stimulated by the occurrence of a layered stratiform basic intrusive, within the Yilgarn Block.

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## Estimation of Nuclear Explosion Energies from Microbarograph Records

FOLLOWING the US and USSR atmospheric test series in 1954-1962, numerous microbarograph records<sup>1-8</sup> of air waves generated by nuclear bomb tests were published. Previous theoretical interpretations<sup>7,9</sup> of such waveforms have required some explicit knowledge of the average atmospheric temperature and wind profiles above the path connecting source to microbarograph. Such profiles are never sufficiently well known and vary from point to point, and as seemingly small changes in the profiles cause relatively large changes in the waveforms, it would seem to be difficult to estimate the explosion energy yield to even order of magnitude accuracy from such records. Recently, however, in a further account of this work to be published elsewhere, we have succeeded in deriving an approximate theoretical relationship between certain waveform features and energy yield which is insensitive to changes in atmospheric structure. This relationship is given by

$$E = 13 p_{\rm FPT} \left[ r_{\rm e} \sin(r/r_{\rm e}) \right]^{\frac{1}{2}} H_{\rm s} \left( c T_{1,2} \right)^{3/2} \tag{1}$$

where E is energy release,  $p_{FPT}$  is the first peak to trough pressure amplitude (see Fig. 1),  $r_e$  is radius of the Earth, r is the great circle distance from burst point to observation point,  $H_s$ is a lower atmosphere scale height, c is a representative sound speed, and  $T_{1,2}$  is the time interval between first and second peaks. The purpose of the present communication is to describe the extent to which the above relation agrees with the existing available data.

The various points shown in Fig. 1 correspond to individual microbarograms recorded at Pasadena, California; Berkeley, California; Terceiva, Azores; Fletcher's Ice Island; Whippany, New Jersey; Ewa Beach, Hawaii, and Palisades, New York, after the Soviet explosions of (a) September 10 (10 MT), (b) September 11 (9 MT), (c) September 14 (7 MT), (d) October 4 (8 MT), (e) October 6 (11 MT), (f) October 20 (5 MT), (g) October 23 (25 MT), (h) October 30 (58 MT), and (i) October 31, 1961 (8 MT) and the US explosions of (j) May 4 (3 MT), (k) June 10 (9 MT), (l) June 12 (6 MT), (m) June 27 (24 MT), and (n) July 11, 1962 (12 MT). Here the estimate of the yield (in equivalent megatons of TNT where one MT equals  $4.2 \times 10^{22}$  ergs) is taken from Båth<sup>10</sup>. All the records used are taken from the articles of Harkrider<sup>7</sup> and of Donn and Shaw<sup>8</sup>. Pressure amplitudes for Harkrider's records were computed using his microbarograph response data. Pressure amplitudes for the Donn and Shaw records were determined according to the premises (W. Donn, private communication) that (a) all records recorded by Lamont type A microbaro-





Fig. 1 Comparison of data with the theoretical relationship between amplitude and period of infrasonic waveforms generated by nuclear explosions. The data points are lettered a to ncorresponding to particular events defined in the text. Donn and Shaw;  $\triangle$ , Harkrider.

graphs are to the same scale and (b) the clip to clip amplitude of off scale oscillations was 350 µbars. The ordinate in Fig. 1 gives  $p_{\rm FPT}/[Y\sin^{-\frac{1}{2}}(r/r_{\rm e})]$  in µbar MT<sup>-1</sup> where Y is the explosion yield in MT. The abscissa gives the period  $T_{1,2}$  in s. Note that the plot is full logarithmic. The solid line represents the theoretical relation, equation (1) with c and  $H_s$  taken as 310 m s<sup>-1</sup> and 8 km, respectively.

The scatter about the theoretical curve could be due to various causes; one which seems especially likely is the undulation in amplitude due to the horizontal refraction and subsequent focusing or defocusing caused by departures of the atmosphere from perfect stratification. We may note also that much of the scatter would not be present if we had omitted data corresponding to explosions of greater than 11 MT. The general trend of longer period signals being of lower amplitudes than signals recorded elsewhere but which were generated by the same event seems to be amply substantiated by the data.

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