# Origin of nanosized diamonds in interstellar space and low-pressure-temperature Earth rocks 

S. K. Simakov ${ }^{\text {a }}$,<br>${ }^{\text {a}}$ Geological Department, St.Petersburg University, 7/9 Universitetskaya Nab., St.Petersburg, 199034, Russia


#### Abstract

Nanosized diamond particles in the interstellar space and in the Earth rocks related with water presence. In the paper proposed the model of the nanosized diamond particle formation oxidized water-carbon dioxide gaseous mixtures.


The origin of diamonds in interstellar space has been a topic of intense discussion since the discovery of presolar nanodiamonds in chondrites ${ }^{1}$. Meteoritic nanodiamonds provide information on the nucleosynthesis of evolved stars and the evolution of the astrophysical environment from which the solar system formed. Sellgren ${ }^{2}$ identified the relationship between the interstellar diamond and water ice. Nakano et al. ${ }^{3}$ related interstellar diamond formation with organic matter. Follow from these relationships Kouch et al. ${ }^{4}$ identified new formation routes of diamond in interstellar clouds and parent bodies of carbonaceous chondrites in laboratory experiments. It's the ice mixture of $\mathrm{H} 2 \mathrm{O}, \mathrm{CO}, \mathrm{NH} 3$, and $\mathrm{CH} 4(4: 2: 2: 1)$.

The formation of the bulk of Earth diamonds is due to the deep upper mantle rocks - kimberlites formed at P and T corresponded to diamond thermodynamic stability. Meanwhile in the last 40 years micro- and nanodiamonds have also been founded in shallow earth rocks formed at P-T parameters corresponded to graphite thermodynamic stability ${ }^{5-9}$ including the organic matrix ${ }^{10-11}$. The maximum diamond grade is observed in hydrothermal metasomatic zones of Kokchetave metamorphic massive situated in Northern Kazakhstan ${ }^{12}$. These diamonds have anomaly high concentration of N in comparing with kimberlitic, their main mineral inclusions are graphite, carbonate and water ${ }^{13}$.

For diamond, it has been argued that crystallization in a P-T regime where diamond is actually thermodynamically unstable with respect to graphite, may be possible due to kinetic factors ${ }^{14,15}$. Nanosize diamond particles have energetic preference upon graphitic particles of the same size and could be more stable at low P-T parameters ${ }^{16-19}$.

Fedoseev et al. ${ }^{20}$ show that critical radii of graphite and diamond nuclei depends upon to surface energy ( $\sigma$ ), atomic volume of carbon $(\mathrm{V}$ ) and chemical potential of the resaturation $(\Delta \mu)$ :

$$
\begin{equation*}
\mathrm{r}_{\mathrm{g}} / \mathrm{r}_{\mathrm{d}}=\mathrm{V}_{\mathrm{g}} \sigma_{\mathrm{g}} \Delta \mu_{\mathrm{d}} / \mathrm{V}_{\mathrm{d}} \sigma_{\mathrm{d}} \Delta \mu_{\mathrm{g}} \tag{1}
\end{equation*}
$$

$\sigma$ of nanosized diamond (d) and graphite (g) particles depends upon temperature and size of the particles ${ }^{21}$. From the $\sigma_{\mathrm{g}} / \sigma_{\mathrm{d}}$ relation given in ${ }^{22}$ and from (ref.18) it follows that for nanoparticles $V_{g} \sigma_{g} / V_{d} \sigma_{d} \approx 1$. Than $r_{g} / r_{d}$ ratio could be expressed as:

$$
\begin{equation*}
\mathrm{r}_{\mathrm{g}} / \mathrm{r}_{\mathrm{d}}=\Delta \mu_{\mathrm{d}} / \Delta \mu_{\mathrm{g}} \tag{2}
\end{equation*}
$$

From another side chemical potentials of the resaturation for diamond and graphite could be expressed by the next formulations:

$$
\begin{align*}
& \Delta \mu_{\mathrm{g}}=\mathrm{R} \ln \left(\mathrm{Pi} /^{2 i_{\mathrm{g}}}\right)  \tag{3}\\
& \Delta \mu_{\mathrm{d}}=\mathrm{R} \ln \left(\mathrm{Pi}^{2} / \mathrm{Pi}_{\mathrm{d}}\right) \tag{4}
\end{align*}
$$

where Pi and $\mathrm{Pi}_{\mathrm{g}, \mathrm{d}}$ are the real and equilibrium pressures of carbonaceous gases As it follows from equations (2-4), $\left(\mathrm{r}_{\mathrm{g}} / \mathrm{r}_{\mathrm{d}}\right)$ is connected with the ratio between real and calculated equilibrium pressures of carbonaceous gases. In the field of graphite stability $\mathrm{Pi}_{\mathrm{g}}<\mathrm{Pi}_{\mathrm{d}}$, which corresponds to preferable graphite formation from fluid phase. The condition of preferable diamond formation corresponds to $\mathrm{Pi}_{\mathrm{g}}>\mathrm{Pi}_{\mathrm{d}}$. The difference between $\Delta \mu_{\mathrm{g}}$ and $\Delta \mu_{\mathrm{d}}$ depends upon the difference of $\mathrm{Pi} / \mathrm{Pi}_{\mathrm{g}}$ and $\mathrm{Pi} / \mathrm{Pi}_{\mathrm{d}}$ and
at lower $\mathrm{Pi}_{\text {car }}$ it tends to zero, which corresponds to optimal condition of diamond formation from gaseous in the field of graphite stability.

In the hydrocarbon-hydrogen mixture the gas-solid reaction of hydrocarbon distruction could be proposed for carbon formation in a fluid:

$$
\begin{equation*}
\mathrm{CH}_{4} \rightarrow \mathrm{C}+2 \mathrm{H}_{2} \tag{I}
\end{equation*}
$$

At high temperatures equilibrated pressure of hydrogen would be greater than equilibrated pressure of methane and in vacuum $\mathrm{P}_{\mathrm{CH} 4(\mathrm{~d})}$ tends to $\mathrm{P}_{\mathrm{CH} 4(\mathrm{~g})}$. The predomination of hydrogen under hydrocarbon in the gaseous mixture is also suppressed the growth rate of graphite more than it suppressed the growth rate of diamond (ref.15). As a result, it could stabilize the diamond growth at these conditions. Follow from this effect, Deryagin and Fedoseev (ref.14) have provided the diamond growth on the diamond seeds at vacuum conditions.

Subsequent work has shown that the addition of oxygen to the hydrocarbon gases can stabilize the diamond nucleus formation in the P-T field of graphite stability ${ }^{23}$. Bachman et al. ${ }^{24}$ have shown that the synthesis of diamonds under metastable conditions corresponds to the narrow fluid composition in the $\mathrm{C}-\mathrm{O}-\mathrm{H}$ system. Calculations done for this system show that $\mathrm{P}_{\mathrm{CH} 4(\text { car })}$ is very low in the water and carbon dioxide stability fields of the system, closed to upper limit of carbon stability by oxygen (CCO buffer) ${ }^{25}$. In these fields at lower pressure and temperature $\mathrm{P}_{\mathrm{CH} 4(\mathrm{~g})} \approx \mathrm{P}_{\mathrm{CH} 4(\mathrm{~d})}$ which corresponds to diamond nucleus stabilization -Fig.1. This conclusion coincides with the established fact that diamond is more stable in an oxygen environment than graphite because oxygen reduces graphite to a greater degree than diamond ${ }^{26}$. From this scheme, we can explain the relationship of interstellar and Kokchetave diamonds with water.

Noteworthy for metastable diamond formation is the role of nitrogen ${ }^{27,28}$. It presents in interstellar clouds and parent bodies of carbonaceous chondrites associated with nanodiamonds (ref. 3,4) and in extra-high concentration in Kokchetave diamonds ${ }^{29,30}$. From another side, it's known that N is the main crystallochemical mixture in diamonds ${ }^{31}$ and it has stabilization role at the diamond synthesis at high P-T conditions. The experiments provided at $500^{\circ} \mathrm{C}$ and total pressure of nearly 1000 bar show that nitrogen presence stabilized the diamond formation at these conditions ${ }^{32,33}$. It explains the high concentration of N in Kokchetave diamonds and correlation of ammonium with meteorite and interstellar nanodiamonds.

As a result, it is possible to conclude that nanosized diamond particles could be formed from carbon-bearing fluids at low temperatures and pressures without seeds. The nano- and microdiamonds detected in shallow deep Earth rocks, meteorites and interstellar clouds could be formed from fluids at low P-T parameters corresponding to graphite stability.

## References:

1. Bernatowicz, T., Zinner, E.. Astrophysical Implications of the Laboratory Study of Presolar Materials, AIP Conference Proceedings 402, New York. (1997)
2. Sellgren, K. Aromatic hydrocarbons, diamonds, and fullerence in interstellar spce: puzzles to be solved by laboratory and theoretical astrochemistry. Spectrochimica Acta P.A. 57, 627-642. (2001)
3. Nakano, H., Kouchi, A., Arakawa, M., Kimura, Y., Kaito, C., Ohno, H., Hondoh, T. Alteration of interstellar organic materials in meteorites' parent bodies: a novel route in diamond formation. Proc. Japan Acad. Ser. B. 78, 277-281. (2002)
4. Kouchi, A., Nakano, H., Kimura1, Y., Kaito, C. Novel routes for diamond formation in interstellar ices and meteoritic parent bodies. The Astrophysical Journal. 626, L129-L132. (2005)
5. Novgorodova, M.I. \& Rasskazov, A.V. High-pressure carbon mineral phase formation as a result of heat explosion at shift transformation of graphite. Dokl. Akad. Nauk SSSR. 322, 379-381. (1992)
6. Rozen, O.M., Zorin, U.M. \& Zayachkovsky, A.A. Diamond foundation in connection of precembrian eclogites of Kokchetave massive. Dokl. Akad. Nauk SSSR. 203, 674-676. (1972)
7. Wirth, R. \& Rocholl, A. Nanocrystalline diamonds from the Earth's mantle underneaath Hawaii. Earth Plan. Scie. Lett. 211, 357-362. (2003)
8. Dobrzhinetskaya, L.F., Eide, E.A., Larsen, R.B., Sturt, B.A., Tronnes, R.G., Smith, D.C., Taylor, W.R., Posukhova, T.V. Microdiamonds in high-grade metamorphic rocks of the Western Gneiss region, Norway. Geology. 23, 597-600. (1995)
9. Sobolev, N.V. \& Shatsky, V.S. Diamond inclusions in garnets from metamorphic rocks; a new environment for diamond formation. Nature. 343, 742-746. (1990)
10. Dubinchuk, V.T., Kochenov, A.V., Pen’kov, V.F., Sidorenko, G.A., Uspenski, V.A. New formations in the organic matter of sidementary rock as a result of radiation. Dokl. Akad. Nauk SSSR. 231, 973-976. (1976)
11. Novgorodova, M.I., Usupov, R.G. \& Dmitrieva, M.T. Cubic carbide of silicium in the integrowing with graphite and diamond from mumie. Dokl. Akad. Nauk SSSR. 277, 1222-1227. (1984)
12. Pechnikov, V.A., Kaminsky, F.V. Diamond potential of metamorphic rocks in the Kokchetav Massif, northern Kazakhstan. Eur. J. Mineral. 20, 395-413. (2008)
13.De Corte, K., Cartigny, P., Shatsky, V.S., De Paepe, P., Sobolev, M.V. \& Jovay, M. Characteristics of microdiamond from UHPM rocks of the Kokchetav massif (Kazakhstan). Proc. VII ${ }^{\text {th }}$ Int. Kimb. Conf. 2, 174-182. (1999)
13. Deryagin, B.V. \& Fedoseev, D.V. Growth of diamond and graphite from the gas phase. Nauka, Moscow. 115 p. (1977)
14. Chauhan, S.P., Angus, J.C. \& Gardner, N.C. Kinetics of carbon deposition on diamond powder. J.Appl. Phys. 47, 4746-4754 (1976)
15. Badziag, P., Verwoerd, W.S., Ellis, W.P. \& Greimer N.R. Nanometer-sized diamonds are more stable than graphite. Nature. 343, 244-245. (1990)
16. Chaikovskii, E.F., Rosenberg, G.H. Phase diagram of carbon and possibility of diamond formation at low pressures. Dokl. Akad. Nauk SSSR 279, 1372-1375. (1984) 18. Gamarnik, M.Y. Energetical preference of diamond nanoparticles. Physical Rev. B. 54, 2150-2156. (1996)
17. Tawson, V.L., Abramovich, M.G. Polymorphism of crystals and phases size effect: transformation diamond to graphite. Dokl. Akad. Nauk SSSR 287, 291-295. (1986)
18. Fedoseev, D.V., Deryagin, B.V., Varshavskaya, I.G., Semenova-Tyan-Shanskaya, A.S. Diamond crystallization. Nauka, Moscow. 134 p. (1984)
19. Magomedov, M.N. About the relationship of surface ebergy with size and form of nanocrystals. Phys. Tverd. Tela. 46, 924-937. (2004)
20. Nuth, J. A. Small-particle physics and interstellar diamonds. Nature. 329, 589. (1987)
21. Simakov, S.K. Thermodynamic estimation of oxygen-hydrogen conditions influence on diamond and graphite critical nucleus formation at processes of methane destruction at low pressures. Rus. J. Phys.-Chem. 69, 3460-347. (1995) 24. Bachman, P.K., Leers, D., Lydtin, H. Towards a general concept of diamond chemical vapour deposition. Diam. Relat. Mater., 1, 1-12. (1991)
22. Simakov S.K. Redox state of Earth's upper mantle peridotites under the ancient cratons and its connection with diamond genesis. Geoch. Cosm. Acta. 62, 1811-1820. (1998)
23. Rudenko, A.P., Kulakova, I.I. \& Skvortsova, V.I. Chemical diamond synthesis. Aspects of general theory. Usp. Him. (Rus. Chem. Rev.) 62, 99-117. (1993)
24. Simakov, S.K. Possibility of metastable diamond formation from fluids in continental crust conditions. Dokl. Akad. Nauk SSSR. 278, 953-957. (1984)
25. Simakov, S.K. Physsico-chemical aspects of diamond-bearing eclogite formation in the upper mantle and Earth crust rocks. Magadan, 187 p. (2003)
26. Cartigny, P., De Corte, K., Shatsky, V.S., Ader, M., De Paepe, P., Sobolev, N.V., Javoy, M. The origin and formation of metamorphic microdiamonds from the Kokchetav massif, Kazakhstan: a nitrogen and carbon isotopic study. Chem. Geol. 176, 265-281. (2001)
27. Finney, K.S., Fisher, D., Griffin, W.L., Harris, G.W., Sobolev, N.V. Nitrogen agregates in metamorphic diamonds of Kazakhstane. Geoch. Cosm. Acta. 58, 51735177. (1994)
28. Kaiser, W. \& Bond, W.L. Nitrogen - a major impurity in common type 1 diamond. Phys. Rev. 115, 857-863. (1959)
29. Simakov, S.K., Dubinchuk, V.T., Baidakova, M.V. Synthesis of nanosize diamond and diamondlike phases at lower temperatures and pressures. NDNC2007 Abstract Book. 279. (2007)
30. Simakov S.K., V.T. Dubinchuk, M.P. Novikov, N.N., Drozdova I.A. Formation of Diamond and Diamond-Type Phases from the Carbon-Bearing Fluid at PT Parameters Corresponding to Processes in the Earth's Crust. Dokl. Earth Scie. 421, 835-837. (2008)

## FIGURE CAPTIONS:

Fig. 1. Variation of compositions of C-O-H system : A. - at $500^{\circ} \mathrm{C}$ and $10^{3}$ atm, B - at $250^{\circ} \mathrm{C}$ and $10^{-3} \mathrm{~atm}$. CCO- upper limit of carbon by oxygen, $\mathrm{Dm}-$ field of diamond preferable growth, point line - the boundary of this field.


A


B

Fig.. 1

