

hood of the site. The galactic latitude of the object is $b_{\text{II}} \approx -20^\circ$, so it is consistent with the argument that a supernova remnant at high galactic latitude does not become (or cannot last) as a strong radio source.

(3) The event of AD 437: The star was visible during daylight before 3 p.m. It was orange-red in colour and was the size of an orange. The radio source IC 443 was considered by Shklovsky and Shajn as the remnant of a nova of AD 837, but it is also within the error circle of the event of AD 437.

(4) The event of AD 902: "The star was as big as a peach", that is, an area source of light. It was visible for more than 1 yr. Pskovskii¹⁰ has given a detailed discussion of this object and proposed nebula CTA-1 as the supernova's radio remnant.

Finally, the distance to Vela X is about 500 pc, suggesting a brightness of $m_v < -8$ at explosion, but no suspicious event near the location of the radio source was recorded in Chinese history dating back to 206 BC (the beginning of Han dynasty), and with less certainty to 1000 BC. (According to Needham¹¹ the oldest Chinese record of a nova was inscribed on an oracle-bone, dating from about -1300.) The age of the remnant is therefore most likely older than the five events considered here. The period of the pulsar associated with it is 0.0892 s. This may be an indication of the expected period of historical pulsars.

Part of this work was carried out while I was a visiting lecturer at the Summer Science Institute sponsored by Academia Sinica of the Republic of China. I thank Dr C. Y. Hsu, professor of Chinese history at Taiwan University, for many helpful conversations. The work was supported by a US National Aeronautics and Space Administration grant at Purdue University.

C. S. SHEN

Department of Physics,
Purdue University,
Lafayette, Indiana.

Received January 6, 1969.

¹ Large, M. I., Vaughan, A. E., and Mills, B. Y., *Nature*, **220**, 340 (1968).

² Schonfeld, E., *Astron. Nachr.*, **127**, 153 (1891).

³ Goldstein, B. R., *Astron. J.*, **70**, 105 (1965); Marsden, B. C., *ibid.*, **70**, 126 (1965); Goldstein, B. R., and Ho pen Yoke, *ibid.*, **70**, 738 (1965); Gardner, F. E., and Milne, D. K., *ibid.*, **70**, 754 (1965); Minkowski, R., *ibid.*, **71**, 371 (1966).

⁴ Duyvendak, J. J. L., *Publ. Astron. Soc. Pacific*, **54**, 91 (1945).

⁵ Minkowski, R., *Ann. Rev. Astron. Astrophys.*, **2**, 247 (1964).

⁶ Mills, B. Y., Alec, O. B., and Hill, E. R., *Austral. J. Phys.*, **13**, 676 (1960).

⁷ Poveda, A., and Woltjer, L., *Astron. J.*, **73**, 65 (1968).

⁸ Ho pen Yoke, *Vistas in Astronomy*, **5**, 127 (1962).

⁹ Hsi Tse-Tsung, *Smithson. Contr. Astrophys.*, **2**, 109 (1958).

¹⁰ Pskovskii, Y. P., *Sov. Astron.*, **7**, 501 (1964).

¹¹ Needham, J., *Science and Civilization in China*, **3** (Cambridge Univ. Press, 1959).

Concentration of CO₂ in the Upper Troposphere and Lower Stratosphere

DURING a number of flights of commercial aircraft over the polar route from Frankfurt/Main to Tokyo via Copenhagen the concentration of CO₂ in the atmosphere was continuously recorded using a non-dispersive infrared gas analyser. The flights took place between April 1967 and February 1968. The recorded concentration of CO₂ from all flights shows a remarkable change when the aircraft enters and leaves the stratosphere. A quite sudden change of about 3–5 p.p.m. CO₂ is usually observed at the transition from tropospheric into stratospheric air and vice versa. The striking difference between the CO₂ content of the upper troposphere and the stratosphere was clear enough on the record for the crew of the aircraft to use our CO₂ analyser as an indicator of the position of the

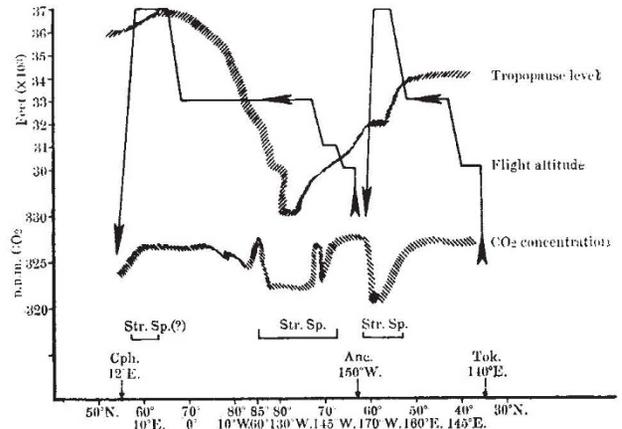


Fig. 1. Transpolar flight Tokyo-Frankfurt, May 29–30, 1967.

aircraft relative to the tropopause. Fig. 1 shows the record obtained during the flight from Tokyo to Copenhagen on May 29–30, 1967, together with the position of the aircraft and the tropopause. Sections of the flight within the stratosphere are marked. They clearly show a drop of the CO₂ content in the stratosphere. The sudden peak at 70°N 145°W during the flight through the stratosphere may be explained by influx of tropospheric air into the stratosphere. The relative accuracy of our measurements is about ± 0.3 p.p.m. CO₂, which gives confidence in the reality of the observed fluctuations. The accuracy of the absolute values of the CO₂ concentration amounts to ± 1.5 p.p.m. due to difficulty in calibrating the reference gases. The marked decrease of the CO₂ concentration in the lower stratosphere compared with the upper troposphere suggests that, contrary to previous practice, it is wrong to assume a constant mixing ratio of CO₂ in the troposphere and stratosphere. Exact knowledge of the CO₂ concentration with altitude is needed for many calculations of the radiation budget of the atmosphere.

Even considering the annual increase in the global CO₂ concentration of 0.7 to 1 p.p.m., our values seem rather high compared with discontinuous CO₂ analyses of air samples from the upper troposphere and lower stratosphere by Bischof and Bolin¹. Agreement is better with the results of a balloon flight by Steinberg and Rohrbough², who collected two air samples at an altitude of 21 km having concentrations of CO₂ of 323 p.p.m. and 338 p.p.m. These authors used an adsorption method giving an accuracy in the CO₂ values of only ± 5 per cent. Further support of our findings comes from some recent (and as yet unpublished) results of Murreray (private communication by F. Möller) which show a rather sharp drop of the CO₂ concentration between 15 and 20 km. More flights with continuous records of the concentration of CO₂ are required to investigate the difference between the CO₂ content of the upper troposphere and lower stratosphere in more detail and in relation to seasonal variations of tropospheric CO₂. Nevertheless the results of the six transpolar flights carried out up to now may have further implications on many theoretical studies of the radiation and heat budget.

We thank Deutsche Lufthansa German Airlines for their support.

H. W. GEORGI
D. JOST

Department of Meteorology and Geophysics,
University of Frankfurt,
Germany.

Received November 25, 1968.

¹ Bischof, W., and Bolin, B., *Tellus*, **18**, 155 (1966).

² Steinberg, S., and Rohrbough, S. F., *J. Appl. Met.*, **1**, 418 (1962).