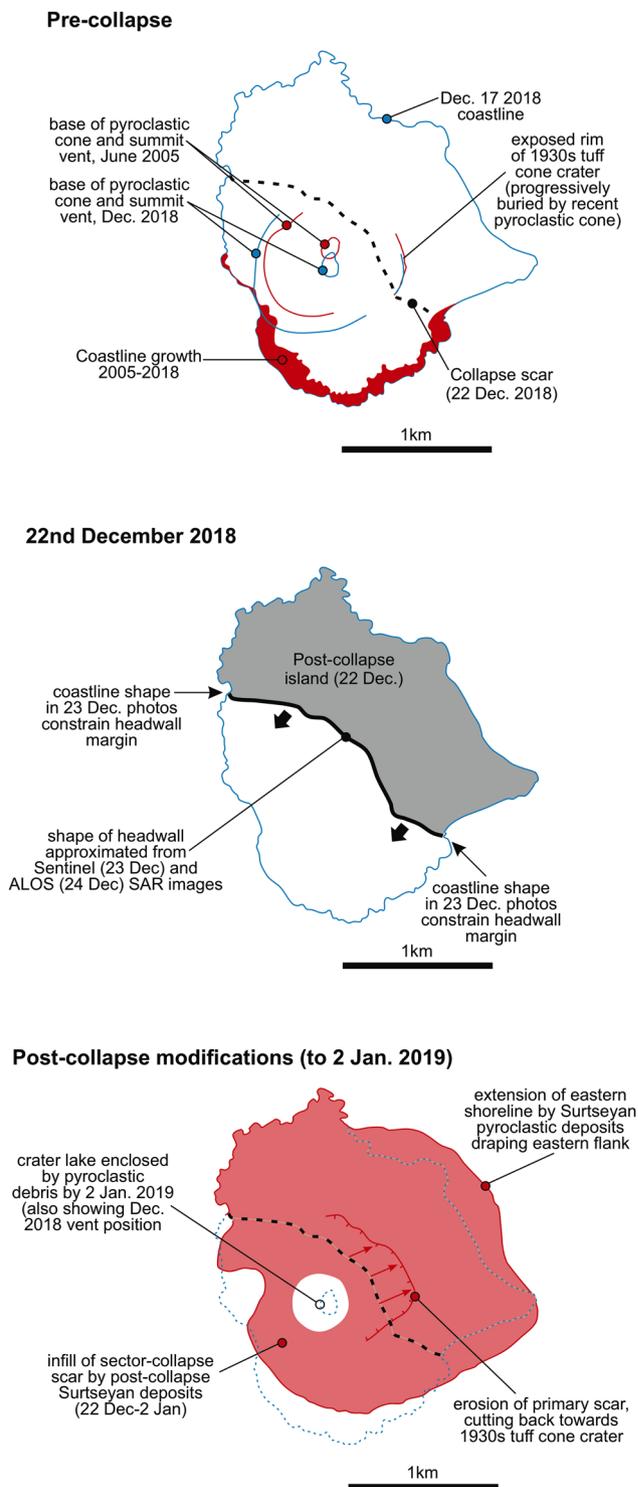


Figure 2. Morphological changes at Anak Krakatau spanning the December 22nd 2018 lateral collapse. Features are interpreted from various satellite images^{29–32}.

of the caldera margin). These assumptions result in a concave and gently dipping form to the basal plane (Fig. 3), typical of subaerial lateral collapse scars^{36,37}. Although the submarine form of the landslide scar is uncertain, a substantially smaller collapse than predicted by the above assumptions would require a sub-horizontal submerged basal plane, which is hard to reconcile with the wide collapse opening angle. We also have no evidence to infer a substantially deeper collapse scar, particularly given the rapid subaerial emergence of post-collapse debris around the vent site. We therefore consider our approximation to be appropriate given an absence of additional

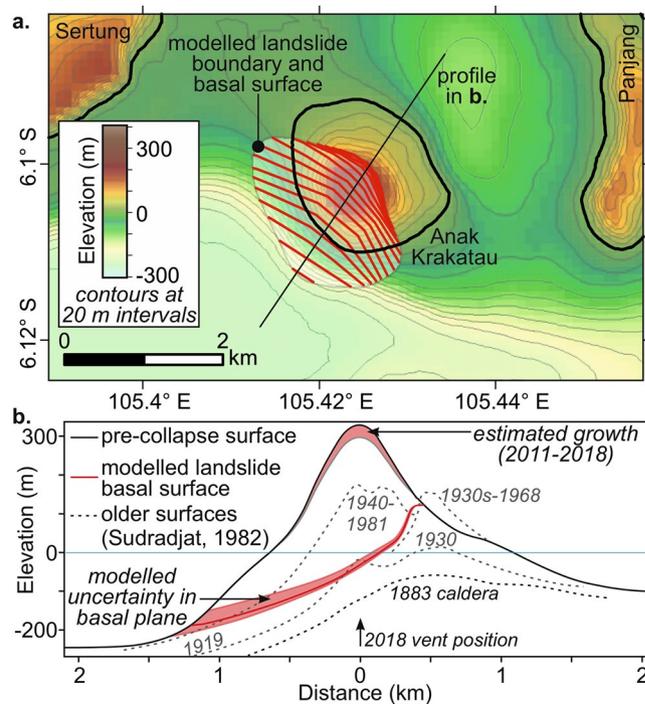


Figure 3. Geometry of modelled AK collapse. **(a)** Pre-collapse bathymetry and topography (colour scale and contours) of parts of Sertung and Panjang, and AK with superimposed geometry of collapsed volume (perimeter and contour levels at 20 m intervals); **(b)** Transect in SW direction (marked by black line in **a**) showing AK's historical growth, the failure surface of the modelled most likely 0.27 km³ volume scenario (consistent with Fig. 2), with associated uncertainty (two additional volumes are modelled for the upper and lower failure surfaces, with a 0.22 (−20%) and 0.30 km³ (+10%) volume, respectively). Maps and surfaces were created by the authors with ArcMap10.2, using the dataset obtained from¹² and the interpretation in Fig. 2.

post-collapse bathymetric or geophysical constraints on the basal plane (noting that this plane is already likely to have been partially buried).

Because the earliest post-collapse satellite images and aerial photographs are from ~12 hours after the event, they provide no evidence on whether the collapse was a single event or occurred in stages. However, the tsunami observations, primarily the single wave train reported and absence of observed subsequent tsunamis, suggest one main phase of failure. This provides support for our numerical tsunami modelling based on a single collapse and using the headwall geometry derived from the 23rd December images.

Landslide tsunami generation and impact modelling in the near-field. Bathymetric and topographic data at an approximate 100 m resolution were obtained from the work of¹², which was interpolated to set-up model grids (90 and 100 m resolutions; Fig. 1a,b and see details below). Comparing Fig. 1b to the pre-collapse aerial image of Fig. 1c shows that AK and the coastline of the three surrounding islands of Sertung, Panjang, and Rakata are well resolved in our grids. A simplified landslide headwall shape and basal plane geometry was defined on this grid based on the constraints outlined above. The likeliest landslide basal plane defined by these constraints is shown in Fig. 3 (solid red line), together with our estimate of the associated uncertainty (shaded red). The modelled collapse volume based on our best-estimated failure surface is calculated at 0.27 km³. To investigate the effects of the failure surface uncertainty on the collapse volume, we also considered and simulated collapse volumes corresponding to the upper and lower bounds of the shaded uncertainty area in Fig. 3, which yielded 10% larger (0.30 km³) and 20% smaller (0.22 km³) volumes. Our best estimated collapse volume is similar to that assumed by¹², albeit for a different failure plane. The three failure surfaces and resulting collapse geometry and volumes were used to initialize simulations with a three-dimensional (3D) slide and hydrodynamic model of tsunami generation and propagation (“Non-Hydrostatic WAVE model” NHWAVE^{38–40}; see methods), as detailed below. For each of the 3 selected collapse volumes and corresponding geometry, the 3D model was used in the near-field (Fig. 1b; grid G0, 90 m horizontal resolution with 5 boundary fitted vertical layers) to simulate the lateral collapse landslide and the corresponding tsunami generation.

As the actual collapse mechanism is unknown, for each collapse volume, simulations were performed assuming two alternative rheologies (granular material and dense viscous fluid), which were available in this model (animations of results are provided in Supplementary #2 for the 0.27 km³ volume scenario). Using these two rheologies allowed assessment, to some extent, of the effect on tsunami simulations of the epistemic uncertainty associated with different model physics, while the range of simulated collapse volumes allowed assessment of the effect on tsunami simulations of the aleatory uncertainty associated with selecting a failure surface and resulting collapse volume.

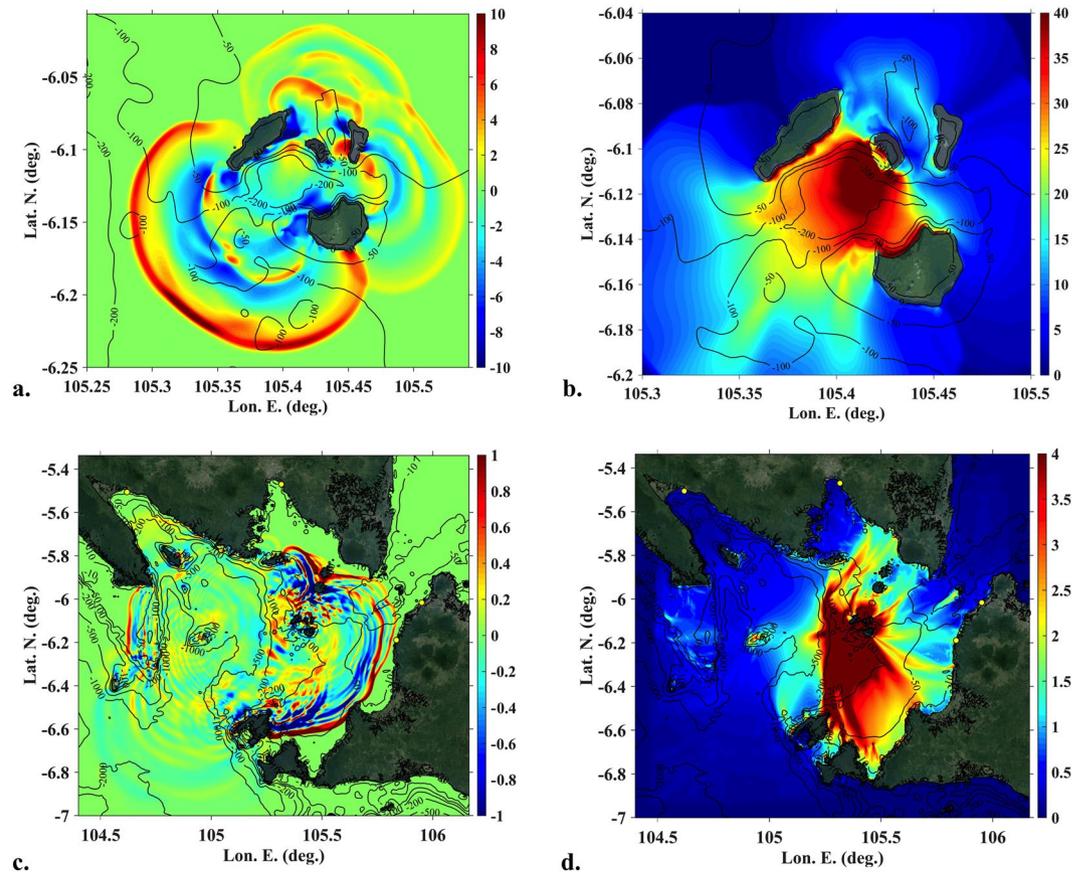


Figure 4. Tsunami surface elevations for AK collapse. **(a,b)** Instantaneous and maximum surface elevations (colour scale in meter), respectively, simulated with NHWAVE in grid G0, for the 0.27 km^3 granular AK collapse at and up to 410 s; **(c)** Instantaneous surface elevation simulated with FUNWAVE-TVD at 1810 s in grid G1; **(d)** Envelope of NHWAVE/FUNWAVE-TVD SE up to 7,610 s. Reference level is MSL + 1.5 m (tide elevation in Marina Jambu⁴⁵ at the time of tsunami arrival). Maps in **(a–d)** were produced by the authors based on the bathymetry/topography dataset, using MATLAB version 2017b; topography from GOOGLE EARTH georeferenced satellite images was embedded in these maps using an API key: GOOGLE EARTH, 2019 DigitalGlobe.

For each rheology, animations of model results show the 3D slide motion with or without the corresponding wave generation, which allows for easier identification of the slide in the latter case (Supplementary #2). For both rheologies, results show a rapid collapse of AK's cone, causing a thick flow of material to the SW, which gradually fills the bottom of the caldera, together with the lateral runout of a thinner layer of slide material. With superimposed wave generation, animations show that a leading elevation wave nearly 50 m high is first generated near AK, that travels in a dominant SW direction, followed by a deeper negative elevation wave at the location formerly occupied by AK's cone. Whereas the leading elevation wave propagates south-westward away from AK, interacting with the bathymetry and the nearby islands on which it runs up, the deep over 50 m trough near the volcano rebounds as a backwash that causes an over 40 m runup onto AK's collapse scar. Subsequently, waves reflected from nearby islands propagate back onto AK and onto other islands. Both these reflected and initial waves cause significant runup on the most exposed parts of the three surrounding islands of Sertung to the W, Panjang to the E, and Rakata to the S (Fig. 1b,c, and Supplementary #4).

For the 0.27 km^3 volume likeliest scenario, Fig. 4a,b show instantaneous surface elevations computed in grid G0 at time $t = 410 \text{ s}$ for the granular slide, and the envelope of maximum elevation up to this time, respectively. In Fig. 4a, large leading waves (3–10 m high) are seen to propagate in all directions, but preferentially in the SW/SSE, N and ENE directions. Surface elevation time series computed for both rheologies at numerical wave gauges 5–9 (Fig. 1b; Table 1) up to 410 s (Fig. 5e) confirm these observations and show overall a good agreement with each other, particularly further away from AK (e.g., leading elevation waves in gauges 6 and 9). However, the granular slide causes significantly larger waves at gauge 7 than the viscous slide, which appears to result from the different behaviour and runout of the slide material in the northern direction (see animations in Supplementary #2). At gauge 5, close to AK in the SW direction, the leading waves are 33 and 42 m high, for the viscous and granular slides respectively. At gauges 6 and 9, further away from the volcano in the same direction, these are closer, in the 20–25 m range. At gauge 7 in the N direction, wave elevation reaches 5 and 13 m for the viscous and granular rheology, respectively, and at gauge 8 in the E direction, elevation reaches 6 and 7 m. At all numerical wave gauges, the large leading waves are followed by smaller oscillations, with a short 100–120 s period. Surface elevation time

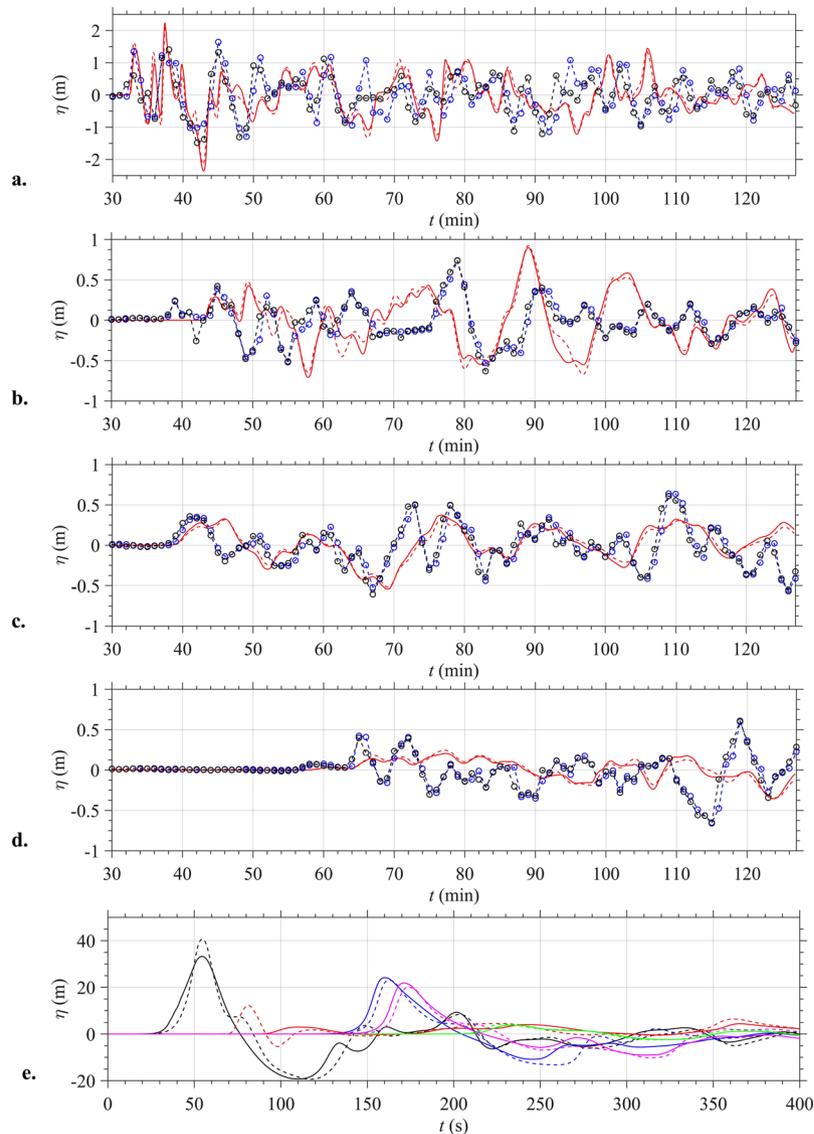


Figure 5. Surface elevation time series at: (a–d) Tide Gauges 1–4 (TG; Table 1; Fig. 1a) computed with FUNWAVE in grid G1; symbols are detided observations using 2 different sensors⁴⁵. (e) Numerical Wave Gauge 5–9 (WG; Table 1; Fig. 1b) computed with NHWAVE in grid G0: (black) 5, (blue) 6, (red) 7, (green) 8, (magenta) 9. Time $t = 0$ is estimated collapse time, 20:57' local time (UTC + 7). Computations are for (solid) viscous slide, or (dashed) granular slide.

series simulated for the three volumes and the two rheologies are compared at the five numerical wave gauges in Fig. S.6d–i Supplementary #3). Overall, results are consistent with those of Fig. 5e, with a larger effect of model rheology than collapse volume (within the tested range) on wave generation in the near-field.

The surface elevation envelope of Fig. 4b shows a 25–40 m maximum runup along AK's SW coast. On the directly exposed steep shores of the nearby islands of Rakata and Sertung runup reaches up to at least 45 m. On Panjang, which is located behind AK's main collapse direction, due to wave refraction around the volcano (see animations in Suppl. #2), runups still reach 15–25 m. Aerial images from a drone survey on Jan. 11th⁴¹ as well as coastline changes evident in ALOS-2⁴² and Planet.com⁴³ satellite images, show clear signs of large runup on the order computed with the model, on the E coast of Sertung, the W/SW coast of Panjang and NW coast of Rakata (see details in Suppl. #4). On Panjang, tree trunks are still *in-situ*, but there is very severe vegetation damage across the entire island from post-collapse, phreatomagmatic ash deposition, which obscures evidence of tsunami-related damage in the north.

Tsunami propagation and coastal impact modelling in the far-field. The 2D model FUNWAVE-TVD (Total Variation Diminishing version of the fully nonlinear Boussinesq wave model FUNWAVE-TVD⁴⁴; see methods) is used (Fig. 1a; grid G1, 100 m resolution) to simulate tsunami propagation from the near-field to the surrounding coasts in the far-field; the model is initialized using NHWAVE results such as shown in Fig. 4a. As for the near-field simulations, the three different collapse volumes were simulated for both

rheologies. Figure 4c shows instantaneous surface elevations computed at $t = 1810$ s for the 0.27 km^3 granular collapse, when the leading tsunami waves are about to reach the Tide Gauge 1 at Marina Jambu⁴⁵, on the W coast of Java (Fig. 1a; Table 1). At this time, the pattern of tsunami waves in the Sunda Straits is already quite complex, with waves propagating in many directions from their interaction with the bathymetry, the coastline, and other islands. This pattern would become even more complex as time increases, due to further multiple wave reflections within the Straits. Figure 4d shows the envelope of maximum surface elevations computed up to $t = 7,610$ s. As a result of the initial directionality of the generated wave train and bathymetric focusing, in particular by the steep linear scarp that divides the shallow eastern half of Sunda Straits from the much deeper Semangka trough to the west, wave elevations are greater in a number of narrow, preferential, directions, particularly in the S, SE, E and NE directions; the model predictions of tsunami coastal impacts are also larger at coastal sites in those directions.

For the 0.27 km^3 volume likeliest scenario (granular case), Fig. 5a–d compare surface elevation time series computed for the viscous and granular slides to field observations made at Tide Gauges 1–4⁴⁵ (Fig. 1a; Table 1). The surface elevation data recorded at 1 min intervals by two functioning sensors was detided using a Butterworth filter (Supplementary #3). Overall, there is a good agreement between the observed and simulated time series, particularly earlier in the time series. As summarized in Table 1, arrival times at each gauge are predicted to within 30–100 s of observations. Considering the 1 min interval of tide gauge data, this is quite a small discrepancy. Although the phase lag between simulations and observations increases with time, the trough to crest elevation of the largest waves is well captured in simulations. Additionally, each tide gauge is located within a harbor/marina or close to or behind protective coastal structures (Supplementary #3), whose complex geometry is not represented in the 100 m resolution grid G1. These surrounding structures would have induced reflections and seicheing that significantly affected the recorded signal. Finally, as reported by eyewitnesses, simulations predict that multiple large waves of short period (initially 5–8 minute period) impacted the coast, with the second or later waves being the largest.

Time series of surface elevation were also computed at the four Tide Gauges for the two additional collapse volumes and each rheology. Only small differences could be observed in these far-field results between the six scenarios (Fig. S.6a–d in Supplementary #3), indicating that our predictions of the tsunami coastal impact are not very sensitive to finer details of the collapse scenario assumed for Anak Krakatau (i.e., both small changes in volume and differences in rheology). Hence, given bathymetric/topographic data and model grids, far-field tsunami impact results are quite robust. Finally, we also note that the arrival time of the first elevation wave at each tide gauge is closely identical for each scenario (Fig. 5a–d, and Fig. S.6a–d in Supplementary #3), indicating that the tsunami travel time to shore is dominated by the wave celerity, which is primarily dependent on water depth and, for these fairly short waves, to some extent, also on wave frequency. For the very large initial waves generated near the volcano, amplitude dispersion effects caused by wave nonlinearity will also affect tsunami wave celerity and thus travel time. However, results at the near-field numerical wave gauges show (Fig. 5e, and S.6e–i in Supplementary #3) that wave time series are quite similar in amplitude (and frequency) and hence effects of wave nonlinearity on tsunami celerity will also be similar. As a result of these observations, a single arrival time at the Tide Gauges was reported in Table 1 for the six simulated scenarios.

Figure 6a,b show zoom-ins of the envelope of maximum surface elevation of Figs 4d, 6c–e show the corresponding flow depths at the shore, obtained by interpolating the envelope along the 0 m contour level, for two highly impacted sections of coast in West Java and Sumatra north of Anak Krakatau. The latter results are compared to data from an initial post-event survey²⁷ (note, the coordinates of surveyed data points were inferred from survey data maps using GOOGLE EARTH). Simulated and measured tsunami flow depths are found in good agreement, particularly in view of the fairly coarse model grid G1 used here. A good agreement is found in particular for the larger flow depth values measured at particular coastal sites, that result from the initial tsunami directionality and subsequent strong bathymetric focusing discussed above (note, at the time our work was performed, field data was lacking from the south-western coast of Java (Panaitan and Ujong Kulon) where the largest flow depths of 9.5 m are simulated in Fig. 6c). Additional closer zoom-ins of maximum runups are consistent with field observations and eyewitness reports (see Fig. S.3 in Supplementary #2).

Discussion

The potential tsunami hazard from the collapse of AK was recognised before the 2018 lateral collapse^{12,18}. Several factors may have predisposed the volcano for collapse of its southwest flank. First, its position on the northeast margin of the 1883 caldera has resulted in steeper submarine slopes to the southwest compared to the northeast^{20,36}. Second, there has been a gradual shift in vent position during its evolution towards the southwest and the caldera interior²⁰. Third, significant lava flow deltas have extended the base of AK to the west, which overlie potentially weak and altered volcanoclastic material produced by the Surtseyan activity of 1927–1960 and later hyaloclastites formed by ocean-entering lava flows as the lava deltas expanded. Finally, the very rapid growth of the volcano during the last 90 years had built a steep-sided summit cone consisting dominantly of unstable volcanoclastic material⁴⁶.

Despite uncertainties in the volume and rheology of the collapsed material, the effects of which were estimated in our simulations based on six scenarios, the good correspondence between the predicted and observed wave characteristics in the far-field at the coast, flow depths, and runups, confirms the relevance of the proposed geometric model for the collapse. In particular, the modelling is consistent with a tsunami source of $0.22\text{--}0.30 \text{ km}^3$, as inferred from our interpretation of post-collapse imagery (Figs 2, 3). Combined with the observation of a single phase of tsunami inundation, our results support an interpretation of an *en masse* failure of the SW flank of Anak Krakatau, rather than a more piecemeal process, and suggest that this was necessary to generate the observed tsunami. Our results also suggest that significant portions of the submarine flank of AK were involved in the landslide. As identified previously by¹² the mechanism of failure (for example, whether a landslide occurs in a multi-stage retrogressive process) is important in determining the characteristics of the generated tsunami,

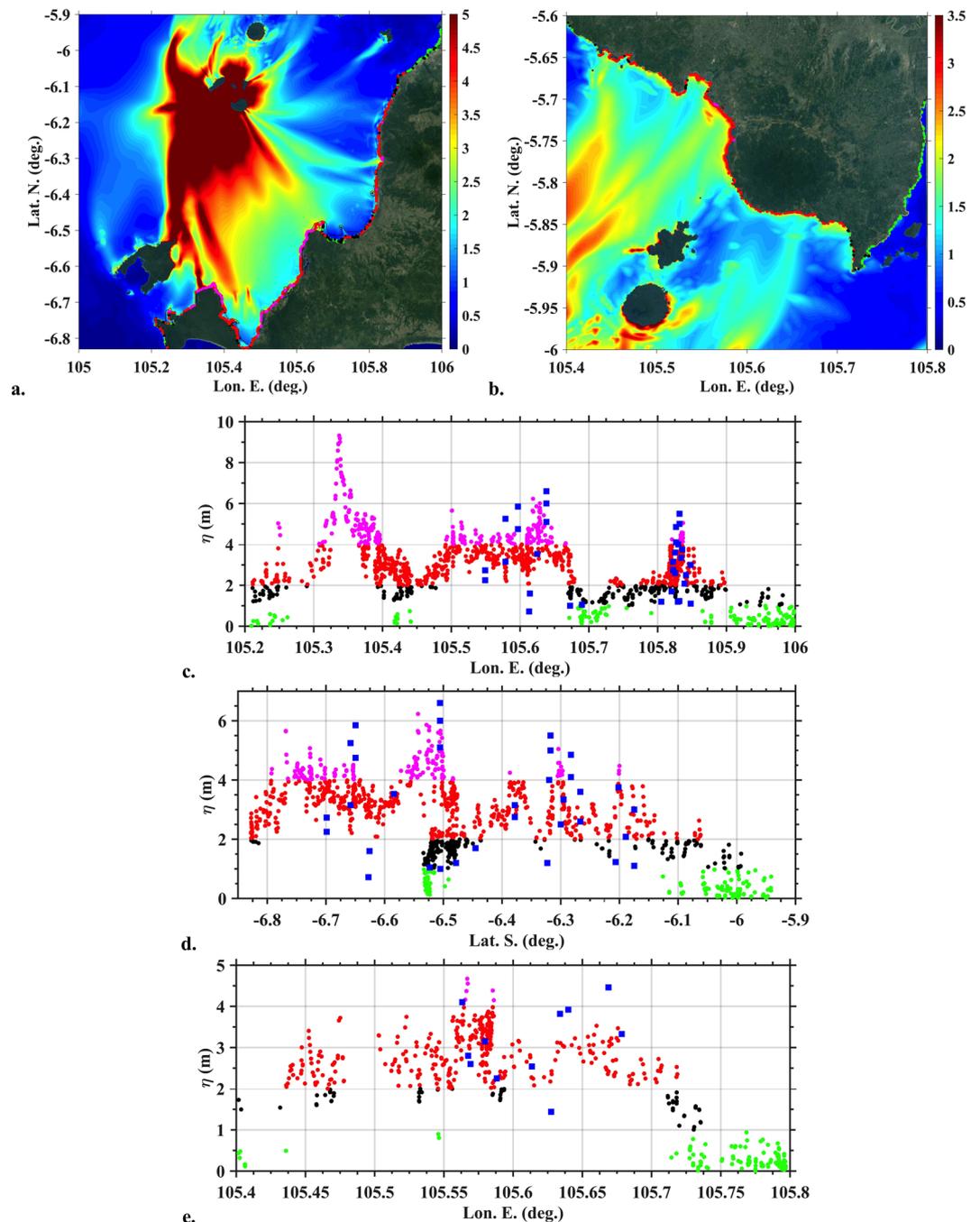


Figure 6. Maximum surface elevations and flow depth at the shore (along 0 m contour), simulated for the 0.27 km^3 scenario (granular slide) based on results of Fig. 4d, in two highly impacted areas (a,b). For clarity of their drawing in (a,b), four arbitrary height classes were selected for bullets in (c–e). (a,b) Coastal data point locations (coloured bullets) for flow depths plotted in (c,e), respectively, shown over the maximum surface elevation (colour scale in meter). (c,d) Flow depth calculated for area (a) (for clarity same results were plotted as a function of Lon. or Lat.); and (e) area (b). Level datum is MSL + 1.5 m. Blue square symbols indicate data measured in field surveys²⁷. Maps in (a,b) were produced by the authors using MATLAB version 2017b; topography from GOOGLE EARTH georeferenced satellite images was embedded in these maps using an API key: GOOGLE EARTH, 2019 DigitalGlobe.

particularly in the near-field. In this instance, our results confirm that failure had to be sufficiently rapid to represent a single tsunami source, which is consistent with the observed tsunami wave train. In addition, based on our simulations, it is not necessary to invoke any other possible tsunami-generating mechanism, such as a lateral blast or submarine explosion, to explain the magnitude and intensity of the tsunami.

It remains unclear what role eruptive activity had in triggering the lateral collapse. The collapse occurred several months into an eruption phase that was relatively intense but comparable in style to the activity that has

characterised preceding decades, and there was no clear indication of increasing or anomalous activity in the days or weeks preceding the collapse^{23,25}. Although a major change in eruptive behaviour, in the form of intense phreatomagmatic explosions, occurred after the collapse, current observations cannot identify whether this change was a response to the collapse, or whether an incipient change in behaviour occurred in the run-up to the collapse and may thus have acted as a trigger for failure. Given that there are examples of volcanic lateral collapses which were not associated with magmatic eruptions⁴⁷, it is plausible that the collapse was simply associated with a critical instability arising from the long-term rapid growth and local topographic and geological characteristics of AK.

Numerical simulations of six collapse scenarios, including both slide and tsunami generation, and propagation to shore, were performed that covered the range of uncertainty in estimated collapse volume, for two different slide rheologies (granular material and dense viscous fluid). Overall, results showed that waves generated in the near-field were quite similar for each scenario, particularly in the dominant direction of propagation (SW/SSE) (Figs 5e, S.6e–i). This similarity was even more pronounced for far-field results, as time series of surface elevations simulated at the four tide gauges were very close for each scenario (Figs 5a–d, S.6a–d). This “loss of memory” of the details of the tsunami source, as a tsunami propagates away from it is expected⁴⁸. Nearshore, this effect is even more pronounced as a result of the strong bathymetric control on the refraction and shoaling of tsunami waves⁴⁸. Given the relatively minor effect of both the collapse volume and slide rheology on the far-field tsunami impact (a feature already pointed out in the hypothetical scenarios simulated by¹²), simulation results were mostly shown here in detail and compared to field data for the likeliest collapse scenario (0.27 km³), assuming a granular rheology. This also implies that the limited choice of two model rheologies was adequate to estimate the far-field tsunami impact. A good agreement with field data was observed for the arrival time and elevation of the first few waves at tide gauges, as well as for the flow depth at the shore, which confirmed the relevance of the estimated collapse scenario(s) and the accuracy of numerical simulations with the NHWAVE/FUNWAVE-TVD model suite. The latter implies that a similar modelling approach could be applied to estimate tsunami hazard from future hypothetical volcanic collapses at other locations. Although these models have been well validated and benchmarked against laboratory data and some field case studies, to date, they had not been applied to a well-documented large scale volcanic collapse such as AK's.

One of the most significant developments following the AK eruption is the rapid post-collapse regrowth of the volcanic edifice by Surtseyan activity, which in the space of a few weeks significantly modified the morphology of the collapse scar. The rapidity of AK's rebuilding following the collapse signals the potential for future collapse. Seawater gained access to the vent in the immediate post-collapse period, generating distinctive Surtseyan ash plumes and in a sense resetting the evolution of the volcano back to the type of activity occurring in the 1930–1960 period. The rebuilding has already led to an enclosed crater lake, and it is likely that further activity will rapidly establish a new cone, with a transition back to magmatic eruptions and lava effusion. The southwest submarine flank is again being supplied with unconsolidated volcanoclastic material and if lava flows extend the coast in this direction, they will overlie these weaker lithologies, as was the situation prior to the December 2018 failure. AK's rapid regrowth emphasizes the need to monitor its activity by the acquisition of satellite data, as well as marine surveys to better determine submarine bathymetry. In both volume and volcano-tectonic setting, the AK event is broadly similar to lateral collapse-generated tsunamis at other arc volcanoes including Oshima–Oshima 1741, Unzen 1792, Ritter Island 1888, and Harimkotan 1930^{5,7,49–52}. These historical events suggest that tsunami-generating volcanic landslides >0.1 km³ have a global recurrence interval of less than 100 years. Although there are likely to be differences in collapse precursors, dominant lithologies and the style of landslide emplacement between all events e.g.^{16,51}, the AK lateral collapse is a modern equivalent of these historical events. As such, it is significant both in improving our understanding of the mechanisms of lateral collapse and as a benchmark against which current landslide-tsunami models can be tested and developed. The AK lateral collapse highlights the major hazard from this volcanic process, both in the impact of the initial landslide and, in island or coastal settings, from the associated tsunami generation.

If the regrowth of AK continues, there is a possibility of future collapses to the southwest and potential tsunami generation. However, the precursors to lateral collapses remain poorly understood, limiting the prospects for forecasting their timing. Given the short tsunami travel time from the AK volcanic source to the adjacent coastlines, mitigating against tsunami impacts after generation based on standard buoy-based measurements is a major challenge. Unlike earthquake-triggered tsunamis, which can be more easily detected, and hazard warnings issued to adjacent coastal communities, volcanic lateral collapses will continue to be a surprise, with the potential for causing significant loss of life and property destruction. One promising avenue for early detection of tsunamis from such non-seismic sources is the use of shore-based high frequency radars, combined with tsunami detection algorithms based on simulations of potential tsunami scenarios cf.⁵³, and literature reviews therein. Following the premonitory study of¹², the present research confirms that state-of-the-art numerical models can accurately simulate the potential hazard from volcanic induced tsunamis. Hence, realistic lateral collapse scenarios could be simulated for future volcanic events, at AK or elsewhere, and serve as a basis for developing tsunami hazard assessment maps and new types of tsunami detection and warning systems.

Methods

Numerical models and grids. We use a suite of hydrodynamic models for 3D landslide tsunami generation (NHWAVE)^{38–40} and 2D propagation (FUNWAVE-TVD)⁴⁴. These models have been extensively validated through benchmarking against laboratory and field data, as well as in the context of operational tsunami hazard assessment, for hypothetical^{38,39,53–56} or actual^{57–60} landslide tsunamis and other events. Much of this past work was performed in the US in parallel with or under the auspice of the National Tsunami Hazard Mitigation (NTHMP) program, including studies of potential volcanic flank collapses in the Canary Islands^{11,48} (see also⁹ for initial work on this event).

NHWAVE simulates 3D wave generation by slides, including frequency dispersion effects (vertical acceleration/non-hydrostatic effects), using only a small number of vertical, boundary fitted, σ -layers. Slides are modelled as an additional (depth-averaged) layer of subaerial and/or underwater material whose motion is fully coupled to that of the surrounding water. This layer can be of arbitrary geometry and also simulate non-hydrostatic effects⁶¹, with the slide material being either represented as a dense Newtonian fluid³⁹ or a saturated granular medium with intergranular and basal stresses governed by Coulomb friction^{40,55}. Both of these *rheologies* were used in the present work. FUNWAVE-TVD is a fully nonlinear and dispersive 2D Boussinesq-type model used (and initialized) here to propagate the initial near-field tsunami generated with NHWAVE to the far-field. FUNWAVE-TVD has been applied and validated for some actual tsunami case studies^{57–60}.

The 100 m resolution bathymetric and topographic data was interpolated to set-up model grids, to which +1.5 m was added, representing the tide elevation at Marina Jambu (tide gauge 1; Fig. 1a) at the time of tsunami arrival: (i) a Cartesian grid G1 used in FUNWAVE-TVD, with a 100 m horizontal resolution (Fig. 1a; 1950 by 1840 cells), and (ii) a 3D grid G0, with a 90 m horizontal resolution and 5 σ -layers, over a smaller footprint nested within G1 (Fig. 1b; 360 by 300 cells). Based on AK's collapse geometry described in Fig. 3 the likeliest collapse volume, and its lower and upper bounds (within the estimated uncertainty) are computed at 0.27, 0.22 and 0.30 km³, respectively in NHWAVE grid G0 (Supplementary #2 provides 3D images of the pre- and post-collapse discretized topography and bathymetry in grid G0).

Tsunami simulations. Following initial sensitivity analyses performed with another model (TSUNAMI-SQUARE⁶²), six collapse cases were simulated with NHWAVE in grid G0 (Fig. 1), for 0.22, 0.27, and 0.30 km³ volume of slide material (Fig. 3) represented: (i) either as a Newtonian fluid of density $\rho_c = 1,900$ kg/m³ and kinematic viscosity $\varphi_c = 0.5$ m²/s; or (ii) a granular medium with $\rho_c = 1,900$ kg/m³ for the solid part, an internal friction angle $\varphi_{ic} = 10^\circ$, a basal friction angle $\varphi_{bc} = 2^\circ$ (similar to that used by¹⁵, and a 40% porosity; with this data, assuming a water density $\rho_w = 1,025$ kg/m³, the average density of the granular medium is $\rho_{ac} = 1,550$ kg/m³). NHWAVE was run up to 410 s, the time when the generated tsunami approaches the boundary of grid G0 (Fig. 4a). Then surface elevation and horizontal velocity (calculated at the required depth $z = -0.531h$) were interpolated to initialize FUNWAVE-TVD simulations in grid G1⁵⁵, which were run for another 7,200 s (Fig. 4c,d). Time series of surface elevations shown in Fig. 5 (and Fig. S.6) are computed at 5 numerical wave gauges in NHWAVE and 4 tide gauges in FUNWAVE-TVD (Table 1).

Data Availability

The FUNWAVE-TVD and NHWAVE models used here (developed by some of the authors) are open source and available on github.com. All data is available from references cited in the text, except for unprocessed raw input-output model results, which have been archived and the authors can make available upon request.

References

1. AHA Centre. *Indonesia, Tsunami in Sunda Strait*, <http://adinet.ahacentre.org/reports/view/1383> (2018).
2. Simkin, T. & Fiske, R. S. *Krakatau 1883: the volcanic eruption and its effects*. Smithsonian Institution Press, Washington, D.C. (1983).
3. Verbeek, R. D. M. *Krakatau*. Government Press Batavia (1885).
4. Johnson, R. W. Large-scale volcanic cone collapse: the 1888 slope failure of Ritter Volcano, and other examples from Papua New Guinea. *Bulletin of Volcanology* **49**, 669–679 (1987).
5. Day, S. In *The Encyclopedia of Volcanoes*, 993–1009, Elsevier (2015).
6. Paris, R. Source mechanisms of volcanic tsunamis. *Philosophical Transactions of the Royal Society of London A: Mathematical, Physical and Engineering Sciences*, **373**, <https://doi.org/10.1098/rsta.2014.0380> (2015).
7. Auker, M. R., Sparks, R. S. J., Siebert, L., Croswell, H. S. & Ewert, J. A. Statistical analysis of the global historical volcanic fatalities record. *Journal of Applied Volcanology* **2**, 1–24, <https://doi.org/10.1186/2191-5040-2-2> (2013).
8. Day, S. *et al.* Submarine landslide deposits of the historical lateral collapse of Ritter Island, Papua New Guinea. *Marine and Petroleum Geology* **67**, 419–438, <https://doi.org/10.1016/j.marpetgeo.2015.05.017> (2015).
9. Ward, S. N. & Day, S. Cumbre Vieja volcano – Potential collapse and tsunami at La Palma, Canary Islands. *Geophysical Research Letters* **28**(17), 3397–3400 (2001).
10. Løvholm, F., Pedersen, G. & Gisler, G. Oceanic propagation of a potential tsunami from the La Palma Island. *Journal of Geophysical Research*, **113**, <https://doi.org/10.1029/2007JC004603> (2008).
11. Abadie, S. M., Harris, J. C., Grilli, S. T. & Fabre, R. Numerical modeling of tsunami waves generated by the flank collapse of the Cumbre Vieja Volcano (La Palma, Canary Islands): Tsunami source and near field effects. *Journal of Geophysical Research* **117**, C05030, <https://doi.org/10.1029/2011jc007646> (2012).
12. Giachetti, T., Paris, R., Kelfoun, K. & Ontowirjo, B. Tsunami hazard related to a flank collapse of Anak Krakatau Volcano, Sunda Strait, Indonesia. *Geological Society, London, Special Publications* **361**, 79–90, <https://doi.org/10.1144/sp361.7> (2012).
13. Tinti, S., Bortolucci, E. & Romagnoli, C. Computer simulations of tsunamis due to sector collapse at Stromboli, Italy. *Journal of Volcanology and Geothermal Research* **96**, 103–128, [https://doi.org/10.1016/S0377-0273\(99\)00138-9](https://doi.org/10.1016/S0377-0273(99)00138-9) (2000).
14. Ward, S. N. & Day, S. Ritter Island Volcano-lateral collapse and the tsunami of 1888. *Geophysical Journal International* **154**, 891–902 (2003).
15. Hunt, J. E., Wynn, R. B., Masson, D. G., Talling, P. J. & Teagle, D. A. H. Sedimentological and geochemical evidence for multistage failure of volcanic island landslides: A case study from Icod landslide on north Tenerife, Canary Islands. *Geochemistry Geophysics Geosystems* **12**, Q12007, <https://doi.org/10.1029/2011gc003740> (2011).
16. Watt, S. F. L. *et al.* From catastrophic collapse to multi-phase deposition: Flow transformation, seafloor interaction and triggered eruption following a volcanic-island landslide. *Earth and Planetary Science Letters* **517**, 135–147, <https://doi.org/10.1016/j.epsl.2019.04.024> (2019).
17. Watt, S. F. L. *et al.* Combinations of volcanic-flank and seafloor-sediment failure offshore Montserrat, and their implications for tsunami generation. *Earth and Planetary Science Letters* **319–320**, 228–240, <https://doi.org/10.1016/j.epsl.2011.11.032> (2012).
18. Camus, G., Gourgaud, A. & Vincent, P. M. Petrologic evolution of Krakatau (Indonesia): Implications for a future activity. *Journal of Volcanology and Geothermal Research* **33**, 299–316, [https://doi.org/10.1016/0377-0273\(87\)90020-5](https://doi.org/10.1016/0377-0273(87)90020-5) (1987).
19. Stehn, C. In *Krakatau, Proc. 4th Pacific Sci. Congr.*, Batavia I. 1–55.
20. Sudradjat, A. The morphological development of Anak Krakatau volcano, Sunda Straits. *Geol. Indonesia* **9**, 1–11 (1982).
21. Siswoidjoyo, S. In *Proc. Symp. on 100 years development Krakatau its surroundings, Jakarta, 23–27 August 1983*. 191–198, Indonesian Inst. Sci. Jakarta (1983).

22. Abdurrachman, M., Widiyantoro, S., Priadi, B. & Ismail, T. Geochemistry and Structure of Krakatoa Volcano in the Sunda Strait, Indonesia. *Geosciences* **8**, 111 (2018).
23. Global Volcanic Program. In *Bulletin of the Global Volcanism Network* Vol. 43:10 (ed Venzke, E.) Smithsonian Institution (2018).
24. Andersen, O. *Krakatau Volcano: Witnessing the eruption, tsunami and the aftermath 22–23th December 2018*, <http://www.oysteinlundandersen.com/krakatau-volcano-witnessing-the-eruption-tsunami-22december2018/> Accessed December 26, 2018 (2018).
25. PVMBG (Pusat Vulkanologi dan Mitigasi Bencana Geologi). *Activity reports for Anak Krakatau from 01/12/2018 – 03/01/2019*, <http://pvmbg.bgl.esdm.go.id/index.php/gunungapi/aktivitas-gunungapi> (2018).
26. Gurney, J. *Low frequency analysis of the 13:55 event (arrival at Bungbulang at 13:57 UTC)*, https://twitter.com/UKEQ_Bulletin. UK Earthquake Bulletin. Accessed on Dec. 28, 2018 (2018).
27. TDMRC (Tsunami Disaster and Mitigation Research Center) Post Sunda Strait Tsunami Survey, <http://tdmrc.unsyiah.ac.id/the-latest-update-from-post-sunda-strait-tsunami-survey/>. Accessed Jan. 8, 2019 (2019).
28. Muhari, A. In *Kompas*, (in Indonesian), <https://sains.kompas.com/read/2018/12/23/180319123/menyoal-dakwaan-pada-anak-krakatau-tentang-kasus-tsunami-selat-sunda?page=all>. Accessed Dec. 23, 2018 (2018).
29. Adristy, C. Kondisi Waspada, Begini Erupsi Gunung Anak Krakatau dari Udara, in *Serambin News*, <https://foto.tempco.com/read/69191/kondisi-waspada-begini-erupsi-gunung-anak-krakatau-dari-udara> (2018).
30. Foto-foto Letusan Gunung Anak Krakatau Sehari setelah Tsunami Menerjang Banten dan Lampung in *Serambin News*, <https://aceh.tribunnews.com/2018/12/24/foto-foto-letusan-gunung-anak-krakatau-sehari-setelah-tsunami-menerjang-banten-dan-lampung> (2018).
31. *Sentinel Hub*, <https://www.eorc.jaxa.jp/ALOS-2> (2018).
32. ALOS-2 (Advanced Land Observing Satellite-2; SAR mission)/Daichi-2, ALOS-2 (Advanced Land Observing Satellite-2; SAR mission) / Daichi-2 (.
33. Cole, P. D., Guest, J. E., Duncan, A. M. & Pacheco, J.-M. Capelinhos 1957–1958, Faial, Azores: deposits formed by an emergent tsurfeyan eruption. *Bulletin of Volcanology* **63**, 204, <https://doi.org/10.1007/s004450100136> (2001).
34. Kokelaar, P. Magma-water interactions in subaqueous and emergent basaltic. *Bulletin of Volcanology* **48**, 275–289, <https://doi.org/10.1007/bf01081756> (1986).
35. Moore, J. G. Structure and eruptive mechanisms at Surtsey Volcano, Iceland. *Geological Magazine* **122**, 649–661, <https://doi.org/10.1017/S0016756800032052> (2009).
36. Deplus, C. *et al.* Inner structure of the Krakatau volcanic complex (Indonesia) from gravity and bathymetry data. *Journal of Volcanology and Geothermal Research* **64**, 23–52, [https://doi.org/10.1016/0377-0273\(94\)00038-I](https://doi.org/10.1016/0377-0273(94)00038-I) (1995).
37. Siebert, L. Large volcanic debris avalanches: Characteristics of source areas, deposits, and associated eruptions. *Journal of Volcanology and Geothermal Research* **22**, 163–197 (1984).
38. Ma, G., Shi, F. & Kirby, J. T. Shock-capturing non-hydrostatic model for fully dispersive surface wave processes. *Ocean Modelling* **43–44**, 22–35, <https://doi.org/10.1016/j.ocemod.2011.12.002> (2012).
39. Kirby, J. T., Shi, F., Nicolsky, D. & Misra, S. The 27 April 1975 Kitimat, British Columbia, submarine landslide tsunami: a comparison of modeling approaches. *Landslides* **13**, 1421–1434, <https://doi.org/10.1007/s10346-016-0682-x> (2016).
40. Ma, G., Kirby, J. T., Hsu, T.-J. & Shi, F. A two-layer granular landslide model for tsunami wave generation: Theory and computation. *Ocean Modelling* **93**, 40–55, <https://doi.org/10.1016/j.ocemod.2015.07.012> (2015).
41. Reynolds, J. *Post-collapse image of Anak Krakatau*. *Earth Uncut TV*, <https://twitter.com/hashtag/Krakatau?src=hash>, Accessed Jan. 11, 2019 (2019).
42. *Observation Result for Eruption of Anak Krakatau Volcano in Indonesia by ALOS-2*, https://www.eorc.jaxa.jp/ALOS-2/en/img_up/dis_pal2_eruption_anak_krakatau_20181226.htm (2018).
43. Team, P. *Planet Application Program Interface: In Space for Life on Earth*. San Francisco, CA, <https://api.planet.com> (2017).
44. Shi, F., Kirby, J. T., Harris, J. C., Geiman, J. D. & Grilli, S. T. A high-order adaptive time-stepping TVD solver for Boussinesq modeling of breaking waves and coastal inundation. *Ocean Modelling* **43–44**, 36–51, <https://doi.org/10.1016/j.ocemod.2011.12.004> (2012).
45. Ina-COAP. *Tide Gauge data in Indonesia*, <http://tides.big.go.id/las/UI.vm> Accessed Jan. 3, 2019 (2019).
46. Hoffman-Rothe, A. *et al.* Monitoring Anak Krakatau Volcano in Indonesia. *Eos, Transactions American Geophysical Union* **87**, 581–586, <https://doi.org/10.1029/2006eo510002> (2006).
47. Cooke, R. J. S. In *Cooke–Ravian Volume of Volcanological Papers* (ed. Johnson, R. W.) 115–123 (Geological Survey of Papua New Guinea Memoir 10, 1981).
48. Tehranirad, B. *et al.* Far-field tsunami impact in the north Atlantic basin from large scale flank collapses of the Cumbre Vieja volcano, La Palma. *Pure and Applied Geophysics* **172**(12), 3,589–3,616, <https://doi.org/10.1007/s00024-015-1135-5> (2015).
49. Tsuji, Y. & Hino, T. Damage and inundation height of the 1792 Shimabara landslide tsunami along the coast of Kumamoto prefecture. *Bulletin of the Earthquake Research Institute, University of Tokyo* **68**, 91–176 (1993).
50. Belousova, M. & Belousov, A. Prehistoric and 1933 debris avalanches and associated eruptions of Harimkotan Volcano (Kurile Islands). *Period. di Mineral*, 99–101 (1995).
51. Inoue, K. Shimabara-shigatusaku earthquake and topographic changes by Shimabara catastrophe in 1792 (in Japanese with English Abstract). *Journal of the Japan Society of Erosion Control Engineering* **52**, 45–54 (2000).
52. Karstens, J. *et al.* From gradual spreading to catastrophic collapse – Reconstruction of the 1888 Ritter Island volcanic sector collapse from high-resolution 3D seismic data. *Earth and Planetary Science Letters* **517**, 1–13, <https://doi.org/10.1016/j.epsl.2019.04.009> (2019).
53. Guérin, C.-A., Grilli, S. T., Moran, P., Grilli, A. R. & Insua, T. L. Tsunami detection by high-frequency radar in British Columbia: performance assessment of the time-correlation algorithm for synthetic and real events. *Ocean Dynamics* **68**, 423–438, <https://doi.org/10.1007/s10236-018-1139-7> (2018).
54. Grilli, S. T. *et al.* Modeling of SMF tsunami hazard along the upper US East Coast: detailed impact around Ocean City, MD. *Natural Hazards* **76**, 705–746, <https://doi.org/10.1007/s11069-014-1522-8> (2015).
55. Grilli, S. T. *et al.* Modeling coastal tsunami hazard from submarine mass failures: effect of slide rheology, experimental validation, and case studies off the US East Coast. *Natural Hazards* **86**, 353–391, <https://doi.org/10.1007/s11069-016-2692-3> (2017).
56. Schambach, L., Grilli, S. T., Kirby, J. T. & Shi, F. Landslide Tsunami Hazard Along the Upper US East Coast: Effects of Slide Deformation, Bottom Friction, and Frequency Dispersion. *Pure and Applied Geophysics*, 1–40, <https://doi.org/10.1007/s00024-018-1978-7> (2018).
57. Grilli, S. T. *et al.* Numerical simulation of the 2011 Tohoku tsunami based on a new transient fem co-seismic source: Comparison to far- and near-field observations. *Pure and Applied Geophysics* **170**, 1333–1359, <https://doi.org/10.1007/s00024-012-0528-y> (2013).
58. Grilli, S. T. *et al.* Source constraints and model simulation of the December 26, 2004 Indian Ocean tsunami. *J. Waterway Port Coastal and Ocean Engineering* **133**(6), 414–428 (2007).
59. Ioualalen, M. *et al.* Modeling the 26th December 2004 Indian Ocean tsunami: Case study of impact in Thailand. *Journal of Geophysical Research* **112**, C07024, <https://doi.org/10.1029/2006JC003850> (2007).
60. Tappin, D. R. *et al.* Did a submarine landslide contribute to the 2011 Tohoku tsunami? *Marine Geology* **357**, 344–361, <https://doi.org/10.1016/j.margeo.2014.09.043> (2014).

61. Zhang, C., Kirby, J., Grilli, S., Shi, F. & Ma, G. In *Proc. 36th Intl. Conf. on Coast. Engng. (ICCE18; Balt. Jul 30 - Aug 3)*, ASCE 1, 74 (2018).
62. Xiao, L., Ward, S. N. & Wang, J. Tsunami Squares Approach to Landslide-Generated Waves: Application to Gongjiafang Landslide, Three Gorges Reservoir, China. *Pure Appl. Geophysics* **172**, 3639–3654, <https://doi.org/10.1007/s00024-015-1045-6> (2015).

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Author Contributions

S.G. performed or led the majority of tsunami simulations reported here, coordinated the work, and led the writing of the manuscript. D.R.T., S.C., S.F.W. and S.E. analysed the pre- and post-collapse morphologies of A.K., defined the failure geometry, and assessed the conditions that may have led to the volcano's failure. S.N.W. performed tsunami simulations with an alternate model and helped assessing results sensitivity to collapse parameters. A.G. and L.S. contributed to tsunami simulations and to the pre- and post-processing of model data and simulation results. C.Z. and J.K. worked on the landslide tsunami simulations. M.M., ran an alternate tsunami model and helped accessing bathymetric data. S.G., D.R.T., S.F.W. and S.C. did the most work on the final manuscript.

Additional Information

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