the latter part of Lower Carboniferous times. Following a withdrawal of the sea, an extensive series of terrestrial beds was deposited in this area in Middle and Upper Carboniferous times; these terrestrial strata consist mainly of conglomerates, volcanic rocks (lavas and tuffs), and glacial beds, of an aggregate thickness approaching 10,000 feet. The glacial beds are of such a thickness and volume as to imply intense and long-continued glaciation. Associated with these beds is a characteristic Rhacopteris fossil flora.

In Permian (Permo-Carboniferous) time an alternating series of marine and freshwater beds was deposited in the north-eastern part of New South Wales, and these extend far northwards into eastern Queensland. The freshwater beds contain the most productive coal-measures of Australia, and associated with the coal-seams is the characteristic Glossopteris flora. The glacial conditions of the Carboniferous Period continued also far into the Permian Period, but with apparently reduced intensity. The Permian Period closed in north-eastern New South Wales and south-eastern Queensland with pronounced orogenic movements, accompanied by granitic intrusions; but elsewhere in New South Wales and throughout Victoria and Tasmania, no earth-folding took place at that time.

In the Trias-Jura Period the whole of eastern Australia stood above the sea, and extended far east of the present shore-line. Upon this land there developed a number of large lake-basins in which several thousands of feet of freshwater strata were deposited, and in some areas productive coal-measures were formed. In the Cretaceous Period a transgression of the sea began in the north, and extended southwards over central Queensland into northern New South Wales, and well into central Australia. At the beginning of the Tertiary Period a tilting of the Australian continent on an east-and-west axis caused the Cretaceous sea to retreat northwards, and allowed of transgression taking place over considerable areas in the south, incidentally separating Tasmania from the mainland.

The close of the Tertiary Period was marked by a great epeirogenic uplift in eastern Australia, which produced the existing tablelands trending parallel to the eastern coast of Australia. The elevation of these tablelands was accompanied by extensive block-faulting. During the Pleistocene Period, limited high areas in New South Wales and Tasmania supported glaciers and ice-sheets; more recently, a subsidence of the land (or raising of sea-level) drowned the shore-line to an extent of about 200 feet, and still later an upward movement of the strand-line of some 10 to 20 feet has taken place.

## The Japanese Earthquake of September 1, 1923.

SHORTLY after the Japanese earthquake of September 1, 1923, Mr. Takeo Kato, on behalf of the Imperial Earthquake Investigation Committee, made reconnaissances through the districts of violent shocks, especially those around Sagami Bay. He has published a preliminary report on this field study in the Journal of the Geological Society of Tokyo (vol. 30, No. 361), including some definite estimates of the loss of life and property. In Tokyo, though not a house escaped some damage, the number of houses that collapsed owing to the shock was comparatively small. The latest official statistics place the number of houses destroyed by fire in Tokyo at 316,087, the number of killed at 67,052, of injured at 32,583, and of missing at 38,980. The shocks were far stronger at Yokohama, and strongest of all in the districts around Sagami Bay. In six towns along and near the borders of the Bay, 84,300 houses were destroyed and 26,370 lives were lost.

From seismographic and other evidence, the epicentre of the great earthquake seems to lie in Sagami Bay, about midway between Oshima and Hiratsuka, near the mouth of the Sagami River, and probably a little nearer the latter place. In the two months following this shock, the number of after-shocks was more than 1350; there were 365 on September 1, and 289, 173, and 143 during the next three days; afterwards the normal decline in frequency was manifested. The epicentres of these after-shocks lay at different places more or less remote from the main one. A few minutes after the great earthquake sea-waves swamped the shores of Sagami Bay, from Shimoda in the Idzu peninsula to Misaki and Uraga in the Miura peninsula, and also along the southern part of the Boso peninsula. They caused much damage only at Ito, Atami, Kamakura, and a few other places, where the height of the waves ranged from 20 to 40 feet. No conspicuous sea-waves were seen in Tokyo Bay. No great fault-scarps or traces have been observed, but in the islet of Hatsu-shima, off Atami, a fissure was traced for about half a mile, with in some places a displacement of 3 feet or more, running in the direction N. 30° W. Another, rather more than a mile in length, was found near Nagasawa in the Miura peninsula. Both fissures may be regarded as minor faults of the main fault along which the great earthquake originated.

A remarkable upheaval of the ground took place, it is said, within a few minutes after this earthquake, over an area of about 2000 square miles, including the islet of Hatsu-shima, the headland of Manazuru, the northern border of Sagami Bay, the Miura and Boso peninsulas. The greatest upheaval (about 8 feet) occurred at Tomizaki, along the southern coast of the Boso peninsula, and it is remarkable that here the ground settled down more than a foot within one month after the upheaval. According to the soundings made by the Navy Hydrographic Office, conspicuous changes in depth, of 50 fathoms and more, have occurred at various places in the deepest portion of Sagami Bay. Soundings are still being carried out in the Bay, while many levelling parties of the Military Department are engaged in ascertaining the changes in elevation throughout the earthquake area.

It is worthy of notice that the epicentre lies in the deepest region of Sagami Bay. This trough, which has a N.N.W. direction from Oshima to Hiratsuka, appears to be the southern continuation of the valley of the Sagami River running in the same direction. Moreover, tectonic lines in this direction are numerous, as indicated by the coast borders, valley courses, etc., and the minor faults at Hatsu-shima and Nagasawa run nearly parallel to this direction. The origin of the earthquake is attributed to a great compressive force, probably from the east, which has been accumulating for a long time. The last great earthquake in the district occurred on November 22, 1703, and seems to have originated in a great fault along the same tectonic line, the epicentre lying apparently a little to the south of that in 1923. A similar upheaval of land is recorded at that time.

C. D.