

## ORIGINAL ARTICLE

# Relevance of a crenarchaeotal subcluster related to *Candidatus Nitrosopumilus maritimus* to ammonia oxidation in the suboxic zone of the central Baltic Sea

Matthias Labrenz<sup>1</sup>, Eva Sintez<sup>2,3</sup>, Falko Toetzke<sup>1</sup>, Anita Zumsteg<sup>1,4</sup>, Gerhard J Herndl<sup>2,3</sup>, Marleen Seidler<sup>1</sup> and Klaus Jürgens<sup>1</sup>

<sup>1</sup>IOW—Leibniz Institute for Baltic Sea Research, Department of Biological Oceanography, Warnemuende, Germany; <sup>2</sup>Department of Biological Oceanography, Royal Netherlands Institute for Sea Research, Texel, The Netherlands; <sup>3</sup>Department of Marine Biology, University of Vienna, Vienna, Austria and <sup>4</sup>Eidg. Forschungsanstalt für Wald, Schnee und Landschaft WSL, Research Unit Soil Sciences, Birmensdorf, Switzerland

Marine pelagic redoxclines are areas of enhanced biogeochemical cycling inhabited by distinct functional groups of prokaryotes. In this study, the diversity and abundance of archaeal and bacterial nitrifying populations throughout a pelagic redoxcline in the central Baltic Sea were examined using a suite of molecular methods. 16S rRNA/rRNA gene as well as bacterial and archaeal *amoA* mRNA/*amoA* gene fingerprints and clone libraries revealed that the putative nitrifying assemblages consisted solely of one crenarchaeotal subcluster, named GD2, which was closely related to *Candidatus Nitrosopumilus maritimus*. Neither distinct differences between transcript- and gene-based fingerprints nor pronounced differences in the crenarchaeotal composition throughout the whole redoxcline were detected. The abundance of this GD2 subgroup, as determined by the oligonucleotide probe Cren537 and the newly developed and more specific probe Cren679 showed that GD2 and total crenarchaeotal cell numbers were nearly identical throughout the redoxcline. The highest GD2 abundance ( $2.3 \times 10^5$  cells ml<sup>-1</sup>) occurred in the suboxic zone, accounting for around 26% of total prokaryotic cells. Below the chemocline, GD2 abundance was relatively stable ( $1.5\text{--}1.9 \times 10^5$  cells ml<sup>-1</sup>). Archaeal *amoA* expression was detected only in the putative nitrification zone and formed a narrow band in the suboxic layer, where ammonium, oxygen, nitrate, nitrite and phosphate concentrations were below 5  $\mu\text{mol l}^{-1}$ . To our knowledge this is the first study to show the dominance of only one crenarchaeotal nitrifying key cluster in a natural habitat. The metabolic properties and survival mechanisms present in this cluster inside and outside the nitrification zone remain to be determined.

*The ISME Journal* (2010) 4, 1496–1508; doi:10.1038/ismej.2010.78; published online 10 June 2010

**Subject Category:** microbial population and community ecology

**Keywords:** central Baltic Sea; Crenarchaea; *amoA* gene; pelagic redoxcline

## Introduction

The Baltic Sea is the world's second largest brackish basin. The Baltic proper comprises a number of deep areas, of which the Gotland Deep is the largest. These areas are characterized by periodically anoxic conditions in their bottom waters. A stable halocline below 50–60 m separates the water column into an upper oxygenated layer and an underlying oxygen-deficient and anoxic/sulfidic layer (Lepland and

Stevens, 1998; Neretin *et al.*, 2003), resulting in a vertically stratified redoxcline that is responsible for a distinct prokaryotic stratification (Rheinheimer *et al.*, 1989; Labrenz *et al.*, 2007). For example, high dark CO<sub>2</sub> fixation is consistently found in the upper sulfidic layer (Jost *et al.*, 2010) and is driven mainly by chemolithoautotrophic Epsilonproteobacteria (Grote *et al.*, 2008; Glaubitz *et al.*, 2009). At the interface between nitrate and reduced sulfur compounds, Epsilonproteobacteria are also responsible for chemoautotrophic denitrification (Labrenz *et al.*, 2005; Brettar *et al.*, 2006), which represents probably the most important N loss process in the central Baltic Sea (Hannig *et al.*, 2007). If the accordant nitrate is not laterally introduced, then it must be produced by nitrification.

Correspondence: M Labrenz, Leibniz-Institut für Ostseeforschung Warnemünde (IOW), Seestr. 15, 18119 Rostock-Warnemünde, Germany.

E-mail: matthias.labrenz@io-warnemuende.de

Received 10 February 2010; revised 3 May 2010; accepted 4 May 2010; published online 10 June 2010

Studying the oxic–anoxic interface of the Gotland Deep from 1998 to 2000, Bauer (2003) determined mean nitrification rates of  $202 \text{ nmol NH}_4^+ \text{ L}^{-1} \text{ d}^{-1}$ . In that study, based on 16S rRNA as well as bacterial ammonia monooxygenase (*amoA*, the key enzyme in aerobic ammonia-oxidizing bacteria) gene analyses, only *Betaproteobacteria* phylogenetically related to *Nitrosomonas* or *Nitrosospira* could be identified as putative nitrifying organisms. However, the highest cell numbers determined by fluorescence *in situ* hybridization (FISH) corresponded to only 0.6–3.9% of total prokaryotic cell counts. In fact, until recently nitrification was thought to be exclusively a bacterial feature, but an *amoA*-like gene has been detected on an insert in a fosmid clone library from soil also carrying a crenarchaeotal 16S rRNA gene, thus demonstrating the nitrifying potential of non-thermophilic *Crenarchaeota* for the first time (Schleper *et al.*, 2005; Treusch *et al.*, 2005). In parallel, *amoA* genes on a presumptive archaeal metagenome fragment from the Sargasso Sea study were detected (Venter *et al.*, 2004), and later on also in the marine sponge symbiont *Cenarchaeum symbiosum* (Hallam *et al.*, 2006). The only free-living cultivated representative of the marine pelagic *Crenarchaeota* is still *C. Nitrosopumilus maritimus*, which is known to be able to oxidize ammonium aerobically to nitrite (Könneke *et al.*, 2005; Martens-Habbena *et al.*, 2009). Phylogenetically, *C. N. maritimus* is a member of marine group I (MGI) (Schleper *et al.*, 2005). MGI *Crenarchaeota* are widely distributed throughout the mesopelagic ocean, reaching cell abundances accounting for up to 30% of the oceanic picoplankton (Massana *et al.*, 2000; Karner *et al.*, 2001; Church *et al.*, 2009). While a considerable fraction of the MGI was suggested to be autotrophic (Herndl *et al.*, 2005; Ingalls *et al.*, 2006), MICRO-CARD-FISH (Herndl *et al.*, 2005; Teira *et al.*, 2006) and archaeal *amoA* analyses (Agogué *et al.*, 2008; De Corte *et al.*, 2008) demonstrated that members of this group are not restricted to an autotrophic lifestyle. In any case, it is likely that non-thermophilic *Crenarchaeota* have a significant role in the global nitrogen cycle (Francis *et al.*, 2005; Wuchter *et al.*, 2006; Erguder *et al.*, 2009).

Recently, the relevance of nitrifying *Crenarchaeota* and *Bacteria* was also shown for the suboxic zones of Black Sea redoxclines. Applying gene abundance and expression analyses, reaction-diffusion modeling, and  $^{15}\text{N}$  incubation experiments, Lam *et al.* (2007) identified ammonium-oxidizing *Crenarchaeota* as well as gammaproteobacterial ammonia-oxidizing bacteria as important nitrifiers and proposed that nitrification is coupled to anammox (anaerobic ammonia oxidation). 16S rRNA gene analyses by Coolen *et al.* (2007) identified different crenarchaeotal phylotypes throughout the Black Sea suboxic zone. The importance of nitrifying organisms, especially in the suboxic zones of Black Sea redoxclines, is also supported by biomarker analyses indicating the presence of

*Crenarchaeota* or nitrifying bacteria (Wakeham *et al.*, 2007; Schubotz *et al.*, 2009).

In principle, pelagic redoxclines of the central Baltic Sea are physicochemically comparable to those of the Black Sea, as these specific habitats share several important microbial organisms involved in the C-, N- and S-cycle (Labrenz *et al.*, 2007). However, in contrast to the Black Sea, central Baltic Sea redoxclines are irregularly oxygenated after inflow events of saline- and oxygen-rich North Sea water (Reissmann *et al.*, 2009), albeit eventually becoming re-established. Additional disturbances include small-scale turbulence, which can produce local mixing events (Lass *et al.*, 2003). Thus, the extended suboxic zones characteristic of the Black Sea (Yakushev *et al.*, 2008) are absent in Baltic redoxclines. Nevertheless, 16S rRNA fingerprinting analyses has indicated that MGI *Crenarchaeota* are also prominent in suboxic zones of the central Baltic Sea (Labrenz *et al.*, 2007).

The main objective of this study was to assess the diversity and abundance of nitrifying assemblages throughout a central Baltic Sea redoxcline. Although *amoA* gene and transcript analyses failed to provide evidence of bacterial ammonium-oxidizing capacities, one highly abundant nitrifying member of the MGI *Crenarchaeota* was instead detected.

## Materials and methods

The complete Materials & Methods part is given as Supplementary information.

### Sampling

Water samples were obtained from the central Baltic Sea onboard the RV *Maria S. Merian* in February 2006 (Baltic Sea monitoring station 271, Figure 1). Samples from different depths around the chemocline, defined as the uppermost presence of sulfide, were collected in free-flow bottles attached to a conductivity, temperature and depth-rosette.

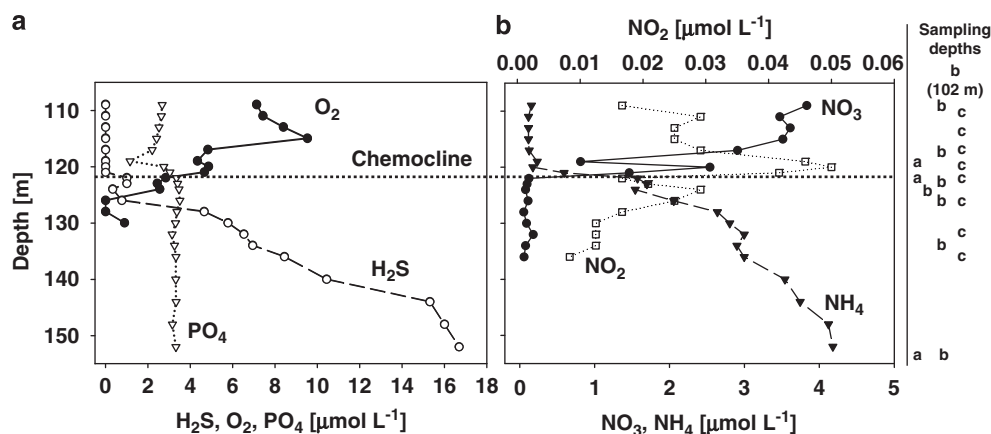
### Physicochemical structure of the water column

Chemical profiles of oxygen, hydrogen sulfide, ammonia, phosphate, nitrite and nitrate were determined as described elsewhere (Grasshoff *et al.*, 1983).

### Bacterioplankton sampling

For catalyzed reporter deposition fluorescence *in situ* hybridization (CARD-FISH), 100 ml of water samples were directly transferred to glass bottles and fixed with particle-free formaldehyde (2% final concentration) at 4 °C for 2–5 h. Portions of 40–45 ml were filtered onto white polycarbonate membrane filters (Millipore GTTP, 0.2- $\mu\text{m}$  pore size, 47-mm diameter, Molsheim, France), which were then rinsed with sterile seawater, air dried and stored at  $-80^\circ\text{C}$  until further processing. For nucleic acid





**Figure 2** Depth profile throughout the water column of the Gotland Deep (station 271) in February 2006. (a)  $\text{H}_2\text{S}$ ,  $\text{O}_2$ ,  $\text{PO}_4^{3-}$ ; (b)  $\text{NH}_4^+$ ,  $\text{NO}_2^-$  and  $\text{NO}_3^-$ . The dotted line denotes the chemocline. Sampling depths for molecular studies: (a) clone libraries (119, 122, 152 m); (b) SSCP fingerprinting (102, 109, 117, 122, 123, 124, 132, 152 m); (c) DGGE fingerprinting (111, 113, 115, 119, 121, 124, 130, 134 m).

consisting of a total of 25 (cDNA) or 35 (DNA) cycles, was done following the method of Leclerc *et al.* (2001). Single-stranded DNA was generated and purified and SSCP analysis was carried out according to Schwieger and Tebbe (1998). Individual bands excised from the SSCP gels were re-amplified as described by Pöhler *et al.* (2002). PCR products were purified using the MiniElute kit (Qiagen, Hilden, Germany) according to the manufacturer's protocol and sequenced by Seqlab (Göttingen, Germany).

**16S rRNA/rRNA gene clone library construction and RFLP.** Archaeal 16S rRNA clone libraries were prepared from genomic DNA or RNA extracts originating from the suboxic zone, suboxic-sulfidic interface, and sulfidic zone (Figure 2). The primer systems Arch21F/1492R, w036/1492R and Arch21F/w039 (Table 1) were used to amplify 16S rRNA for PCR, which was performed as described above. The PCR products were purified also as described above and cloned following the manufacturer's instructions using the pGEM-T-Easy Vector system (Promega) and competent *E. coli* JM109 cells. For restriction fragment length polymorphism (RFLP) and sequencing, the inserted fragment was PCR-amplified with the vector-specific primers T7 and SP6 (Table 1). The unpurified PCR products were digested with restriction enzymes *Hin*6I and *Msp*I and documented as described previously (Bond *et al.*, 2000). Identical RFLP patterns were grouped and representative cloned fragments sequenced by Seqlab using the primers Arch21F, w036, w069 or 1492R.

**Phylogenetic analysis.** 16S rRNA/rRNA gene sequences were examined for accuracy using the software SeqMan (DNASTar) and checked for chimeras using the Bellerophon program Huber *et al.* (2004). Phylogenetic affiliations of the partial sequences were initially estimated with the program BLAST (Altschul *et al.*, 1990). Sequences were aligned using the ARB software package (Ludwig

*et al.*, 2004) and those with a similarity > 99% were grouped together. Basic trees were reconstructed based on related public sequences and generated clones of approximately 1300 nucleotides. These were reduced to 1284 unambiguously alignable positions using group-specific filters. An evolutionary-distance dendrogram was constructed using neighbor-joining (NJ), maximum likelihood (ML) and maximum parsimony (MP) algorithms. Partial SSCP sequences were individually imported using ARB parsimony and group-specific filters.

**AmoA mRNA/amoA gene-based diversity analyses DGGE fingerprinting.** Ten nanograms of DNA or 20 ng of RNA, both originating from eight different depths throughout the redoxcline, were used as template for the PCR (Figure 2). Seven of these depths were slightly different from those depths used for the 16S rRNA-based analysis, but still from the same biogeochemical zone. The method of Coolen *et al.* (2007) was followed, with 25 (cDNA) and 35 (DNA) cycles in the PCR together with the archaeal *amoA* primers ammonia-oxidizing *Archaea* (AOA)-*amoA*-f and AOA-*amoA*-r (Table 1). DGGE was performed according to Coolen *et al.* (2004), using 9% polyacrylamide gels (ratio of acrylamide to bisacrylamide, 37.5:1) submerged in  $0.5 \times$  TAE buffer ( $20 \text{ mmol l}^{-1}$  Tris,  $20 \text{ mmol l}^{-1}$  acetic acid,  $0.5 \text{ mol l}^{-1}$  EDTA; pH 7.4) at  $60^\circ\text{C}$ . Individual bands excised from the DGGE gels were re-amplified using the above described *amoA* PCR protocol, purified using the MiniElute kit (Qiagen) as described by the manufacturer, and sequenced by Seqlab.

**AmoA clone libraries.** Archaeal and bacterial *amoA* clone libraries were prepared from extracts containing 10 ng of genomic DNA or 20 ng of RNA originating from the suboxic zone, suboxic-sulfidic interface and sulfidic zone (Figure 2). For PCR, the primer systems Arch-*amoA*F and Arch-*amoA*R

(archaeal (Table 1; Francis *et al.*, 2005)), *AmoA*-1F and *AmoA*-2R ( $\beta$ -proteobacterial (Rotthauwe *et al.*, 1997)), and *AmoA*-3F and *AmoB*-4R ( $\gamma$ -proteobacterial (Purkhold *et al.*, 2000)) were used (Table 1), with a total of 25 (cDNA) or 30 (DNA) PCR cycles. Cloning and sequencing steps were done as described for the 16S rRNA procedure.

**Phylogenetic analyses.** The procedures were the same as those described for 16S rRNA. Phylogenetic trees were reconstructed based on sequence information from DGGE bands and clones, and related to public sequences. These were reduced to unambiguously alignable positions using group-specific filters, resulting in 166 valid characters. An evolutionary-distance dendrogram was constructed using NJ, ML and MP.

#### Gene probe design

Based on the above-described 16S rRNA and rRNA gene clone libraries, a probe directed against crenarchaeotal cells of the suboxic zone was designed with the PROBE\_FUNCTION tool of the ARB package. Probe specificity was validated with the PROBE\_MATCH tool of the ARB package, the SILVA database (Pruesse *et al.*, 2007) and BLAST. Horseradish peroxidase (HRP)-labeled probes were synthesized by Biomers (Ulm, Germany). The specificity of the newly designed probe S<sup>\*</sup>-Cren-0679-a-A-18 (Table 2), which was named according to the Oligonucleotide Probe Database nomenclature (Alm *et al.*, 1996), was experimentally tested by whole-cell hybridization with different archaeal strains. With only one mismatch at the fourth base position of the gene probe, *C. 'Nitrosopumilus maritimus' SCM1* served as positive control. With 9 and 7 mismatches, respectively, *Halorubrum lacusprofundi* (DSM5036T) and *Pyrococcus furiosus* (DSM3638T) served as negative controls (Table 2). All strains were cultivated following the recommendations of the DSMZ (Braunschweig, Germany).

#### Cellular abundances

Total cell numbers were determined by flow cytometric analyses as described previously (Labrenz *et al.*, 2007). All HRP-labeled probes, appropriate sequences, target sites and hybridization conditions used in this study are described in Table 1 of the Supplementary information. Non-specific binding was determined using the NonEUB probe. By varying the formamide concentration, the hybridization conditions for the newly designed probe Cren679 were optimized (formamide concentrations tested: 0, 15, 20, 25, 30, 35, 50 and 60%). A concentration of 35% formamide in the hybridization buffer was found to be optimal for probe specificity.

The euryarchaeotal CARD-FISH enumeration was carried out according to the proteinase K protocol of

**Table 2** Probe sequence, appropriate target sites and corresponding sequences in other *Archaea*, including reference organisms

S <sup>*</sup> -Cren-0679-a-A-18 target/reference organisms	Mismatch	Probe 3'GCCTTC CTTCCCGATTTT' Target 5 'CGGAAGGAAGGGUAAAA'
Uncultured Baltic crenarchaeote clone_9	0	-----
Uncultured Baltic crenarchaeote SSCP_band_4	0	-----
Uncultured crenarchaeote MERTZ_2CM_56	0	-----
Uncultured archaeon CS_R008	1	----- U----
Uncultured archaeon clone BSA14-89m	1	--- U-----
' <i>Candidatus Nitrosopumilus maritimus</i> ' SCM1	1	--- U-----
Uncultured crenarchaeote MERTZ_0CM_362	1	----- G---
Uncultured Baltic crenarchaeote SSCP_band_3	2	--- U----- U-----
' <i>Cenarchaeum symbiosum</i> '	2	--- U----- U-----
' <i>Candidatus Nitrosocaldus yellowstonii</i> ' HL72	5	- C C - G - - U - - - - U - - - -
<i>Pyrococcus furiosus</i> (DSM3638 <sup>T</sup> )	7	- C C G G - - U - - - - U G - - -
<i>Halorubrum lacusprofundi</i> (DSM5036 <sup>T</sup> )	9	U C C G G - - U - - - A - U G - - -

The probe was named according to the Oligonucleotide Probe Database nomenclature (Alm *et al.*, 1996) and is abbreviated in the text as Cren679.

Teira *et al.* (2004), but this method was not successful for the Cren537 and Cren679 gene probes. Therefore, the crenarchaeotal cell wall was permeabilized and the subsequent steps carried out as described by Sekar *et al.* (2003). Specifically, filter sections were washed in pre-warmed washing buffer and the tyramide signal was amplified with 5-(and 6-) carboxyfluorescein-labeled tyramides. The preparations were counterstained with a previously described mixture of 4', 6'-diamidino-2-phenylindol (DAPI), Citifluor and VectaShield (Perntaler *et al.*, 2002). Stained filter sections were examined with an epifluorescence microscope (Axioskop 2 mot plus, Zeiss) as described previously (Grote *et al.*, 2007). Ratios of gene probe versus DAPI-stained cells were determined. Group-specific cell numbers were calculated by adjusting the ratios to flow cytometry-based total cell numbers.

#### Quantification of *amoA* gene transcripts by real-time PCR

Quantitative PCR (qPCR) standards were prepared as described by Agogué *et al.* (2008). All qPCR

analyses were conducted on an iCycler iQ5 thermocycler (Bio-Rad). Archaeal or  $\beta$ -proteobacterial *amoA* transcripts were determined in triplicate on non-diluted or 1:5 and 1:25 diluted cDNA and RNA samples. The methods of Coolen *et al.* (2006) and Agogu e *et al.* (2008) were used for the qPCR, together with the primers Arch-*amoA*-for and Arch-*amoA*-rev and the primers *amoA*-1F and *amoA*Ar NEW (Table 1). *AmoA* abundance of the samples were calculated by simultaneously quantifying the standard, prepared as 10-fold serial dilutions ranging from  $10^7$  to  $10^1$  copies.

#### Nucleotide sequence accession numbers

Sequences were deposited in the GenBank database under accession numbers GU386315 to GU386337, and GU455423.

## Results

### Physicochemical structure of the redoxcline

The pelagic redoxcline was established from approximately 110 m to 130 m. Sulfide was detectable from a depth of 122 m onwards and increased to  $17 \mu\text{mol l}^{-1}$  at a depth of 150 m (Figures 2a and b). At the chemocline, oxygen was still detectable and was present at concentrations of  $3 \mu\text{mol l}^{-1}$ , with an overlap of  $\text{O}_2$  and  $\text{H}_2\text{S}$  at lower concentrations measurable at a depth ranging from 122 m to 130 m (Figure 2a). Phosphate concentrations were generally between 2.2 and  $3.3 \mu\text{mol l}^{-1}$ ; however, at 119 m, they dropped to  $1.2 \mu\text{mol l}^{-1}$  (Figure 2a). Nitrate concentrations decreased sharply between 110 and 120 m but, together with nitrite, peaked at 122 m. Ammonium increased strongly below 120 m, reaching a maximum value of  $4.2 \mu\text{mol l}^{-1}$  at 150 m (Figure 2b).

### Gene-based diversity

Archaeal 16S rRNA/rRNA gene SSCP fingerprinting revealed 10 different and identifiable bands for the whole redoxcline. In general, neither distinct differences between 16S rRNA and 16S rRNA gene fingerprints nor pronounced differences in the archaeal composition throughout the redoxcline were detected for bands of highest relative abundance (Figure 3a; Supplementary information Figure 1A, B). Others, as SSCP bands 7 and 11 or SSCP band 5 were more pronounced in the sulfidic or in the suboxic zone, respectively. Phylogenetically, most of the SSCP bands belonged to the *Euryarchaeota*, with SSCP bands 1, 6, 7, 10, 11 and 12 representing a predominantly Baltic-Sea-specific cluster. SSCP bands 3 and 4 were members of the crenarchaeotal MGI and closely related to *C. Nitrosopumilus maritimus*, but constituting a specific phylogenetic cluster (Figure 4a).

From 16S rRNA/rRNA gene clone libraries based on different archaeal primer combinations,

91 clones were evaluated by RFLP analyses, but only one operational taxonomic unit could be detected for each of the three primer pairs. Subsequent sequencing of eight representative clones revealed that these were nearly 100% identical; thus, clone 9 was chosen as a representative sequence. Phylogenetically, clone 9 belonged to crenarchaeotal MGI, clustering with SSCP bands 3 and 4 and named Gotland Deep cluster 2 (GD2 (Figure 4a)). GD2 was 2% distant to *C. N. maritimus*. *Euryarchaeotal* 16S rRNA clones were not detected.

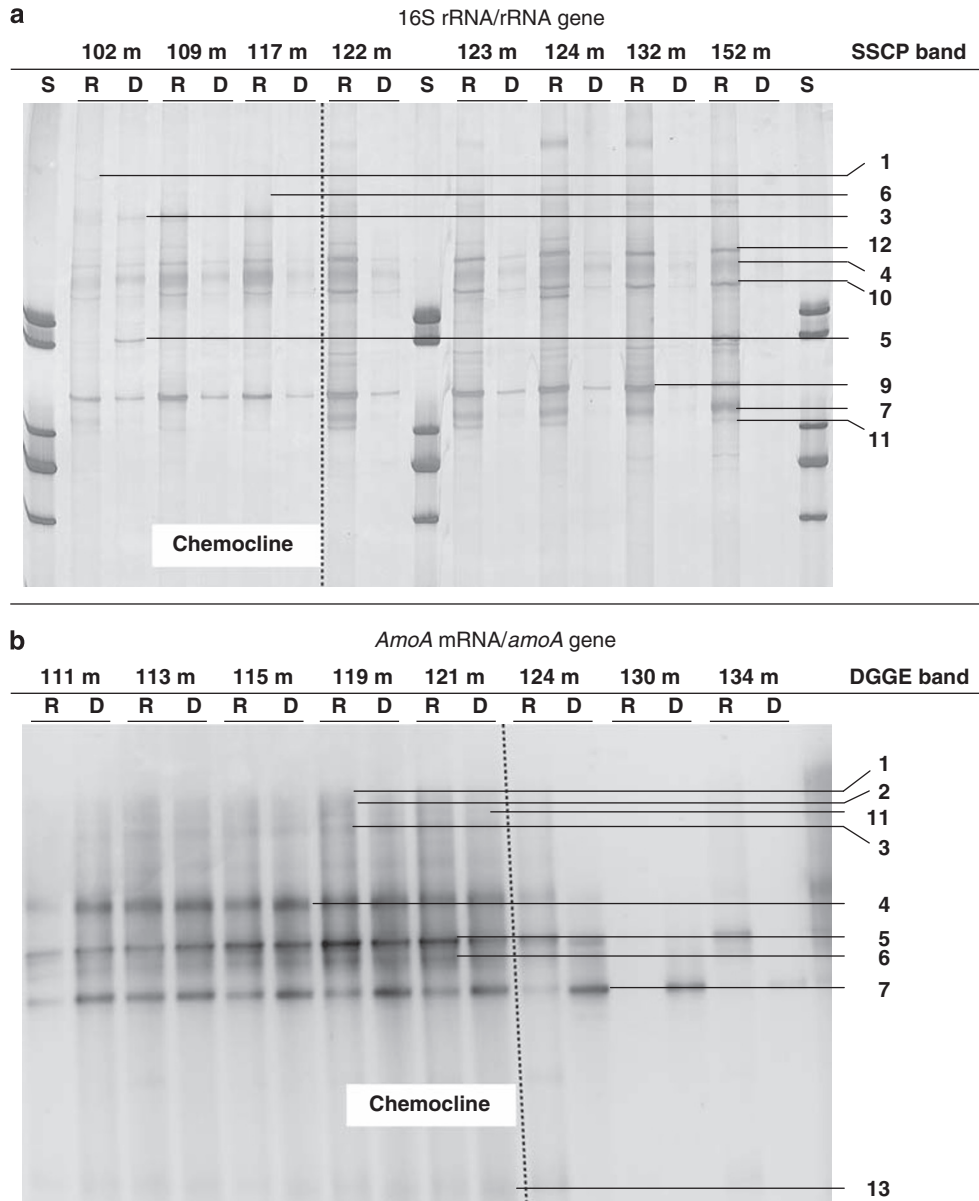
Archaeal *amoA* transcript and gene DGGE fingerprinting revealed eight distinct bands for the whole redoxcline (Figure 3b). No distinct differences between *amoA* transcript and gene fingerprints were detected (Figure 3b). Phylogenetically, all *amoA* sequences were related to *C. Nitrosopumilus*, but they again formed a distinct cluster (Figure 4b). With the exception of DGGE\_band\_7, all generated DGGE sequences were identical on the amino acid level. Their distances to DGGE\_band\_7 and *C. Nitrosopumilus amoA* were 2 and 3% on the amino acid and 3 and 8% on the nucleic acid level, respectively.

Of the clones generated by *amoA* transcript and gene clone libraries, 123 were evaluated by RFLP analyses, 9 OTUs identified and 31 representative clones finally sequenced. Comparable to the DGGE results, most of the clone sequences were identical or at least highly similar to each other (>98.5%). On the amino acid level, all *amoA* sequences were identical, with clones 14 and 15 chosen as representative sequences. Phylogenetically, these were similarly related to *C. Nitrosopumilus* and grouped with the *amoA* sequences generated by DGGE fingerprinting (Figure 4b).

$\beta$ - or  $\gamma$ -proteobacterial *amoA* cDNA or genes could not be amplified.

### Gene probe design

The group-specific rRNA-targeting oligonucleotide probe S-*\*-Cren-0679-a-A-18* (abbreviated as Cren679) was designed based on the eight nearly identical full-length crenarchaeotal 16S rRNA clones. Probe details and the number of mismatches with cultured organisms or related clone sequences are presented in Table 2. Cren679 was specific for all generated clones, SSCP\_band\_4, and, according to the SILVA database of August 2009, 44 other members of the GD2 cluster of organisms (Figure 4a). Within the target region of SSCP\_band\_3, two mismatches were found; however, it is unclear whether cells corresponding to SSCP\_band\_3 are detectable with Cren679. Despite one mismatch with *C. N. maritimus* SCM1 (Table 2), binding with this strain was recorded at all formamide concentrations tested, enabling the use of SCM1 as a positive control for the optimization procedure. Taking the *C. N. maritimus* mismatch

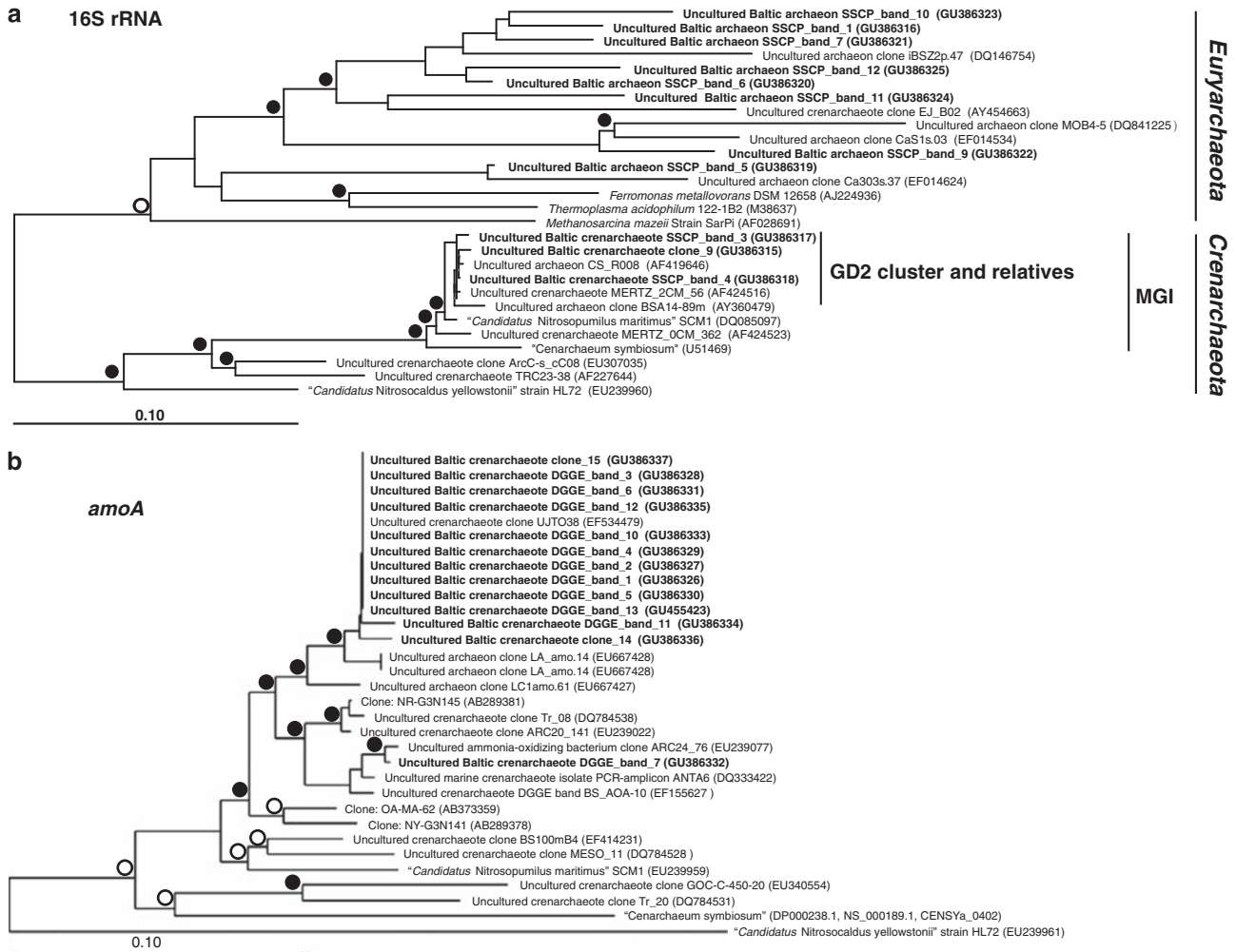


**Figure 3** Genetic fingerprints of samples covering the pelagic redoxcline. Sequenced bands are labeled by numbers. The phylogenetic affiliation of these bands is presented in Figure 4a and b. (a) Archaeal 16S rRNA/rRNA gene single-strand conformation polymorphism (SSCP); (b) Crenarchaeotal *amoA* mRNA/*amoA* gene denaturing gradient gel electrophoresis (DGGE). S, standard; R, cDNA; D, DNA.

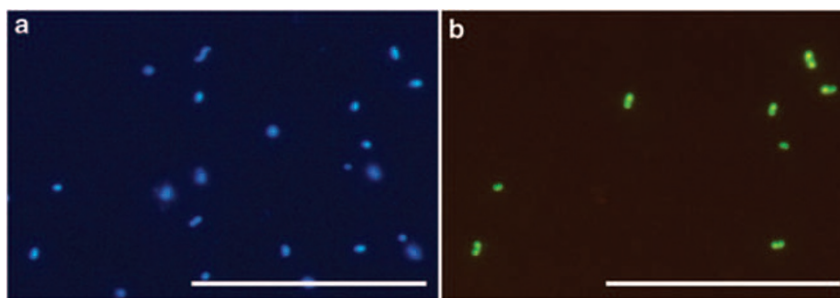
into account, gene probe Cren679 matched 838 out of 1712 MGI sequences included in the SILVA database of August 2009, but only one crenarchaeotal sequence (out of 4540) outside this group. Therefore, the specificity of gene probe Cren679 should be described as detecting members of the *C. Nitrosopumilus* and MGI cluster. *In situ*, cells detected by Cren679 showed a peanut-shaped cell morphology (Figure 5), similar to the morphology of *C. N. maritimus* cells (Könneke *et al.*, 2005). The probe did not hybridize with the negative control strains *Pyrococcus furiosus* (DSM3638<sup>T</sup>) and *Halorubrum lacusprofundi* (DSM5036<sup>T</sup>).

#### *Abundance of prokaryotes and specific groups*

Total prokaryotic cell numbers (TCN) varied between 0.5 and  $1.3 \times 10^6$  cells ml<sup>-1</sup>, with the highest numbers reached below the chemocline (Figure 6a). Bacterial cell numbers followed the prokaryotic pattern, with 60–80% EUB338-positive cells below the chemocline. Total crenarchaeotal cell numbers varied between 0.8 and  $2.3 \times 10^5$  cells ml<sup>-1</sup>, reaching highest numbers 1 m above the chemocline. In this zone, 25–26% of the TCN were Cren537-positive. Below the chemocline, crenarchaeotal cells remained relatively stable, with abundances of  $1.5\text{--}1.9 \times 10^5$  cells ml<sup>-1</sup> that decreased



**Figure 4** Phylogenetic relationships of representative sequences from the studied redoxcline samples. The evolutionary distance dendrogram was reconstructed using the Jukes-Cantor correction and NJ. Branching points supported by NJ, ML and MP algorithms are marked by a filled circle. Branching points supported by two algorithms are marked by an open circle. **(a)** Unrooted 16S rRNA sequence-based phylogenetic tree of members of the domain *Archaea*. **(b)** Unrooted archaeal *amoA* mRNA sequence-based phylogenetic tree of members of the *Crenarchaeota*.



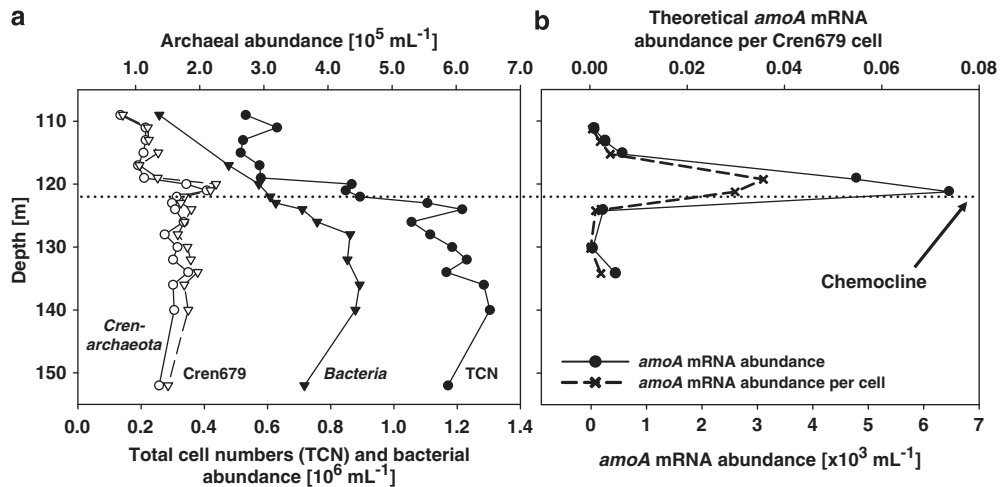
**Figure 5** Photomicrographs of DAPI- or CARD-FISH-stained microorganisms from the suboxic zone of the Gotland Deep redoxcline. Bar denotes 20 µm. **(a)** DAPI-stained cells; **(b)** Cells from the same microscopic field stained with the gene probe Cren679 exhibit a peanut-shaped morphology.

only slightly in the deeper sulfidic zone. Above the suboxic zone, Cren537 cell numbers declined to less than  $1.0 \times 10^5$  cells ml<sup>-1</sup>. Cren679-positive cells closely followed the total crenarchaeotal

distribution (Figure 6a). *Euryarchaeota* could not be detected by CARD-FISH analyses.

In a second profile of the Gotland Deep obtained in February 2006, the highest Cren679 abundance of





**Figure 6** Prokaryotic cell numbers and *amoA* mRNA abundances along the Gotland Deep redoxcline in February 2006. The dashed horizontal line indicates the chemocline. (a) Vertical distribution of total cell numbers (DAPI counts), bacterial abundance, crenarchaeotal, and Cren679 abundance. (b) Vertical distribution of *amoA* mRNA abundances and theoretical *amoA* mRNA abundances per Cren679 cell.

$2.1 \times 10^5 \text{ cells mL}^{-1}$ , accounting for 26% of the TCN, was also determined in the suboxic zone, where concentrations of  $4 \mu\text{mol O}_2 \text{ l}^{-1}$  and  $0.3 \mu\text{mol H}_2\text{S l}^{-1}$  were measured (data not shown). As in the first profile, Cren679-positive cells closely followed crenarchaeotal abundance (Cren537-positive cells) and vertical distribution, whereas *Euryarchaeota* were not detectable (data not shown).

#### Quantification of *amoA* transcripts throughout the redoxcline

Comparable to the *amoA* cloning results, no  $\beta$ -proteobacterial but only archaeal *amoA* transcripts were amplifiable using RT-qPCR primer combinations. The highest number of archaeal *amoA* transcripts ( $6.5 \times 10^3 \text{ mL}^{-1}$ ) coincided with the highest number of Cren679-positive cells in the suboxic zone (Figures 6a and b). Calculation of the ratios of *amoA* transcript numbers per Cren679 cell showed the highest ratios (0.04) in the suboxic zone and generally following the distribution of number of transcripts per  $\text{mL}^{-1}$  (Figure 6b).

## Discussion

Taken together, our data set shows that not only the nitrifying community but also the crenarchaeotal assemblage itself was characterized by an astonishing low diversity (Figure 4), consisting of only one crenarchaeotal subcluster (or potentially even a single species), which we named subgroup GD2. Phylogenetically, GD2 is closely related to, but clearly separated from, *C. Nitrosopumilus maritimus* (Figure 4a). GD2 abundance was highest in the suboxic zone, where it accounted up to 26% of the total prokaryotic abundance, but relatively high cell

numbers were also detected in the deeper sulfidic waters (Figure 6a). Based on maximal archaeal *amoA* transcript abundances, archaeal ammonia oxidation occurred in a narrow suboxic zone close to the chemocline, where ammonium, oxygen, nitrate, nitrite and phosphate concentrations were below  $5 \mu\text{mol l}^{-1}$ . The observation that the crenarchaeal assemblage was almost exclusively composed of the GD2 cluster leads to the conclusion that these cells were also responsible for *amoA* expression. This corroborates results obtained for the isolate *C. N. maritimus*, identified as an ammonia oxidizer (Könneke *et al.*, 2005; Martens-Habbena *et al.*, 2009).

#### Archaeal diversity

According to the results of both 16S rRNA and rRNA gene SSCP fingerprinting, members of the *Euryarchaeota* and *Crenarchaeota* were present throughout the pelagic redoxcline (Figures 3a and 4a). *Euryarchaeota* diversity was higher than that of *Crenarchaeota*, but summing up relative cren- and euryarchaeotal SSCP band areas as an indicator for accordant abundances, these were principally comparable throughout the redoxcline (Supplementary information Figure 1c and d). Moreover, potentially most abundant members of both phyla (SSCP bands 4 and 9) seemed to be uniformly distributed from the suboxic to the sulfidic zone, without pronounced differences between 16S rRNA and 16S rRNA gene fingerprints (Figure 3a; Supplementary information Figure 1a and b). The observed differences between eury- and crenarchaeotal diversities in the Baltic Sea might be a common phenomenon and is consistent with the findings of other studies, for instance, the lower diversity of pelagic MGI *Crenarchaeota* compared with *Euryarchaeota* described for the southern North Sea (Herfort *et al.*, 2007). Interestingly,

most of the euryarchaeotal sequences formed a nearly Baltic Sea-specific cluster that was at least 9% distant from the most similar environmental 16S rRNA sequence clone iBSZ2p.47 (Figure 4a). On the cellular level, *Euryarchaeota* were not detectable using the Eury806 gene probe. Euryarchaeotal 16S rRNA SSCP sequences were diverse and showed 0–4 mismatches to gene probe Eury806. The most abundant euryarchaeotal SSCP band 9 (Figure 3a) holds 1 mismatch at the edge of the probe and should be detected, but could be below the detection limit or the cell walls not properly permeabilized by the fixation procedure. Based on the data set of this study the exact cellular euryarchaeotal abundance remains unclear; however, most recent metagenomic analyses of a comparable Baltic Sea redoxcline showed that euryarchaeotal gene abundance was below 2% of the total genes (unpublished data), supporting low euryarchaeotal abundances. This is in contrast to the redoxclines from the Black Sea and the Cariaco Trench, where both *Eury-* and *Crenarchaeota* were present in considerable numbers (Lin *et al.*, 2006).

Only crenarchaeotal sequences were generated in the 16S rRNA/rRNA gene clone libraries but they were also detected in the SSCP fingerprints. The lack of euryarchaeotal clones is probably due to the lower coverage of the primer combinations used. For instance, according to the SILVA database of August 2009 Arch21F matches only 1176 out of 10936 sequences, in comparison with 5443 matches of the w036 SSCP forward primer. All gained crenarchaeotal sequences clustered together and, with the exception of the partial sequence SSCP\_band\_3 (98% sequence similarity), were >99% similar to each other. Assuming that 16S rRNA sequence similarities above 98.7–99% indicate an affiliation with a single species (Stackebrandt and Ebers, 2006), then nearly all of the Baltic crenarchaeotal 16S rRNA sequences could have arisen from one species. Archaeal *amoA* gene and transcript diversities supported the 16S rRNA analyses: the *amoA* sequences were also related to *C. Nitrosopumilus* (Figure 4b) and, with the exception of DGGE\_band\_7, all of the generated DGGE sequences were identical on the amino acid level. However, with a 16S rRNA distance of 2% and a phylogenetic separation supported by NJ, MP and ML algorithms (Figure 4a), the relationship to *C. N. maritimus* remains unexplained. Interestingly, the Baltic 16S rRNA sequences cluster together with sequences reported by Vetriani *et al.* (2003) (Figure 4a). Those sequences were obtained from comparable physico-chemical conditions in the Black Sea, supporting recent observations that marine pelagic redoxclines share similar microbial key organisms (Labrenz *et al.*, 2007).

Subcluster GD2 was quantified by CARD-FISH analyses. The highest cell numbers ( $2.3 \times 10^5$  cells ml<sup>-1</sup>) detected by the newly developed probe Cren679 were in the suboxic zone, accounting up to

26% of the TCN and matching the Cren537-positive cell abundance covering most of the known *Crenarchaeota* (De Corte *et al.*, 2008, Figure 6a). These results were reproduced in a second profile, in which Cren679 contributed up to 26% of the TCN. Considering the *C. N. maritimus* mismatch (Table 2), Cren679 matches approximately 45% of all MGI sequences included in the SILVA database of August 2009. Thus, it potentially covers different members of the MGI *Crenarchaeota*. Nonetheless, the results of our ribosomal, and functional gene and transcript-based diversity analyses were obtained using several primer, fingerprinting and cloning systems, all of which identified one 16S rRNA and one *amoA* cluster, both closely related to *C. N. maritimus* (Figures 4a and b). Also, on the morphological level, the specific detection of peanut-shaped cells (Figure 5) was an indication that *C. Nitrosopumilus* relatives are targeted by probe Cren679. Hence, there is strong evidence that only subcluster GD2 was detected by both gene probes, Cren679 and Cren537 (Figure 6a). Taken together, these data point to the high likelihood that: (1) the crenarchaeotal assemblages consist only of one phylotype, named GD2, which dominates the archaeal community throughout the entire redoxcline, (2) the detected *amoA* gene cluster belongs to this group and, consequently, (3) GD2 is a nitrifying *Crenarchaeota*.

#### Nitrification zone

Higher nitrite concentrations combined with lower oxygen or ammonium concentrations are indicative of a nitrification zone in pelagic redoxclines. In our study, these conditions were fulfilled for a narrow zone ranging from a depth of 120 m to 122 m just above the chemocline (Figures 2a and b). This area was designated as the nitrification zone because it contained the maximal number of archaeal *amoA* transcripts (Figure 6b). Analogously, archaeal *amoA* expression in Black Sea redoxclines is confined to a narrow nitrification zone with low oxygen and ammonium concentrations and high crenarchaeotal *amoA* expression (Lam *et al.*, 2007). Interestingly, the highest ratio of *amoA* transcript numbers per cell was in the nitrification zone (Figure 6b), but the value of 0.04 is, overall, lower than what might be expected. The rates of transcriptional regulation or *amoA* mRNA synthesis with respect to degradation are unknown for Cren679 cells, and absolute transcript numbers do not necessarily reflect their physiological state (Lam *et al.*, 2007). However, we developed a new sampling device for deeper redoxcline waters, which is based on *in situ* fixation of water samples, and most recent analyses showed that *amoA* transcript numbers are approximately 20 times higher when this system is used (unpublished data). Transferring these data to our study, absolute *amoA* transcript abundances per cell should be around 1 (assuming 100% active cells) to 4 (20% active cells) in the nitrification zone. This is in the

range what could be expected for a population actively oxidizing ammonia.

It is not surprising that archaeal *amoA* abundances were substantially lower in the sulfidic zone, which is inadequate for nitrification, than in the chemocline. It remains unclear, however, why the abundance of the GD2 subcluster is still high below the putative nitrification zone. This may be due to distribution processes (for example, sinking, mixing events) or to other metabolic pathways that enable a broad vertical distribution of the GD2 subgroup.

#### Selection factors

Despite the fact that different bacterial *amoA* primer systems were used (Table 1), no  $\beta$ - or  $\gamma$ -proteobacterial *amoA* genes or transcripts could be detected in any of the samples investigated. This finding was also supported by bacterial 16S rRNA clone libraries generated from the suboxic zone, which were devoid of  $\beta$ - or  $\gamma$ -proteobacterial nitrifiers (data not shown), and is in contrast to earlier studies of the Baltic Sea (Bauer, 2003) or the Black Sea (Lam *et al.*, 2007), where proteobacterial ammonia oxidizers were also found in the lower oxic and suboxic zones of the pelagic redoxcline.

In our investigated redoxcline, the environmental factors favoring the dominance of the GD2 cluster are as yet unknown. Martens-Habbena *et al.* (2009) demonstrated that the high affinity of *C. N. maritimus* for ammonium enables it to out-compete bacterial nitrifiers at low ammonium concentrations. Therefore, the authors proposed that ammonia oxidation kinetics can determine the niche separation of nitrifying *Archaea* and *Bacteria*. In our study, the highest Cren679 cell numbers (Figure 6a) and archaeal *amoA* abundances (Figure 6b) occurred at ammonium concentrations of less than  $1 \mu\text{mol l}^{-1}$ , which would support the assumption of Martens-Habbena *et al.* (2009). However, similar concentrations have been determined for the Black Sea where  $\beta$ - and  $\gamma$ -proteobacterial nitrifiers were detected (Lam *et al.*, 2007).

Erguder *et al.* (2009) reviewed the environmental factors shaping the ecological niches of AOA. These authors proposed that sulfide-containing waters favor AOA, because the inhibitory effect of sulfide or other S-containing compounds is more severe for bacterial than for archaeal nitrification. In our study, the overlap of sulfide- and oxygen-containing waters (Figure 2a), indicative of local mixing events or lateral intrusions, also supports a presumed inhibition of bacterial nitrifiers. It is known that these events are seasonally occurring in the Gotland Basin of the central Baltic Sea, but are more pronounced in fall and winter than in spring and summer (Axell, 1998). Thus, at the end of February, nitrifying GD2 could represent a sulfide-tolerating crenarchaeotal subcluster that is favored because of a longer period of sulfide pulses into the suboxic nitrification zones.

## Acknowledgements

We are grateful to the captain and crew of RV 'Maria S. Merian'. The excellent technical assistance of Christian Meeske, Heike Brockmüller, Bärbel Buuk, and Annett Grützmüller is greatly appreciated. Special thanks to Martin Könneke for providing cells of '*C. Nitrosopumilus maritimus*'. ML was supported by the Deutsche Forschungsgemeinschaft (DFG), project number LA 1466/4-1,4-2, and ES, and GJH by the Dutch Science Fund (NWO-ALW), project number 835.20.023.

## References

- Agogué H, Brink M, Dinasquet J, Herndl GJ. (2008). Major gradients in putatively nitrifying and non-nitrifying Archaea in the deep North Atlantic. *Nature* **456**: 788–791.
- Alm EW, Oerther DB, Larsen N, Stahl DA, Raskin L. (1996). The oligonucleotide probe database. *Appl Environ Microbiol* **62**: 3557–3559.
- Altschul SF, Gish W, Miller W, Myers EW, Lipman DJ. (1990). Basic local alignment search tool. *J Mol Biol* **215**: 403–410.
- Axell LB. (1998). On the variability of Baltic Sea deep-water mixing. *J Geophysical Res* **103**: 21667–21682.
- Bauer S. (2003). Structure and function of nitrifying bacterial communities in the eastern Gotland basin (central Baltic Sea). University of Rostock, Germany.
- Bond PL, Smriga SP, Banfield JF. (2000). Phylogeny of microorganisms populating a thick, subaerial, predominantly lithotrophic biofilm at an extreme acid mine drainage site. *Appl Environ Microbiol* **66**: 3842–3849.
- Brettar I, Labrenz M, Flavier S, Bötzel J, Kuosa H, Christen R *et al.* (2006). Identification of a *Thiomicrospira denitrificans*-like epsilonproteobacterium as a catalyst for autotrophic denitrification in the central Baltic Sea. *Appl Environ Microbiol* **72**: 1364–1372.
- Church MJ, Wai B, Karl DM, Delong EF. (2009). Abundances of crenarchaeal *amoA* genes and transcripts in the Pacific Ocean. *Environ Microbiol*: <http://dx.doi.org/10.1111/j.1462-2920.2009.02108.x>.
- Coolen MJL, Abbas B, van Bleijswijk J, Hopmans EC, Kuypers MMM, Wakeham SG *et al.* (2007). Putative ammonia-oxidizing Crenarchaeota in suboxic waters of the Black Sea: a basin-wide ecological study using 16S ribosomal and functional genes and membrane lipids. *Environ Microbiol* **9**: 1001–1016.
- Coolen MJL, Hopmans EC, Rijpstra WIC, Muyzer G, Schouten S, Volkman JK *et al.* (2004). Evolution of the methane cycle in Ace Lake (Antarctica) during the Holocene: response of methanogens and methanotrophs to environmental change. *Org Geochem* **35**: 1151–1167.
- Coolen MJL, Muyzer G, Schouten S, Volkman JK, Sinninghe Damsté JS. (2006). Sulfur and methane cycling during the holocene in Ace Lake (Antarctica) revealed by lipid and DNA stratigraphy. In: Neretin LN (ed). *Past and Present Marine Water Column Anoxia*, Vol 64. Springer: Dordrecht, The Netherlands, pp. 41–65.
- De Corte D, Yokokawa T, Varela MM, Agogue H, Herndl GJ. (2008). Spatial distribution of *Bacteria* and *Archaea* and *amoA* gene copy numbers throughout the water

- column of the Eastern Mediterranean Sea. *ISME J* **3**: 147–158.
- Delong EF. (1992). Archaea in coastal marine environments. *Proc Natl Acad Sci USA* **89**: 5685–5689.
- Erguder TH, Boon N, Wittebolle L, Marzorati M, Verstraete W. (2009). Environmental factors shaping the ecological niches of ammonia-oxidizing archaea. *FEMS Microbiol Rev* **33**: 855–869.
- Francis CA, Roberts KJ, Beman JM, Santoro AE, Oakley BB. (2005). Ubiquity and diversity of ammonia-oxidizing archaea in water columns and sediments of the ocean. *Proc Natl Acad Sci USA* **102**: 14683–14688.
- Glaubitx S, Lueders T, Abraham W-R, Jost G, Jürgens K, Labrenz M. (2009). <sup>13</sup>C-isotope analyses reveal that chemolithoautotrophic *Gamma*- and *Epsilonproteobacteria* feed a microbial food web in a pelagic redoxcline of the central Baltic Sea. *Environ Microbiol* **11**: 326–337.
- Grasshoff K, Erhardt M, Kremling K. (1983). Methods of seawater analysis, Vol 2nd, Verlag Chemie Weinheim: New York.
- Grote J, Jost G, Labrenz M, Herndl GJ, Jürgens K. (2008). *Epsilonproteobacteria* represent the major portion of chemoautotrophic bacteria in sulfidic waters of pelagic redoxclines of the Baltic and Black Seas. *Appl Environ Microbiol* **74**: 7546–7551.
- Grote J, Labrenz M, Pfeiffer B, Jost G, Jürgens K. (2007). Quantitative distributions of *Epsilonproteobacteria* and a *Sulfurimonas* subgroup in pelagic redoxclines of the central Baltic Sea. *Appl Environ Microbiol* **73**: 7155–7161.
- Hallam SJ, Konstantinidis KT, Putnam N, Schleper C, Watanabe Y-i, Sugahara J *et al*. (2006). Genomic analysis of the uncultivated marine crenarchaeote *Cenarchaeum symbiosum*. *Proc Natl Acad Sci USA* **103**: 18296–18301.
- Hannig M, Lavik G, Kuypers MMM, Woebken D, Martens-Habben W, Jürgens K. (2007). Shift from denitrification to anammox after inflow events in the central Baltic Sea. *Limnol Oceanogr* **52**: 1336–1345.
- Herfort L, Schouten S, Abbas B, Veldhuis MJW, Coolen MJL, Wuchter C *et al*. (2007). Variations in spatial and temporal distribution of Archaea in the North Sea in relation to environmental variables. *FEMS Microbiol Ecol* **62**: 242–257.
- Herndl GJ, Reinthaler T, Teira E, van Aken H, Veth C, Pernthaler A *et al*. (2005). Contribution of Archaea to total prokaryotic production in the deep Atlantic Ocean. *Appl Environ Microbiol* **71**: 2303–2309.
- Hornek R, Pommerening-Röser A, Koops H-P, Farnleitner AH, Kreuzinger N, Kirschner A *et al*. (2006). Primers containing universal bases reduce multiple *amoA* gene specific DGGE band patterns when analysing the diversity of beta-ammonia oxidizers in the environment. *J Microbiol Methods* **66**: 147–155.
- Huber T, Faulkner G, Hugenholtz P. (2004). Bellerophon: a program to detect chimeric sequences in multiple sequence alignments. *Bioinformatics* **20**: 2317–2319.
- Ingalls AE, Shah SR, Hansman RL, Aluwihare LI, Santos GM, Druffel ERM *et al*. (2006). Quantifying archaeal community autotrophy in the mesopelagic ocean using natural radiocarbon. *Proc Natl Acad Sci USA* **103**: 6442–6447.
- Jost G, Martens-Habben W, Pollehne F, Schnetger B, Labrenz M. (2010). Anaerobic sulfur oxidation in the absence of nitrate dominates microbial chemoautotrophy beneath the pelagic chemocline of the eastern Gotland Basin, Baltic Sea. *FEMS Microbiol Ecol* **71**: 226–236.
- Karner MB, DeLong EF, Karl DM. (2001). Archaeal dominance in the mesopelagic zone of the Pacific Ocean. *Nature* **409**: 507–510.
- Könneke M, Bernhard AE, de la Torre JR, Walker CB, Waterbury JB, Stahl DA. (2005). Isolation of an autotrophic ammonia-oxidizing marine archaeon. *Nature* **437**: 543–546.
- Labrenz M, Jost G, Jürgens K. (2007). Distribution of abundant prokaryotic organisms in the water column of the central Baltic Sea with an oxic-anoxic interface. *Aquat Microb Ecol* **46**: 177–190.
- Labrenz M, Jost G, Pohl C, Beckmann S, Martens-Habben W, Jürgens K. (2005). Impact of different *in vitro* electron donor/acceptor conditions on potential chemolithoautotrophic communities from marine pelagic redoxclines. *Appl Environ Microbiol* **71**: 6664–6672.
- Lam P, Jensen MM, Lavik G, McGinnis DF, Müller B, Schubert CJ *et al*. (2007). Linking crenarchaeal and bacterial nitrification to anammox in the Black Sea. *Proc Natl Acad Sci USA* **104**: 7104–7109.
- Lane DJ. (1991). 16S/23S rRNA sequencing, Vol. John Wiley and Sons: Chichester.
- Lass HU, Prandke H, Liljebladh B. (2003). Dissipation in the Baltic proper during winter stratification. *J Geophys Res-Oceans* **108**: doi:10.1029/2002JC001401.
- Leclerc M, Delbes C, Moletta R, Godon J-J. (2001). Single strand conformation polymorphism monitoring of 16S rDNA Archaea during start-up of an anaerobic digester. *FEMS Microbiol Ecol* **34**: 213–220.
- Lepland A, Stevens RL. (1998). Manganese authigenesis in the Landsort Deep, Baltic Sea. *Mar Geol* **151**: 1–25.
- Lin X, Wakeham SG, Putnam IF, Astor YM, Scranton MI, Chistoserdov AY *et al*. (2006). Comparison of vertical distributions of prokaryotic assemblages in the anoxic Cariaco Basin and Black Sea by use of fluorescence *in situ* hybridization. *Appl Environ Microbiol* **72**: 2679–2690.
- Ludwig W, Strunk O, Westram R, Richter L, Meier H, Yadhukumar *et al*. (2004). ARB: a software environment for sequence data. *Nucleic Acids Res* **32**: 1363–1371.
- Martens-Habben W, Berube PM, Urakawa H, de la Torre JR, Stahl DA. (2009). Ammonia oxidation kinetics determine niche separation of nitrifying Archaea and Bacteria. *Nature* **461**: 976–979.
- Massana R, DeLong EF, Pedros-Alio C. (2000). A few cosmopolitan phylotypes dominate planktonic archaeal assemblages in widely different oceanic provinces. *Appl Environ Microbiol* **66**: 1777–1787.
- Muyzer G, de Waal EC, Uitterlinden AG. (1993). Profiling of complex microbial populations by denaturing gradient gel electrophoresis analysis of polymerase chain reaction-amplified genes coding for 16S rRNA. *Appl Environ Microbiol* **59**: 695–700.
- Neretin LN, Pohl C, Jost G, Leipe T, Pollehne F. (2003). Manganese cycling in the Gotland Deep, Baltic Sea. *Mar Chem* **82**: 125–143.
- Pernthaler A, Pernthaler J, Amann R. (2002). Fluorescence *in situ* hybridization and catalyzed reporter deposition for the identification of marine bacteria. *Appl Environ Microbiol* **68**: 3094–3101.
- Pöhler I, Wenderoth DF, Wendt-Potthoff K, Höfle MG. (2002). Bacterioplankton community structure and dynamics in enclosures during bio-remediation

- experiments in an acid mining lake. *Water, Air, and Soil Poll: Focus 2*: 111–121.
- Pruesse E, Quast C, Knittel K, Fuchs BM, Ludwig W, Peplies J *et al.* (2007). SILVA: a comprehensive online resource for quality checked and aligned ribosomal RNA sequence data compatible with ARB. *Nucl Acids Res* **35**: 7188–7196.
- Purkhold U, Pommerening-Röser A, Juretschko S, Schmid MC, Koops HP, Wagner M. (2000). Phylogeny of all recognized species of ammonia oxidizers based on comparative 16S rRNA and *amoA* sequence analysis: implications for molecular diversity surveys. *Appl Environ Microbiol* **66**: 5368–5382.
- Reissmann J, Burchard H, Feistel R, Hagen E, Lass HU, Mohrholz V *et al.* (2009). Vertical mixing in the Baltic Sea and consequences for eutrophication—A review. *Progr Oceanogr* **82**: 47–80.
- Rheinheimer G, Gocke K, Hoppe H-G. (1989). Vertical distribution of microbiological and hydrographic-chemical parameters in different areas of the Baltic Sea. *Mar Ecol Progr Ser* **52**: 55–70.
- Rotthauwe JH, Witzel KP, Liesack W. (1997). The ammonia monooxygenase structural gene *amoA* as a functional marker: molecular fine-scale analysis of natural ammonia-oxidizing populations. *Appl Environ Microbiol* **63**: 4704–4712.
- Schleper C, Jurgens G, Jonuscheit M. (2005). Genomic studies of uncultivated archaea. *Nature Rev Microbiol* **3**: 479–488.
- Schubotz F, Wakeham SG, Lipp JS, Fredricks HF, Hinrichs K-U. (2009). Detection of microbial biomass by intact polar membrane lipid analysis in the water column and surface sediments of the Black Sea. *Environ Microbiol* **11**: 2720–2734.
- Schwieger F, Tebbe CC. (1998). A new approach to utilize PCR-single-strand-conformation polymorphism for 16S rRNA gene-based microbial community analysis. *Appl Environ Microbiol* **64**: 4870–4876.
- Sekar R, Pernthaler A, Pernthaler J, Warnecke F, Posch T, Amann R. (2003). An improved protocol for quantification of freshwater *Actinobacteria* by fluorescence *in situ* hybridization. *Appl Environ Microbiol* **69**: 2928–2935.
- Stackebrandt E, Ebers J. (2006). Taxonomic parameters revisited: tarnished gold standards. *Microbiol Today* **33**: 152–155.
- Teira E, Reinthaler T, Pernthaler A, Pernthaler J, Herndl GJ. (2004). Combining catalyzed reporter deposition-fluorescence *in situ* hybridization and microautoradiography to detect substrate utilization by bacteria and archaea in the deep ocean. *Appl Environ Microbiol* **70**: 4411–4414.
- Teira E, van Aken H, Veth C, Herndl GJ. (2006). Archaeal uptake of enantiomeric amino acids in the meso- and bathypelagic waters of the North Atlantic. *Limnol Oceanogr* **51**: 60–69.
- Treusch AH, Leininger S, Kletzin A, Schuster SC, Klenk H-P, Schleper C. (2005). Novel genes for nitrite reductase and Amo-related proteins indicate a role of uncultivated mesophilic crenarchaeota in nitrogen cycling. *Environ Microbiol* **7**: 1985–1995.
- Venter JC, Remington K, Heidelberg JF, Halpern AL, Rusch D, Eisen JA *et al.* (2004). Environmental genome shotgun sequencing of the Sargasso Sea. *Science* **304**: 66–74.
- Vetriani C, Tran HV, Kerkhof LJ. (2003). Fingerprinting microbial assemblages from the oxic/anoxic chemocline of the Black Sea. *Appl Environ Microbiol* **69**: 6481–6488.
- Wakeham SG, Amann R, Freeman KH, Hopmans EC, Jørgensen BB, Putnam IF *et al.* (2007). Microbial ecology of the stratified water column of the Black Sea as revealed by a comprehensive biomarker study. *Org Geochem* **38**: 2070–2097.
- Weinbauer MG, Fritz I, Wenderoth DF, Höfle MG. (2002). Simultaneous extraction from bacterioplankton of total RNA and DNA suitable for quantitative structure and function analyses. *Appl Environ Microbiol* **68**: 1082–1087.
- Wuchter C, Abbas B, Coolen MJL, Herfort L, van Bleijswijk J, Timmers P *et al.* (2006). Archaeal nitrification in the ocean. *Proc Natl Acad Sci USA* **103**: 12317–12322.
- Yakushev E, Chasovnikov V, Murray J, Pakhomova S, Podymov O, Stunzhas P. (2008). Vertical hydrochemical structure of the Black Sea. In: Kostianoy AG, Kosarev AN (eds). *The Black Sea Environment*, Vol 5Q Springer: Berlin/Heidelberg, pp. 277–307.

Supplementary Information accompanies the paper on The ISME Journal website (<http://www.nature.com/ismej>)