

effects mainly through hypothalamic growth-hormone-releasing hormone (GHRH) and somatostatin-releasing-inhibiting factor (SRIF)⁷. Perhaps Grf-1-dependent regulation of growth hormone synthesis and release is modified in our *grf1* mutant mice by deregulation of the neuronal network in the hypothalamus. The molecular events that link muscarinic receptors and Ras proteins through Grf-1 (ref. 8) might also be involved in controlling the GHRH/SRIF balance in the hypothalamus and in memory consolidation in the amygdala⁹.

Like other paternally expressed genes, *grf1* is involved in growth stimulation, whereas maternally expressed genes are responsible for growth suppression¹⁰. In contrast to these imprinted genes that are implicated in fetal growth, *grf1* is the first imprinted gene to be implicated exclusively in postnatal growth control.

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Miocene/Pliocene shift: one step or several?

Cerling *et al.*¹ provide evidence of a world-wide expansion of biomass of plants using the C₄ photosynthetic pathway about 8–6 million years (Myr) ago, recorded in the carbon isotope composition of the dental enamel of fossil herbivores. The authors claim that the expansion of C₄ plants is accompanied by a “worldwide faunal change” and infer that “an important global

ecological change was under way at this time”. For these statements to be true, the period of transition from C₃ to C₄ plants must prove to be contemporary with the period of faunal turnover, and both events must have taken place in the same geographic areas. At least, this is what the correlation of faunal and vegetation events provided by Cerling *et al.* seems to indicate.

But this model lacks a detailed analysis of extinction/immigration rates in the framework of precise palaeomagnetic and biostratigraphical data. In this context, a comparison of the late Miocene–early Pliocene records of artiodactyl and rodent communities from Spain and Pakistan is of special interest. First, the amount of palaeomagnetic and biostratigraphical data from both areas^{2–8} provides a high resolution of the shifts in faunal communities, and second, according to Cerling *et al.*¹, the C₃/C₄ transition took place all over the world except in western Europe.

Our comparison of extinction and immigration rates from sites in Spain and Pakistan (Fig. 1) reveals that first, in Pakistan's Siwalik sediments, not one but two faunal changes occurred, affecting especially artiodactyls. The first one, at about 10–9 Myr ago, largely preceded the C₃/C₄ transition (8–6 Myr ago); the second one occurred at its end (6.5 Myr). However, none of them exactly coincides with the initial expansion of C₄ biomass (8–7 Myr ago), despite a slight increase in rodent extinctions.

Second, both faunal changes in the Siwaliks are contemporary with faunal turnovers in Spain, although in western Europe there is no indication of C₄ diet at any time. In Spain, these faunal changes turn out to have been even more dramatic than in the Siwaliks as the extinction/immigration rate is twice as high. In Spain, woodland or forest indicators (hominoids and tragulids) became extinct about 9 Myr ago, whereas in the Siwaliks hominoids survived their European counterparts by about 1 Myr, and tragulids by as much as 5 Myr.

In the light of these data, we cannot find synchronicity or even a causal link between faunal and vegetation change. Faunal turnovers of different magnitudes took place before and after the appearance of C₄ plants, suggesting that C₃/C₄ transition and faunal turnovers occurred independently.

The “important global ecological change” instead looks like a succession of faunal and vegetation changes, scattered over a long period of at least 5 Myr. We feel that there is insufficient information to permit a clear interpretation of these phenomena. They probably occurred under the influence of a variety of events such as alpine orogenesis with Himalayan and Tibetan uplift, monsoonal dynamics, increasing seasonality, and other factors not yet fully investigated.

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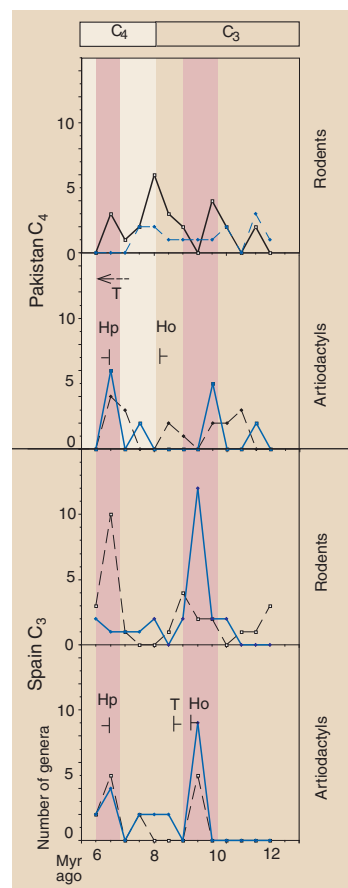


Figure 1 Faunal turnovers in the Eurasian Upper Miocene epoch under C₃ (Spain) and C₄ (Pakistan) conditions. Dotted lines, appearance of genera; continuous lines, disappearance of genera. Hp, hippopotamids; Ho, hominoids; T, tragulids. Data on Pakistan's Siwalik sediments derive from ref. 8; the palaeomagnetic scale is from ref. 9. Considered are herbivore communities with frequent changes at the generic level (artiodactyls and rodents); excluded are groups with low generic diversity such as proboscideans, rhinos and equids. Between 12 and 5.5 Myr ago, two faunal turnovers took place, the older one between 10 and 9 Myr ago and the younger one between 7 and 6 Myr ago. 60% of all faunal changes (immigrations/extinctions) that happened between 12 and 5.5 Myr ago occurred in these times of faunal turnover, when the rate of extinction/immigration increased up to four times the 'normal' value. With the exception of the extinction pattern of Siwalik rodents in the late Miocene epoch, there is a high correlation between the faunal turnover maxima of both the Spanish and the Siwalik herbivore communities, despite the transition from C₃ to C₄ plants at 8 Myr ago in Pakistan.

Cerling *et al.* reply — Köhler *et al.* suggest that phenomena other than floral change may be involved in the late Miocene global vegetation change, such as monsoonal dynamics or unnamed “other factors”. Citing evidence from Spain and Pakistan, they do not believe that there is necessarily a synchronicity or a causal link between faunal and vegetation change in the late Miocene epoch. However, on the contrary, it seems highly unlikely that a vegetation change on the scale documented¹ would be uncorrelated with faunal change.

Widespread faunal change in the late Miocene epoch was recognized^{7,10–12} long before the carbon-isotope shift was identified; our work was the first to attempt to link these widespread faunal changes to global vegetation change^{1,13}. For example, in North America, Webb *et al.*¹⁴ state that the “boundary between the Early and Late Hemphillian (about 6 Myr ago) records a mass extinction event for equids, when about ten of the existing 18 lineages vanished”. Although it is difficult to ‘prove’ causality in historical events, it seems likely that widespread faunal changes are linked to widespread vegetation changes.

The data from the Siwalik sediments in Pakistan are especially informative, because only from this region are there coeval data on faunal turnover, isotope palaeoecology, and upwelling related to monsoon dynamics (Fig. 1). Smoothed palaeosol data for carbon-13 content ($\delta^{13}\text{C}$) show a sharp change starting about 7 Myr ago and continuing to about 5 Myr ago, denoting the shift from C_3 - to C_4 -dominated vegetation.

The $\delta^{13}\text{C}$ data for tooth enamel show that the dietary change, which enhances the C_3 or C_4 signal by selective feeding, can be seen somewhat earlier than in the palaeosols, a result to be expected. Smoothed $\delta^{18}\text{O}$ data from palaeosols indicates a change in soil waters that precedes the $\delta^{13}\text{C}$ shift and which is correlated with increased abundance of upwelling indicators in the Arabian Sea at about 8.5 Myr ago.

Therefore the isotope record in the Siwaliks records two signals: a change in monsoonal dynamics at about 8.5 Myr ago and a pronounced vegetation change at about 7 Myr ago. Detailed faunal collections from the same region document several important turnover events. The two biggest events are at about 7 and 8.5 Myr ago (Fig. 1) and correspond to the two periods of change recorded in the isotope record.

Although the record is indeed complicated, the stable isotope record documents two important events affecting faunal change in the Siwaliks: one starting about 8.5 Myr ago that is related to the monsoon intensification, and a slightly later event related to expansion in C_4 biomass. Earlier faunal changes, such as those before 10 Myr ago as mentioned by Köhler *et al.*, are unre-

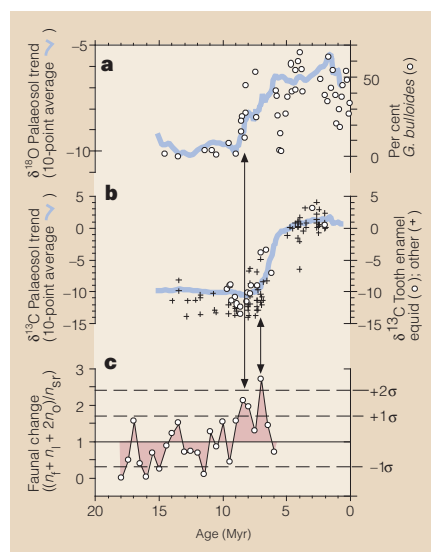


Figure 1 Data from Pakistan's Siwalik sediments show the two biggest events occurring at about 7 and 8.5 Myr ago. **a**, $\delta^{18}\text{O}$ data from Siwalik palaeosols, representing a trend determined by taking a 10-point running average of the roughly 200 palaeosols from the interval 16 to 0 Myr ago¹⁶. Also shown is the fraction of *Globigerina bulloides* from the Arabian Sea, an indicator of upwelling related to monsoon dynamics¹⁷. **b**, $\delta^{13}\text{C}$ data for palaeosols and for mammals' tooth enamel^{18–20} in the Siwaliks, representing a trend determined by taking a 10-point running average of the 200 or so palaeosols from the interval 16 to 0 Myr ago¹⁶. **c**, Faunal change index from the Siwaliks, represented by the number of first (n_1) and last (n_l) occurrences, including only occurrences (n_o), normalized to species richness (n_s). Data from ref. 7. The index is normalized to 1.0 for the total data set.

lated to the global expansion of C_4 biomass.

C_4 photosynthesis is an adaptation to low atmospheric CO_2 levels. Because CO_2 gain and water loss both occur through stomata in C_3 plants, we expect that C_3 plants adapted to aridity would prosper in periods of lower atmospheric CO_2 . We would therefore expect that global changes within C_3 flora accompanied the C_4 expansion at the end of the Miocene epoch. Changes within C_3 ecosystems can be related to changes in atmospheric CO_2 levels (for example, the Pleistocene/Holocene transition¹⁵).

So, although C_4 plants did not flourish in Europe or in other high-latitude regions, it is likely that floral change occurred in those regions within C_3 ecosystems through the Miocene/Pliocene transition. The absence of evidence for C_4 expansion in Europe should not be taken to mean that floral change did not take place in Europe at the end of the Miocene; the isotope record is silent on that issue.

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Life-support system benefits from noise

Mechanical ventilators are used to provide life support for patients with respiratory failure. But over the long term, these machines can damage the lungs, causing them to collapse and the partial pressure of oxygen in the arteries to drop to abnormally low values¹. In conventional mechanical ventilation, the respiratory rate and volume of air inspired per breath are fixed, although during natural breathing these parameters vary appreciably². A computer-controlled ventilator has now been introduced³ that can use noise to mimic this variability. We describe a conceptual model of lung injury in which the partial pressure of arterial oxygen is improved significantly by computer-controlled rather than conventional mechanical ventilation, in agreement with recent experimental data³.