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Nuclear volume effects in equilibrium stable isotope fractionations of mercury, thallium and lead

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The nuclear volume effects (NVEs) of Hg, Tl and Pb isotope systems are investigated with careful evaluation on quantum relativistic effects via the Dirac's formalism of full-electron wave function. Equilibrium $^{202}\text{Hg}/^{198}\text{Hg}$, $^{205}\text{Tl}/^{203}\text{Tl}$, $^{207}\text{Pb}/^{206}\text{Pb}$ and $^{208}\text{Pb}/^{206}\text{Pb}$ isotope fractionations are found can be up to 3.61‰, 2.54‰, 1.48‰ and 3.72‰ at room temperature, respectively, larger than fractionations predicted by classical mass-dependent isotope fractionations theory. Moreover, the NVE can cause mass-independent fractionations (MIF) for odd-mass isotopes and even-mass isotopes. The plot of $\Delta_{\text{NV}}^{199}\text{Hg}$ vs. $\Delta_{\text{NV}}^{201}\text{Hg}$ for Hg-bearing species falls into a straight line with the slope of 1.66, which is close to previous experimental results. For the first time, Pb^{4+} -bearing species are found can enrich heavier Pb isotopes than Pb^{2+} -bearing species to a surprising extent, e.g., the enrichment can be up to 4.34‰ in terms of $^{208}\text{Pb}/^{206}\text{Pb}$ at room temperature, due to their NVEs are in opposite directions. In contrast, fractionations among Pb^{2+} -bearing species are trivial. Therefore, the large Pb fractionation changes provide a potential new tracer for redox conditions in young and closed geologic systems. The magnitudes of NVE-driven even-mass MIFs of Pb isotopes (i.e., $\Delta_{\text{NV}}^{204}\text{Pb}$) and odd-mass MIFs (i.e., $\Delta_{\text{NV}}^{207}\text{Pb}$) are almost the same but with opposite signs.

With rapid progresses in mass-spectrometer, great interests on stable isotope fractionations of heavy elements have been aroused. Evidences showed that heavy elements could have surprising isotopic fractionations as the consequence of the NVE^{1–9}. The NVE is originated from differences in nuclear size and nuclear shape of isotopes^{2,10}. It doesn't belong to the well-known driving forces of equilibrium isotope fractionation, which are governed by the conventional Bigeleisen-Mayer theory¹¹ or Urey method¹².

The concept of NVE was proposed in spectrometric studies¹⁰. However, those early studies have not investigated its influences on isotopic fractionations. Fujii *et al.*¹³ found anomalous isotope fractionations in uranium isotope exchange experiments which violated the Bigeleisen-Mayer equation (or Urey model) but suggested the cause to be the difference in nuclear spin. Nishizawa *et al.*¹ correctly interpreted the anomalous isotope effects of strontium by using isotope shift in atomic spectra (field shift). It was probably the first isotope NVE study. Hereafter, Bigeleisen² and Nomura *et al.*³ independently recognized that those anomalous isotope fractionation phenomena, which were caused by the NVE, could lead to large stable isotope fractionations of heavy elements. Bigeleisen² accordingly added the NVE as an important contribution into a modified calculation formula of the equilibrium isotope fractionation factor. He pointed out that the NVE is only a second order correction in chemical bonds, which suggested the NVE can have minor effect on vibrational frequencies¹⁴. Importantly, he emphasized that the NVE can change isotopic fractionation largely alone via the change of ground-state electronic energy.

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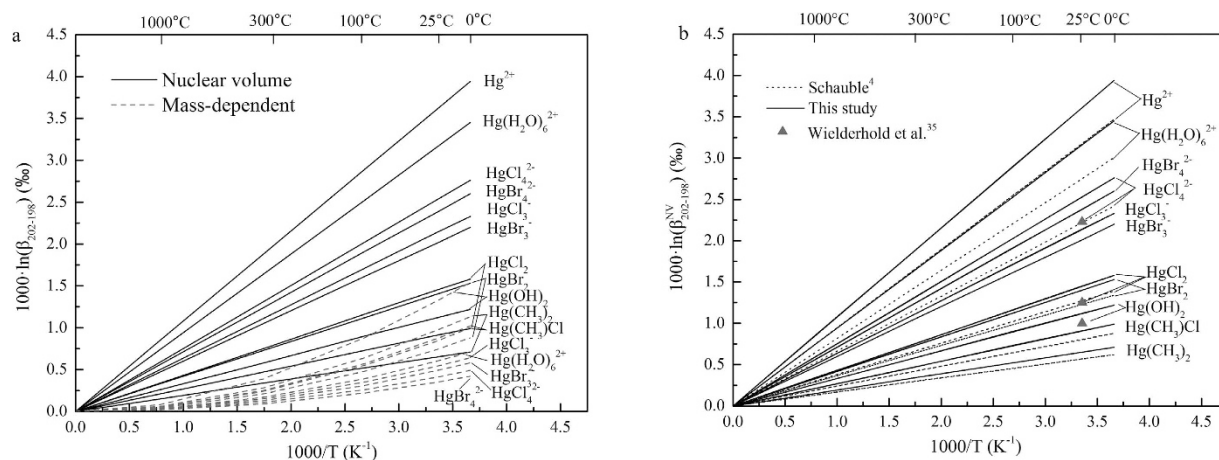


Figure 1. β -factors for Hg-bearing species relative to Hg⁰. (a) Contributions from CMDE and NVE, respectively. (b) Nuclear volume isotope fractionation factors ($\beta_{202-198}^{NV}$ -factors relative to Hg⁰ vapor) compared to the results of Schauble⁴ and Wiederhold *et al.*³⁵.

Toshiyuki Fujii and his co-workers have made tremendous efforts on experimental evaluations of NVEs for isotope systems, including Ti, Sn, Zr, Ni, Zn, Gd, Nd, Cr, Sr, Te and Cd etc.^{15–29} Meanwhile, they performed quantum chemistry calculations for a few isotope systems, such as Zn^{20,21}, Ni¹⁸, Tl³⁰ and Pb³¹. Schauble⁴ used quantum chemistry methods to calculate NVEs of some heavy elements (e.g., Hg, Tl) and showed that the NVE could affect isotope fractionations of heavy elements to surprising degrees. Then, Abe *et al.*^{5–7} independently calculated the NVE-driven fractionation factors of U-bearing species. Schauble³² developed a new method to model nuclear volume effects in crystals. His new method was based on density functional theory (DFT), using the projector augmented wave method (DFT-PAW) with a three-dimensional periodic boundary condition for greater speed and compatibility.

In addition, Zheng *et al.*³³ and Ghosh *et al.*³⁴ did different experiments to estimate the NVE of mercury isotopes in the absence of light. They both assigned those mass-independent isotope fractionation signals as the consequence of the NVE. Wiederhold *et al.*³⁵ also did experimental and theoretical investigations on Hg mass-independent isotope fractionations. Schauble⁴ and Wiederhold *et al.*³⁵ have explored small NVE-driven Hg isotope fractionations in organic Hg-bearing species in depth. Moynier *et al.*⁹ reviewed the NVEs of Tl and U isotope systems in different natural environments, such as under low- or high-temperature conditions and in meteorites. The necessity of careful NVE evaluation during the exploration of new heavy elements is recognized by most people.

Right now, there are a few different computational methods used to investigate quantum relativistic effects associated with the NVE, e.g., Schauble^{4,32} used the DIRAC and ABINIT software package, Abe *et al.*^{5–7} used a four-component relativistic atomic program package-GRASP2K, Fujii *et al.*^{18,20,21,30,31} used a software provided by Tokyo University (UTchem). Recently, Nemoto *et al.*³⁶ found a two-component relativistic method (the finite-order Douglas-Kroll-Hess method with infinite-order spin-orbit interactions for the one-electron term and atomic-mean-field spin-same-orbit interaction for the two-electron term, i.e., IODKH-IOSO-MFSO) with almost equivalent accuracy but 30 times faster than the previous four-component method by DIRAC software package. They also predicted the IODKH-IOSO-MFSO method could compute larger system for future NVE calculation.

Here we calculate the NVE-driven fractionation factors of Hg-, Tl- and Pb-bearing species by using full-electron quantum chemistry calculation methods. Our method is similar to that of Schauble⁴, in which quantum relativistic effects have been carefully evaluated via four-component Dirac equation formalism^{37,38}. Not only more new Hg- and Pb-bearing species (e.g., HgBr₄²⁻, HgCl₃⁻, HgBr₃⁻ and many Pb⁴⁺-bearing species) are calculated here, but more mass-independent fractionations are investigated in light of recent findings on even-number Hg isotope MIFs^{39,40}. Large fractionations (up to ca. 4‰ at room temperature) between Pb⁴⁺- and Pb²⁺-bearing species are found for the first time.

Results

Equilibrium stable isotope fractionations of Hg-, Tl-, and Pb-bearing species are shown in Figures 1, 2, and 3 and Tables 1, 2, and 3 relative to Hg⁰, Tl⁰, and Pb⁰ in terms of $1000 \cdot \ln\beta$, including conventional mass-dependent ($1000 \cdot \ln\beta^{MD}$) and nuclear volume effect fractionation factors ($1000 \cdot \ln\beta^{NV}$).

Hg isotope system. NVEs alone can fractionate ²⁰²Hg/¹⁹⁸Hg isotopes up to 3.61‰ at 25 °C. However, the largest classical mass-dependent fractionation are only 1.32‰ for ²⁰²Hg/¹⁹⁸Hg at 25 °C. All Hg-bearing species enrich heavier isotope (²⁰²Hg) relative to Hg⁰ vapor. The NVE-driven isotope fractionations of inorganic species, such as Hg²⁺, HgCl₄²⁻, HgBr₄²⁻, HgCl₃⁻, HgBr₃⁻, HgCl₂, HgBr₂, Hg(H₂O)₆²⁺ or

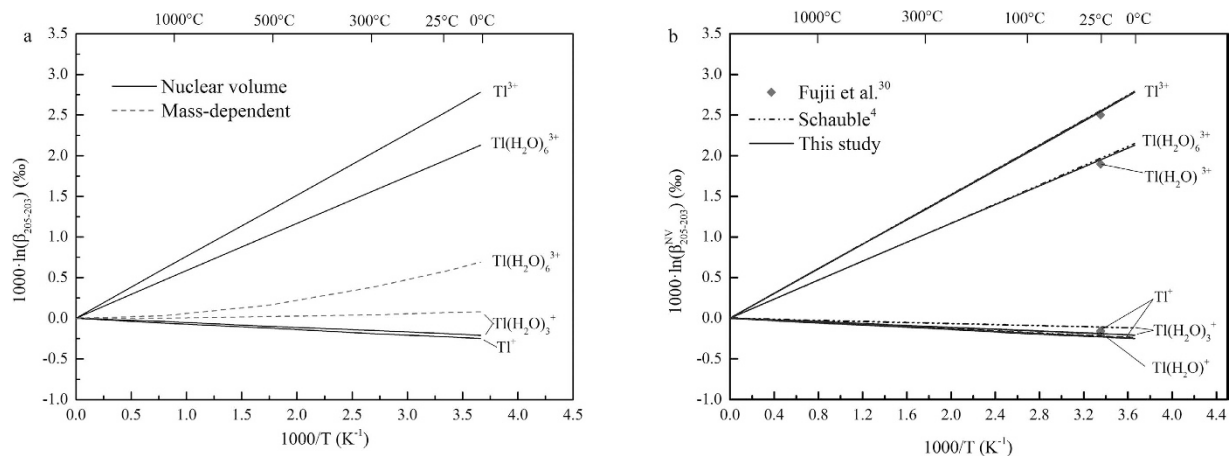


Figure 2. β -factors for Tl-bearing species relative to Tl⁰. (a) Contributions from CMDE and NVE, respectively. (b) Nuclear volume isotope fractionation factors ($\beta_{205-203}^{2525}$ -factors relative to Tl⁰) compared to the results of Schauble⁴ and Fujii *et al.*³⁰.

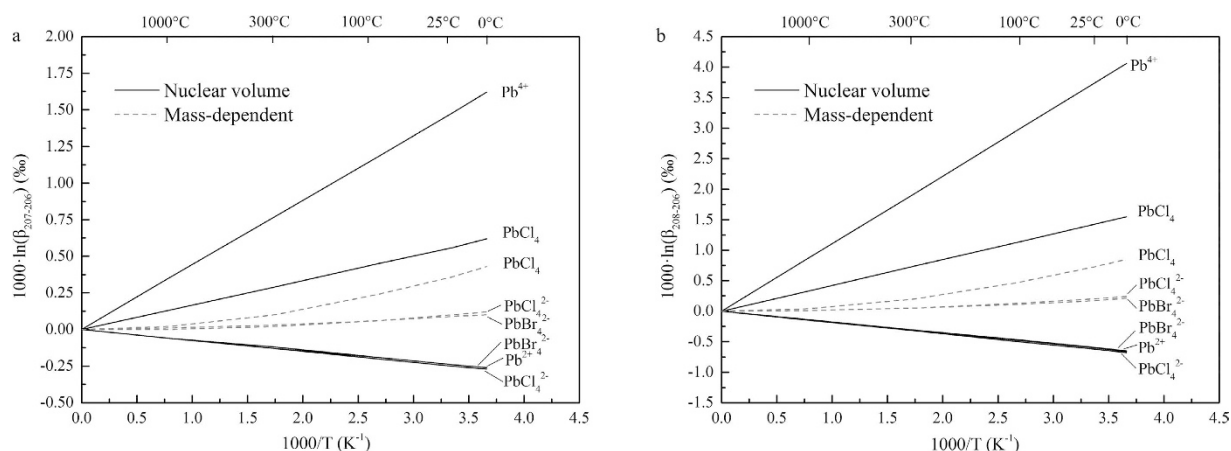


Figure 3. β -factors for Pb-bearing species relative to Pb⁰, including CMDE and NVE. (a) $\beta_{208-206}$ -factors of ²⁰⁷Pb/²⁰⁶Pb. (b) $\beta_{208-206}$ -factors of ²⁰⁸Pb/²⁰⁶Pb.

Hg(OH)₂, are larger than those of organic molecules (e.g., Hg(CH₃)Cl and Hg(CH₃)₂). On the contrary, CMDE fractionations of inorganic species are smaller than organic compounds except for Hg(OH)₂.

Tl isotope system. The NVE-driven fractionation of ²⁰⁵Tl/²⁰³Tl isotopes is up to 2.54‰ relative to Tl⁰ and the CMDE is only 0.58‰ for Tl(H₂O)₆³⁺ and 0.07‰ for Tl(H₂O)₃⁺ at 25 °C. Our NVE results show that Tl³⁺ ion and Tl³⁺-bearing compounds enriches heavier isotope (²⁰⁵Tl) relative to Tl⁰. However, Tl⁺ ion and Tl⁺-bearing compounds enriches lighter isotope (²⁰³Tl) compared to Tl⁰. Note that β -values of Tl⁺-bearing species are even smaller than the unity. This is because NVE tends to let heavier isotopes to be enriched in those atoms or ions with fewer s electrons or with more p, d and f electrons. Tl⁰ has more p electrons than Tl⁺-bearing species does.

Pb isotope system. NVEs induce Pb isotope fractionations up to 1.48‰ (²⁰⁷Pb/²⁰⁶Pb) and 3.72‰ (²⁰⁸Pb/²⁰⁶Pb) relative to Pb⁰, at 25 °C. However, contributions from classical mass-dependent fractionation are small, about 0.1–0.4‰ for ²⁰⁷Pb/²⁰⁶Pb and 0–2–0.7‰ for ²⁰⁸Pb/²⁰⁶Pb at 25 °C. The isotope fractionations of Pb⁴⁺-bearing species (e.g., PbCl₄) are larger than those of Pb²⁺-bearing species (e.g., PbCl₄²⁻ or PbBr₄²⁻) in terms of NVE or CMDE.

Mass-independent fractionation of Hg and Pb isotopes. Table 4 shows NVE-driven mass-independent fractionations for ¹⁹⁹Hg, ²⁰⁰Hg, and ²⁰¹Hg isotopes at room temperature. Those MIFs are relative to the MIF of Hg vapor (i.e., Hg⁰). If the real MIF value of a specific Hg-bearing species is needed, one needs convert the number listed in Table 4 via the aid of experimental MIF data of Hg vapor.

	This study ^{a)}					Schauble ^{b)}					Wiederhold <i>et al.</i> ^{c)}
	0 °C	25 °C	100 °C	300 °C	1000 °C	0 °C	25 °C	100 °C	300 °C	1000 °C	25 °C
CMDE fractionation factor ($1000 \cdot \ln \beta_{202-198}^{MD}$)											
Hg ⁰	0	0	0	0	0	0	0	0	0	0	
Hg ²⁺	0	0	0	0	0	0	0	0	0	0	
HgCl ₂	1.00	0.84	0.55	0.24	0.05	1.08	0.92 ± 0.1	0.60	0.26	0.05	0.84
HgBr ₂	0.88	0.74	0.48	0.21	0.04	0.95	0.80 ± 0.1	0.51	0.22	0.04	
Hg(CH ₃)Cl	1.02	0.87	0.57	0.25	0.05	1.04	0.89 ± 0.1	0.58	0.25	0.05	
Hg(CH ₃) ₂	1.13	0.97	0.64	0.28	0.06	1.14	0.97 ± 0.1	0.65	0.29	0.06	
HgCl ₃ ⁻	0.64	0.54	0.35	0.15	0.03						
HgCl ₄ ²⁻	0.49	0.41	0.26	0.11	0.02	0.67	0.56+0.7/-0.1	0.36	0.15	0.03	0.40
HgBr ₃ ⁻	0.58	0.49	0.31	0.13	0.03						
HgBr ₄ ²⁻	0.42	0.35	0.22	0.09	0.02						
Hg(H ₂ O) ₆ ²⁺	0.71	0.60	0.39	0.17	0.03	1.13	0.96 ± 0.4	0.62	0.27	0.05	
Hg(OH) ₂	1.54	1.32	0.88	0.39	0.08						1.19
NVE fractionation factor ($1000 \cdot \ln \beta_{202-198}^{NV}$)											
Hg ⁰	0	0	0	0	0	0	0	0	0	0	
Hg ²⁺	3.94	3.61	2.89	1.88	0.85	3.47	3.17 ± 0.6	2.54	1.65	0.74	
HgCl ₂	1.58	1.45	1.16	0.75	0.34	1.39	1.27 ± 0.3	1.02	0.66	0.30	1.25
HgBr ₂	1.53	1.40	1.12	0.73	0.33	1.34	1.23 ± 0.2	0.98	0.64	0.29	
Hg(CH ₃)Cl	0.99	0.91	0.73	0.47	0.21	0.88	0.80 ± 0.2	0.64	0.42	0.19	
Hg(CH ₃) ₂	0.71	0.65	0.52	0.34	0.15	0.62	0.57 ± 0.1	0.45	0.30	0.13	
HgCl ₃ ⁻	2.33	2.14	1.71	1.11	0.50						
HgCl ₄ ²⁻	2.76	2.53	2.02	1.31	0.59	2.42	2.22 ± 0.4	1.77	1.16	0.52	2.23
HgBr ₃ ⁻	2.20	2.01	1.61	1.05	0.47						
HgBr ₄ ²⁻	2.60	2.38	1.91	1.24	0.56						
Hg(H ₂ O) ₆ ²⁺	3.45	3.16	2.52	1.64	0.74	3.01	2.75 ± 0.6	2.20	1.43	0.64	
Hg(OH) ₂	1.22	1.12	0.90	0.58	0.26						1.00
Total fractionation factor ($1000 \cdot \ln \beta_{202-198}$)											
Hg ⁰	0	0	0	0	0	0	0	0	0	0	
Hg ²⁺	3.94	3.61	2.89	1.88	0.85	3.47	3.17 ± 0.6	2.54	1.65	0.74	
HgCl ₂	2.58	2.29	1.71	0.99	0.39	2.47	2.19 ± 0.3	1.62	0.92	0.35	2.09
HgBr ₂	2.41	2.14	1.60	0.94	0.37	2.29	2.03 ± 0.2	1.49	0.86	0.33	
Hg(CH ₃)Cl	2.01	1.78	1.30	0.72	0.26	1.92	1.69 ± 0.2	1.22	0.67	0.24	
Hg(CH ₃) ₂	1.84	1.62	1.16	0.62	0.21	1.76	1.54 ± 0.1	1.10	0.59	0.19	
HgCl ₃ ⁻	2.97	2.68	2.06	1.26	0.53						
HgCl ₄ ²⁻	3.25	2.94	2.28	1.42	0.61	3.09	2.78+0.8/0.4	2.13	1.31	0.55	2.63
HgBr ₃ ⁻	2.78	2.50	1.92	1.18	0.50						
HgBr ₄ ²⁻	3.02	2.73	2.13	1.33	0.58						
Hg(H ₂ O) ₆ ²⁺	4.16	3.76	2.91	1.81	0.77	4.14	3.71 ± 0.7	2.82	1.70	0.69	
Hg(OH) ₂	2.76	2.44	1.78	0.97	0.34						2.19

Table 1. Calculated stable isotope fractionation factors for Hg-bearing species relative to Hg⁰ (in per mil), including conventional mass-dependent effect (CMDE) fractionation factors ($1000 \cdot \ln \beta_{202-198}^{MD}$) and nuclear volume effect (NVE) fractionation factors ($1000 \cdot \ln \beta_{202-198}^{NV}$). ^aCalculated with $\langle r^2 \rangle$ values of Fricke and Heilig⁴¹ by using the software package DIRAC13.1. ^bCalculated with $\langle r^2 \rangle$ values of Angeli⁴² by using the software package DIRAC04 by Schauble⁴. ^cCalculated with $\langle r^2 \rangle$ values of Fricke and Heilig⁴¹ by using the software package DIRAC08 by Wiederhold *et al.*³⁵.

	This study					Schauble ^{a)}					Fujii <i>et al.</i> ^{b)}
	0°C	25°C	100°C	300°C	1000°C	0°C	25°C	100°C	300°C	1000°C	25°C
CMDE fractionation factor ($1000 \cdot \ln \beta_{205-203}^{MD}$)											
Tl ⁰	0	0	0	0	0	0	0	0	0	0	
Tl ⁺	0	0	0	0	0	0	0	0	0	0	
Tl ³⁺	0	0	0	0	0	0	0	0	0	0	
Tl(H ₂ O) ₃ ⁺	0.08	0.07	0.04	0.02	0.00	0.08	0.07 ± 0.1	0.04	0.02	0.00	0.063 ^{c)}
Tl(H ₂ O) ₆ ³⁺	0.69	0.58	0.38	0.16	0.03	0.77	0.65 ± 0.2	0.44	0.18	0.04	0.423
NVE fractionation factor ($1000 \cdot \ln \beta_{205-203}^{NV}$)											
Tl ⁰	0	0	0	0	0	0	0	0	0	0	
Tl ⁺	-0.25	-0.23	-0.19	-0.12	-0.06	-0.24	-0.22 ± 0.04	-0.18	-0.11	-0.05	-0.157
Tl ³⁺	2.78	2.54	2.03	1.32	0.60	2.79	2.55 ± 0.5	2.04	1.32	0.60	2.501
Tl(H ₂ O) ₃ ⁺	-0.21	-0.19	-0.15	-0.10	-0.04	-0.12	-0.11 ± 0.02	-0.09	-0.06	-0.03	-0.168 ^{d)}
Tl(H ₂ O) ₆ ³⁺	2.13	1.95	1.56	1.02	0.46	2.15	1.97 ± 0.4	1.57	1.02	0.46	1.898 ^{e)}
Total fractionation factor ($1000 \cdot \ln \beta_{205-203}$)											
Tl ⁰	0	0	0	0	0	0	0	0	0	0	
Tl ⁺	-0.25	-0.23	-0.19	-0.12	-0.06	-0.24	-0.22 ± 0.04	-0.18	-0.11	-0.05	
Tl ³⁺	2.78	2.54	2.03	1.32	0.60	2.79	2.55 ± 0.5	2.04	1.32	0.60	
Tl(H ₂ O) ₃ ⁺	-0.13	-0.12	-0.11	-0.08	-0.04	-0.04	-0.04 ± 0.1	-0.05	-0.04	-0.03	
Tl(H ₂ O) ₆ ³⁺	2.82	2.53	1.94	1.18	0.49	2.92	2.62 ± 0.4	2.01	1.20	0.50	

Table 2. Calculated stable isotope fractionation factors for Tl-bearing species relative to Tl⁰ (in per mil), including conventional mass-dependent effect (CMDE) fractionation factors ($1000 \cdot \ln \beta_{205-203}^{MD}$) and nuclear volume effect (NVE) fractionation factors ($1000 \cdot \ln \beta_{205-203}^{NV}$). ^aCalculated by Schauble⁴ using the software package DIRAC04. ^bCalculated by Fujii *et al.*³⁰ using UTChem program at 298 K. ^cCalculated for Tl(H₂O)₆³⁺. ^dCalculated for Tl(H₂O)₃⁺. ^eCalculated for Tl(H₂O)₆³⁺.

For example, according to Ghosh *et al.*³⁴, the NVE-driven $\Delta_{NV}^{199}\text{Hg}$ of Hg⁰ is about 0.14‰, therefore, $\Delta_{NV}^{199}\text{Hg}$ of Hg²⁺ should be -0.59‰ (i.e., -0.73‰ of Hg²⁺ listed in Table 4 plus 0.14‰).

All $\Delta_{NV}^{199}\text{Hg}$ and $\Delta_{NV}^{201}\text{Hg}$ values listed in Table 4 except for Hg⁰ are negative. For all studied Hg species, the MIF ratio of two odd-mass isotopes (i.e., $\Delta_{NV}^{199}\text{Hg}/\Delta_{NV}^{201}\text{Hg}$) will fall on a straight line with the slope of 1.66 (Fig. 4), suggesting they will be changed in a proportional way. This result is almost identical to a previous theoretical result³⁵ (i.e., with the slope of 1.65). This special relationship can be used to study MIFs caused by other reason via distinguishing the NVE signals from them.

Moreover, NVE can also cause mass-independent fractionations for odd-mass isotope (²⁰⁷Pb) and even-mass isotope (²⁰⁴Pb) (Table 5). The largest signals of NVE-driven MIF are up to -0.39‰ ($\Delta_{NV}^{207}\text{Pb}$) and 0.41‰ ($\Delta_{NV}^{204}\text{Pb}$) among all the studied species relative to Pb⁰ at 25 °C. The signs of even-mass isotope MIF ($\Delta_{NV}^{204}\text{Pb}$) and odd-mass isotope MIF ($\Delta_{NV}^{207}\text{Pb}$) are opposite to each other although their magnitudes are almost the same (Table 5).

The calculation details, including optimized geometries, energies, harmonic vibrational frequencies, *et al.*, have been documented in the Supplementary file for interested reader.

Discussion

One of the special features of NVE is that it can cause large isotope fractionations between isolated atoms and ions (e.g., Hg²⁺-Hg⁰, Tl³⁺-Tl⁺ and Pb⁴⁺-Pb²⁺), which there would be no fractionation at all if based on the classical mass-dependent isotope fractionation theory, because there is no difference in terms of kinetic energies for them. Moreover, it seems that ions with more extra charges (e.g., with fewer s orbital electrons) can have larger NVEs and isotope fractionation potential than those with lesser charges (e.g., Tl³⁺ vs. Tl⁺, Pb⁴⁺ vs. Pb²⁺).

Comparing with previous studies (Fig. 1b and Table 1), our NVE-driven Hg isotope fractionation results are noticeably different from those of Schauble⁴ and Wiederhold *et al.*³⁵. The NVE is proportional to difference in mean square nuclear charge radius of different nuclei (i.e., $NVE \propto \delta \langle r^2 \rangle$ and $\delta \langle r^2 \rangle = \langle r^2 \rangle_A - \langle r^2 \rangle_{A'}$), as King¹⁰ has pointed out based on spectrometric results. Therefore, we can explain the difference between Schauble⁴ and our results very clearly. Schauble⁴ used the nuclear charge radii of Angeli⁴² (i.e., $\langle r^2 \rangle^{1/2}$ of ²⁰²Hg and ¹⁹⁸Hg are 5.4633fm and 5.4466fm) and the nuclear charge radius difference ($\delta \langle r^2 \rangle = \langle r^2 \rangle_A - \langle r^2 \rangle_{A'}$) is 0.182fm². But we use the nuclear charge radii from

species	0 °C	25 °C	100 °C	300 °C	1000 °C	0 °C	25 °C	100 °C	300 °C	1000 °C
	CMDE fractionation factor ($1000 \cdot \ln \beta_{207-206}^{MD}$)					CMDE fractionation factor ($1000 \cdot \ln \beta_{208-206}^{MD}$)				
Pb ⁰	0	0	0	0	0	0	0	0	0	0
Pb ²⁺	0	0	0	0	0	0	0	0	0	0
Pb ⁴⁺	0	0	0	0	0	0	0	0	0	0
PbCl ₄ ²⁻	0.12	0.10	0.06	0.03	0.01	0.24	0.20	0.13	0.05	0.01
PbBr ₄ ²⁻	0.10	0.09	0.06	0.02	0.00	0.21	0.17	0.11	0.05	0.01
PbCl ₄	0.43	0.36	0.24	0.10	0.02	0.85	0.72	0.47	0.20	0.04
	NVE fractionation factor ($1000 \cdot \ln \beta_{207-206}^{NV}$)					NVE fractionation factor ($1000 \cdot \ln \beta_{208-206}^{NV}$)				
Pb ⁰	0	0	0	0	0	0	0	0	0	0
Pb ²⁺	-0.26	-0.24	-0.19	-0.13	-0.06	-0.66	-0.60	-0.48	-0.31	-0.14
Pb ⁴⁺	1.61	1.48	1.18	0.77	0.35	4.06	3.72	2.97	1.93	0.87
PbCl ₄ ²⁻	-0.27	-0.25	-0.20	-0.13	-0.06	-0.68	-0.62	-0.50	-0.32	-0.15
PbBr ₄ ²⁻	-0.26	-0.24	-0.19	-0.12	-0.06	-0.65	-0.59	-0.47	-0.31	-0.14
PbCl ₄	0.62	0.56	0.45	0.29	0.13	1.55	1.42	1.13	0.74	0.33
	Total fractionation factor ($1000 \cdot \ln \beta_{207-206}$)					Total fractionation factor ($1000 \cdot \ln \beta_{208-206}$)				
Pb ⁰	0	0	0	0	0	0	0	0	0	0
Pb ²⁺	-0.26	-0.24	-0.19	-0.13	-0.06	-0.66	-0.60	-0.48	-0.31	-0.14
Pb ⁴⁺	1.61	1.48	1.18	0.77	0.35	4.06	3.72	2.97	1.93	0.87
PbCl ₄ ²⁻	-0.15	-0.15	-0.14	-0.10	-0.05	-0.44	-0.42	-0.37	-0.27	-0.14
PbBr ₄ ²⁻	-0.16	-0.15	-0.13	-0.10	-0.06	-0.44	-0.42	-0.36	-0.26	-0.13
PbCl ₄	1.05	0.92	0.69	0.39	0.15	2.40	2.14	1.60	0.94	0.37

Table 3. Calculated stable isotope fractionation factors for Pb-bearing species relative to Pb⁰ (in per mil), including conventional mass-dependent effect (CMDE) and nuclear volume effect (NVE).

	¹⁹⁸ Hg/ ¹⁹⁸ Hg	¹⁹⁹ Hg/ ¹⁹⁸ Hg	²⁰⁰ Hg/ ¹⁹⁸ Hg	²⁰¹ Hg/ ¹⁹⁸ Hg	²⁰² Hg/ ¹⁹⁸ Hg
NV scaling factor	0.0	0.0525	0.4732	0.6312	1.0
CMD scaling factor	0.0	0.2539	0.5049	0.7539	1.0
	$\Delta_{NV}^{198}Hg$	$\Delta_{NV}^{199}Hg$	$\Delta_{NV}^{200}Hg$	$\Delta_{NV}^{201}Hg$	$\Delta_{NV}^{202}Hg$
Hg ⁰	0.0	0.0	0.0	0.0	0.0
Hg ²⁺	0.0	-0.73	-0.11	-0.44	0.0
HgCl ₂	0.0	-0.29	-0.05	-0.18	0.0
HgBr ₂	0.0	-0.28	-0.04	-0.17	0.0
Hg(CH ₃)Cl	0.0	-0.18	-0.03	-0.11	0.0
Hg(CH ₃) ₂	0.0	-0.13	-0.02	-0.08	0.0
HgCl ₃ ⁻	0.0	-0.43	-0.07	-0.26	0.0
HgCl ₄ ²⁻	0.0	-0.51	-0.08	-0.31	0.0
HgBr ₃ ⁻	0.0	-0.40	-0.06	-0.25	0.0
HgBr ₄ ²⁻	0.0	-0.48	-0.08	-0.29	0.0
Hg(H ₂ O) ₆ ²⁺	0.0	-0.64	-0.10	-0.39	0.0
Hg(OH) ₂	0.0	-0.23	-0.04	-0.14	0.0

Table 4. Conventional mass-dependent (CMD) and nuclear volume (NV) scaling factors and Δ_{NV}^A Hg values for different Hg-bearing species relative to Hg⁰ at 25 °C.

Fricke and Heilig⁴¹ (i.e., $\langle r^2 \rangle^{1/2}$ of ²⁰²Hg and ¹⁹⁸Hg are 5.462fm and 5.443fm) and the nuclear charge radius difference ($\delta \langle r^2 \rangle = \langle r^2 \rangle_A - \langle r^2 \rangle_{A'}$) is 0.207fm². Our results are roughly 1.137 times of those of Schauble⁴, consistent with the radii difference ratio, i.e., 0.207/0.182 = 1.137.

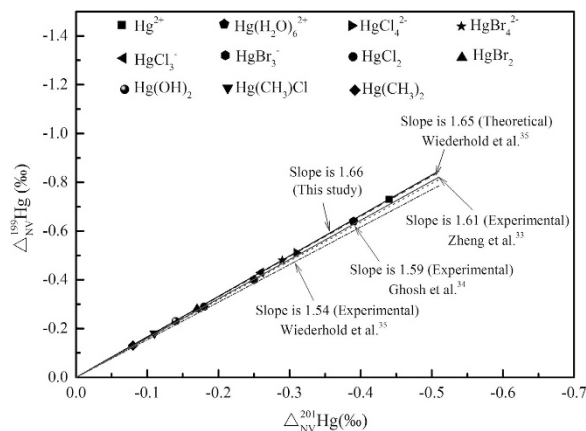


Figure 4. $\Delta_{\text{NV}}^{199}\text{Hg}$ versus $\Delta_{\text{NV}}^{201}\text{Hg}$ from NVE ($\Delta_{\text{NV}}^{\text{A}}\text{Hg}$ -factors relative to Hg^0 vapor) for Hg-bearing species at 25 °C. The black line is the slope of $\Delta_{\text{NV}}^{199}\text{Hg}/\Delta_{\text{NV}}^{201}\text{Hg}$ of this study. The black dash, dark gray solid, dark gray dot, dark gray short dash dot lines are the slope of $\Delta_{\text{NV}}^{199}\text{Hg}/\Delta_{\text{NV}}^{201}\text{Hg}$ based on theoretical NVE-driven MIF calculated with radii of Landolt-Boernstein Databas⁴¹ (calculated by Wiederhold *et al.*³⁵), experimental NVE-driven MIF of Zheng *et al.*³³, Ghosh *et al.*³⁴, Wiederhold *et al.*³⁵, respectively.

	²⁰⁴ Pb/ ²⁰⁶ Pb	²⁰⁶ Pb/ ²⁰⁶ Pb	²⁰⁷ Pb/ ²⁰⁶ Pb	²⁰⁸ Pb/ ²⁰⁶ Pb
NV scaling factor	-0.9097	0.0	0.3980	1.0
CMD scaling factor	-1.0193	0.0	0.5026	1.0
	$\Delta_{\text{NV}}^{204}\text{Pb}$	$\Delta_{\text{NV}}^{206}\text{Pb}$	$\Delta_{\text{NV}}^{207}\text{Pb}$	$\Delta_{\text{NV}}^{208}\text{Pb}$
Pb ⁰	0.0	0.0	0.0	0.0
Pb ²⁺	-0.07	0.0	0.06	0.0
Pb ⁴⁺	0.41	0.0	-0.39	0.0
PbCl ₄ ²⁻	-0.07	0.0	0.06	0.0
PbBr ₄ ²⁻	-0.06	0.0	0.06	0.0
PbCl ₄	0.16	0.0	-0.15	0.0

Table 5. Conventional mass-dependent (CMD) and nuclear volume (NV) scaling factors and $\Delta_{\text{NV}}^{\text{A}}\text{Pb}$ values for different Pb-bearing species relative to Pb⁰ at 25 °C.

We use the same mean square nucleare radii as Wiederhold *et al.*³⁵, but different methods (i.e., DHF vs. MP2, respectively), which lead to different results. In addition, there are suggestions that calculated results of Hg-bearing species used the mean square nuclear charge radii of Fricke and Heilig⁴¹ are closer to the experiment results³⁵. Note that different versions of the calculation software package (i.e., DIRAC04 and DIRAC13.1) have little impact on the calculated results (see Table S6).

Our NVE-driven Tl isotope fractionation results are in comparison with those of Schauble⁴ and Fujii *et al.*³⁰ in terms of $1000 \cdot \ln \beta^{\text{NV}}$ (Fig. 2b and Table 2). Our results agree with those of Schauble⁴ perfectly because of using the similar methods and the same mean square nuclear charge radii (i.e., those radii from Angeli⁴²). The only one exception is for the Tl(H₂O)₃⁺ case. The fractionations between Tl(H₂O)₃⁺ and Tl⁰ are larger than those of Schabule⁴, i.e., our result is -0.19‰ and their result is -0.11‰, at 25 °C. Our results are indeed very close to those of Fujii *et al.*³⁰ with small differences (ca. 0.04–0.07‰).

Previous researches have shown mercury can undergo mass-dependent fractionation (MDF) as well as mass-independent fractionation (MIF) for odd-mass isotopes ($\Delta^{199}\text{Hg}$ or $\Delta^{201}\text{Hg}$)^{43–46} and even-mass isotope ($\Delta^{200}\text{Hg}$)^{39,40}. The mechanism leading to the even-number Hg isotope mass-independent fractionation is still unclear.

Gratz *et al.*³⁹ firstly reported $\Delta^{200}\text{Hg}$ in Great Lakes precipitation and ambient air up to 0.25‰. Later, Chen *et al.*⁴⁰ found larger $\Delta^{200}\text{Hg}$ in precipitation from Peterborough where is located in subarctic zone. They showed that snow samples obtained in winter have surprisingly large $\Delta^{200}\text{Hg}$ values (up to 1.24‰) and rain water obtained in other seasons has much smaller $\Delta^{200}\text{Hg}$ values (about 0.21‰~0.42‰). However, there is no convincing evidence can explain the even-mass number Hg MIF enigma.

With the calculation data from this study, we find that the NVE-driven $\Delta_{\text{NV}}^{200}\text{Hg}$ cannot be the reason to explain those even-mass number Hg MIF results. First, the magnitudes of NVE-driven $\Delta_{\text{NV}}^{200}\text{Hg}$ are much smaller than those found by Gratz *et al.*³⁹ and Chen *et al.*⁴⁰. Second, the sign of NVE-driven even-mass number Hg MIFs calculated here is opposite to those reported $\Delta^{200}\text{Hg}$ results, meaning the NVE causes depletion of ^{200}Hg instead of enrichment of ^{200}Hg relative to ^{198}Hg . Therefore, the observed large positive $\Delta^{200}\text{Hg}$ signals must have other reasons or processes to be produced.

Because the half-life times of uranium isotopes are all very long, e.g., 4.5Ga for ^{238}U and 0.7Ga for ^{235}U , people actually treat uranium isotope system as a stable one in many young geologic systems^{47,48}. As the decayed products of uranium, Pb isotope system can also be treated as a regular stable isotope system for young and closed geologic systems with homogenized formation processes. For example, in some rocks formed less than 10 million years (or younger), or in some plants, or in any system which is young and homogenized before its formation. The equilibrium Pb isotope fractionations between two compounds in such systems can be meaningful and useful. In such systems, the radiogenic Pb isotope differences are no longer existing but homogenized to a background value. For instance, a system with inherited very high $^{208/206}\text{Pb}$ value has been homogenized in some processes. The compounds in such system will all have very high $^{208/206}\text{Pb}$ values. Meanwhile, the small differences of $^{208/206}\text{Pb}$ values among different compounds are caused by mass-driven and NVE-driven isotope fractionations. Our results can be used to explain such differences.

Fujii *et al.*³¹ firstly reported calculated NVE-driven Pb isotope fractionation factors for Pb^0 and Pb^{2+} -bearing species. We provide results of several new Pb-bearing species especially for Pb^{4+} -bearing species. If comparing the NVE results between Pb^0 and Pb^{2+} of Fujii *et al.*³¹ and ours, our results are marginally larger than theirs ($^{208}\text{Pb}/^{206}\text{Pb}$: 0.60‰ vs. 0.393‰ and $^{207}\text{Pb}/^{206}\text{Pb}$: 0.24‰ vs. 0.156‰) due to different methods and software packages used. In general, Pb isotope fractionations among Pb^{2+} -species are very small even with the driving force of NVE. However, we find surprisingly large fractionations (ca. 2 to 4‰) between Pb^{4+} -bearing species and Pb^{2+} -bearing species at room temperature. The fractionation magnitudes are similar or even larger than those Fe isotope fractionations between ferric and ferrous Fe-bearing species (e.g., $\text{Fe}^{3+}_{(\text{aq})}$ vs. $\text{Fe}^{2+}_{(\text{aq})}$) at low temperature, which have been broadly used as tracer for the change of redox conditions. Therefore, Pb isotope fractionations probably can also be used as a new tracer to study redox condition changes in young and closed geologic systems.

The occurrence of such large isotope fractionations is because the β -values of Pb^{4+} -bearing and Pb^{2+} -bearing species are in different directions, as the consequence of unique nuclear volume effects. Pb^{4+} -bearing species enrich heavy isotopes relative to Pb^0 . However, Pb^{2+} -bearing species enrich light isotope compared to Pb^0 , meaning β -values of Pb^{2+} -bearing species are smaller than the unity. It is similar to the case of Tl^+ -bearing species. Pb^0 is the one has more p electrons than Pb^{2+} -bearing species. This finding cannot be explained if only based on classical isotope fractionation theory, which suggests all β values of any kind of isotope systems must be equal or larger than the unity.

Conclusions

In this study, quantum chemical calculations (Dirac-Hartree-Fock) confirm that the nuclear volume effect plays a dominant role in equilibrium isotope fractionation for mercury, thallium and lead systems compared to the contributions of conventional mass-dependent effect, and agree with those conclusions of previous studies^{4,30,31}. NVE-driven $^{202}\text{Hg}/^{198}\text{Hg}$, $^{205}\text{Tl}/^{203}\text{Tl}$, $^{207}\text{Pb}/^{206}\text{Pb}$ and $^{208}\text{Pb}/^{206}\text{Pb}$ fractionations for Hg-, Tl- and Pb-bearing species can be up to 3.61‰, 2.54‰, 1.48‰ and 3.72‰ at 25 °C, respectively. Moreover, the NVE-driven mass-independent fractionations of ^{199}Hg are larger than those of ^{201}Hg and ^{200}Hg which is up to -0.73% . The ratio of $\Delta_{\text{NV}}^{199}\text{Hg}/\Delta_{\text{NV}}^{201}\text{Hg}$ is 1.66 which agrees well with previous experimental and theoretical results. Furthermore, those NVE-driven MIFs of ^{200}Hg calculated here are not only too small to be compared with the $\Delta^{200}\text{Hg}$ results reported in snow and water samples^{39,40}, but also with the opposite sign, meaning the NVE is not the reason of those $\Delta^{200}\text{Hg}$ signals.

Surprisingly, we find Pb isotope fractionations between Pb^{4+} -bearing and Pb^{2+} -bearing species can be up to 2 - 4‰ at room temperatures, suggesting a potential new tracer for redox condition changes in young and closed geologic systems. The NVE-driven MIFs of $\Delta_{\text{NV}}^{207}\text{Pb}$ and $\Delta_{\text{NV}}^{204}\text{Pb}$ are with moderate magnitudes but in opposite signs (i.e., $\Delta_{\text{NV}}^{207}\text{Pb} \approx -\Delta_{\text{NV}}^{204}\text{Pb}$).

Methods

Conventional mass-dependent effect (CMDE). Bigeleisen and Mayer¹¹ and Urey¹² suggested a well-known method for calculating the isotope fractionation factor, which is called the Bigeleisen-Mayer equation (hereafter B-M equation) or the Urey model. The B-M equation was based on the Born-Oppenheimer and harmonic approximations. According to the B-M equation, the natural logarithm of the isotope fractionation factor for an isotope exchange reaction under high-temperature approximations is

$$\varepsilon \approx \ln \alpha_0 \propto \Delta m / (mm'T^2) \quad (1)$$

where ε is the isotope enrichment factor and is roughly equal to $\ln \alpha_0$; α_0 is the isotope fractionation factor; m and m' are the masses of the heavy and light isotopes, respectively; Δm is the relative mass

difference of isotopes (i.e., $\Delta m = m - m'$). When the temperature is constant, enrichment factor is proportional to $\Delta m / mm'$. According to this equation, the isotope fractionation of heavy elements (e.g., Hg, Tl or Pb) would be small.

For an exchange reaction $A'Y + AX = A'X + AY$, the equilibrium CMDE fractionation factors is calculated^{11,12}

$$\alpha^{\text{CMDE}} = \text{RPFR} \left(\frac{AY}{A'Y} \right) / \text{RPFR} \left(\frac{AX}{A'X} \right) \quad (2)$$

where RPFR is the reduced partition function ratio and it is expressed in term of the harmonic vibrational frequencies with isotope substitution

$$\text{RPFR} = \prod_i^{3n-6} \frac{u_i(A'X) \exp(-u_i(A'X)/2)}{u_i(AX) \exp(-u_i(AX)/2)} \frac{1 - e^{-u_i(AX)}}{1 - e^{-u_i(A'X)}} \quad (3)$$

where A and A' are the heavy and light isotopes of the element A; $u_i(AX) = h\nu_i(AX)/kT$; $\nu_i(AX)$ is the *i*th harmonic vibrational frequency of AX molecule; h and k are Planck and Boltzmann constant; T is the absolute temperature.

Nuclear volume effect (NVE). Based on spectrometric results, King¹⁰ proposed that the NVE was proportional to difference in mean square nuclear charge radius of different nuclei (i.e., $\text{NVE} \propto \delta \langle r^2 \rangle$ and $\delta \langle r^2 \rangle = \langle r^2 \rangle_A - \langle r^2 \rangle_{A'}$). Upon the inspiration of U isotope exchange experiments, Bigeleisen² revised the B-M equation and added the NVE term into it. The logarithm of the corrected isotope fractionation factor became

$$\varepsilon \approx \ln \alpha = \ln \alpha_0 + \ln K_{\text{anh}} + \ln K_{\text{BOELE}} + \ln K_{\text{fs}} + \ln K_{\text{hf}} \quad (4)$$

where $\ln \alpha_0$ is the isotope fractionation factor under the B-M equation approximations; $\ln K_{\text{anh}}$ is the anharmonic correction term; $\ln K_{\text{BOELE}}$ is the correction to the Born-Oppenheimer approximation; $\ln K_{\text{fs}}$ is the NVE term (also called nuclear field shift); $\ln K_{\text{hf}}$ is the term for nuclear spin effect. In the terminology of Bigeleisen, the nuclear field shift actually includes both shape and size effects. However, the contribution from nuclear size is easy to calculate but that from nuclear shape is very difficult to evaluate and trivial. Therefore, people trend to use NVE instead of nuclear field shift for more precise description⁴.

Because of extremely small anharmonic corrections for heavy elements, $\ln K_{\text{anh}}$ can be safely neglected². The correction to the Born-Oppenheimer approximation is related to $\Delta m / mm'$ ^{49,50}. Therefore, $\ln \alpha_0$ and $\ln K_{\text{BOELE}}$ are both proportional to $\Delta m / mm'$ when temperature is a constant. Based on the investigations on U isotope exchange reactions, Bigeleisen² showed that nuclear spin effect was also very small and could be safely neglected.

Because the NVE is related to the difference in ground-state electronic energies, it can be written as²

$$\ln K_{\text{fs}} = (kT)^{-1} \{ [E^0(AX) - E^0(A'X)] - E^0(AY) - E^0(A'Y) \} \quad (5)$$

where E^0 is the ground-state electronic energy; AX and A'X represent different isotopologues; k is the Boltzmann's constant and T is in absolute temperature (K). We can see the magnitude of NVE is proportional to $1/T$ and to ground-state electronic energy differences due to isotopic substitutions.

Mass-independent fractionation (MIF). Here we use Hg isotopes as an example to introduce the concept of mass-independent isotope fractionation (MIF). If we define $\delta^A\text{Hg}$ as

$$\delta^A\text{Hg} = [({}^A\text{Hg}/{}^{198}\text{Hg})_{\text{sample}} / ({}^A\text{Hg}/{}^{198}\text{Hg})_{\text{ref}} - 1] \times 1000 \quad (6)$$

Then the mass-independent isotope fractionation (MIF) of any pair of Hg isotopes (e.g., ${}^A\text{Hg}/{}^{198}\text{Hg}$) will be

$$\begin{aligned} \Delta^A\text{Hg} &= \delta'^A\text{Hg} - \lambda_{\text{MD}} \delta'^{202}\text{Hg} \\ &= \lambda_{\text{TOTAL}} \delta'^{202}\text{Hg} - \lambda_{\text{MD}} \delta'^{202}\text{Hg} \end{aligned} \quad (7)$$

where $\delta'^{202}\text{Hg} = 1000 \cdot \ln \left(1 + \frac{\delta^{202}\text{Hg}}{1000} \right)$ ^{51,52}, λ_{TOTAL} includes λ_{MD} (the conventional mass-dependent scaling factor), λ_{NV} (the nuclear volume scaling factor), λ_{MIE} (the magnetic isotope effect scaling factor) and other scaling factors. And if we just consider the MIF caused by the NVE, it would be

$$\Delta_{\text{NV}}^A\text{Hg} = (\lambda_{\text{NV}} - \lambda_{\text{MD}}) \delta'^{202}\text{Hg} \quad (8)$$

where λ_{MD} is actually calculated using the high temperature approximation of equilibrium fractionation⁵²

$$\lambda_{\text{MD}} = \frac{\left(\frac{1}{m_i} - \frac{1}{m_j}\right)}{\left(\frac{1}{m_i} - \frac{1}{m_k}\right)} \quad (9)$$

This is because λ_{MD} values for heavy metal isotope systems are only weakly temperature-dependent⁵³. λ_{NV} is calculated from the mean square nuclear charge radii⁴

$$\lambda_{\text{NV}} = \frac{\langle r_i^2 \rangle - \langle r_j^2 \rangle}{\langle r_i^2 \rangle - \langle r_k^2 \rangle} \quad (10)$$

where m_i , m_j and m_k are the masses of isotopes i , j and k , respectively; $\langle r_i^2 \rangle$, $\langle r_j^2 \rangle$ and $\langle r_k^2 \rangle$ are their mean square nuclear charge radii.

Unfortunately, $\Delta_{\text{NV}}^{\text{A}}\text{Hg}$ cannot be calculated theoretically because the value of $\delta^{202}\text{Hg}$ for a specific Hg species is unknown. Instead, we calculate the relative MIF in comparison of Hg vapor (Hg^0):

$$\Delta_{\text{NV}}^{\text{A}}\text{Hg}_{\text{HgX}} - \Delta_{\text{NV}}^{\text{A}}\text{Hg}_{\text{Hg}^0} = (\lambda_{\text{NV}}^{\text{A}} - \lambda_{\text{MD}}^{\text{A}}) \cdot 1000 \cdot \ln \alpha_{\text{HgX-Hg}^0}^{\text{NV}} \quad (11)$$

If experimental results of MIF for Hg^0 vapor caused by NVE are available, we can obtain MIFs of other Hg species through “equation (11)”. For example, Ghosh *et al.*³⁴ observed MIFs for odd isotopes (^{199}Hg and ^{201}Hg) and small MIFs for even isotope (^{200}Hg) in the vapor phase (Hg^0) caused by NVE at room temperature. Their average $\Delta_{\text{NV}}^{199}\text{Hg}$, $\Delta_{\text{NV}}^{201}\text{Hg}$ and $\Delta_{\text{NV}}^{200}\text{Hg}$ values for Hg^0 were $0.14 \pm 0.01\%$, $0.09 \pm 0.01\%$ and $0.01 \pm 0.03\%$, respectively. The results of equilibrium evaporation experiments of Estrade *et al.*⁴⁵ were similar to those of Ghosh *et al.*³⁴ and their $\Delta_{\text{NV}}^{199}\text{Hg}$, $\Delta_{\text{NV}}^{201}\text{Hg}$ and $\Delta_{\text{NV}}^{200}\text{Hg}$ values for Hg^0 were 0.12% , 0.07% and 0.01% , respectively, in the temperature range of 2–22 °C.

Computational quantum chemistry methods. Ground-state electronic energies calculations are performed with DIRAC13.1 software package⁵⁴. All-electron Dirac–Hartree–Fock (DHF) theory is used to calculate relativistic electronic structures of Hg-, Tl- and Pb-bearing species with four-component wave functions. Our calculation details are similar to those of Schauble⁴. “Double-zeta” basis sets^{55,56} are used for Hg, Tl and Pb atoms and uncontracted cc-pVDZ basis sets⁵⁷ are used for other light atoms (H, O, C, Cl and Br). The molecular geometries were firstly optimized at pseudo-potential HF calculations (by Gaussian 03 software⁵⁸) as initial guesses. Following the methods of Schauble⁴, we optimize structures by using the iteratively quadratic fitting method (i.e., energy vs. bond-length fitting) instead of free geometry optimization using Dirac 13.1 to save computing time. $\text{Hg}(\text{H}_2\text{O})_6^{2+}$ (T_h), HgCl_4^{2-} (T_d), $\text{Tl}(\text{H}_2\text{O})_6^{2+}$ (T_h), $\text{Tl}(\text{H}_2\text{O})_3^+$ (C_3) and PbCl_4^{2-} (T_d) are chosen to compare their results calculated by the iteratively quadratic fitting method and by the free optimization method. The results show that these two methods can produce almost identical geometries but the former consumes much lesser time.

After geometry optimization, all Hg-, Tl- and Pb-bearing species are calculated for obtaining their ground-state electronic energies by using DIRAC 13.1. Different isotopologues will use their own Gaussian exponent ξ as in this form⁵⁹:

$$\xi = \frac{3}{2 \langle r^2 \rangle} \quad (12)$$

where the mean square nuclear charge radii ($\langle r^2 \rangle$) can be found from the Landolt-Boernstein Database⁴¹ for Hg and from Angeli⁴² for Tl and Pb.

Different from closed shell species, we also use the complete open shell configuration interaction (COSCI) method to calculate the ground-state energies of opened shell species (Tl^0 and Pb^0 with the electron configuration as $[\text{Xe}]4f^{14}5d^{10}6s^26p^1$ and $[\text{Xe}]4f^{14}5d^{10}6s^26p^2$, respectively).

With the calculated ground-state electronic energies, the NVE can be calculated from “equation (5)”. For example, the NVE on isotope fractionation of an HgX-Hg^0 isotope exchange reaction is⁴

$$\begin{aligned} 1000 \cdot \ln \beta_{202-198}^{\text{NV}} &= 10^3 \{-\Delta E_{202-198}[\text{HgX-Hg}^0]\}/kT \\ &= 10^3 \{(E[^{202}\text{Hg}^0] - E[^{198}\text{Hg}^0]) - (E[^{202}\text{HgX}] - E[^{198}\text{HgX}])\}/kT \end{aligned} \quad (13)$$

where the $\beta(X)$ factor is the equilibrium fractionation factor between substance X and an ideal mono-atomic gas⁶⁰.

The molecular geometries and harmonic interatomic vibrational frequency are calculated at pseudo-potential Hartree-Fock (HF) level by Gaussian03 software package⁵⁸. We treat inner-shell electrons of Hg, Tl, and Pb atom by using relativistic pseudo-potentials. However, valance and intermediate-shell electrons are treated with a double-zeta basis sets (cc-pVDZ-PP) and cc-pVDZ basis sets are used for H, C, O, Cl and Br atoms.

The usual isotope fractionation between substance A and substance B is defined as

$$\Delta_{A-B} \approx 1000 \cdot \ln \alpha_{A-B} = 1000 \cdot \ln \beta_A - 1000 \times \ln \beta_B \quad (14)$$

where α is the equilibrium isotope fractionation factor.

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Author Contributions

Y.L. proposed and organized the project. S.Y. carried out the theoretical derivation and calculation work. S.Y. and Y.L. analyzed and interpreted the data together. S.Y. and Y.L. wrote the paper.

Additional Information

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