

Received: 14 March 2018 Accepted: 6 November 2018 Published online: 13 February 2019

OPEN In-situ behavioural and physiological responses of Antarctic microphytobenthos to ocean acidification

James G. Black¹, Jonathan S. Stark ², Glenn J. Johnstone², Andrew McMinn^{1,5}, Philip Boyd¹, John McKinlay², Simon Wootherspoon² & John W. Runcie^{3,4}

Ocean acidification (OA) is predicted to alter benthic marine community structure and function, however, there is a paucity of field experiments in benthic soft sediment communities and ecosystems. Benthic diatoms are important components of Antarctic coastal ecosystems, however very little is known of how they will respond to ocean acidification. Ocean acidification conditions were maintained by incremental computer controlled addition of high fCO₂ seawater representing OA conditions predicted for the year 2100. Respiration chambers and PAM fluorescence techniques were used to investigate acute behavioural, photosynthetic and net production responses of benthic microalgae communities to OA in in-situ field experiments. We demonstrate how OA can modify behavioural ecology, which changes photo-physiology and net production of benthic microalgae. Ocean acidification treatments significantly altered behavioural ecology, which in turn altered photophysiology. The ecological trends presented here have the potential to manifest into significant ecological change over longer time periods.

Ocean acidification (OA) is the change in seawater carbonate chemistry, including a reduction in pH, as a result of the absorption of atmospheric CO₂ into the oceans¹. Ocean acidification is predicted to cause a 0.4 decrease in ocean pH and a 190% increase in dissolved inorganic carbon (DIC) by 2100 under a scenario of "business as usual" CO₂ emissions, as modelled by the IPCC AR5 RCP 8.5². This change is expected to substantially alter the structure of most marine communities3-6.

The majority of OA research to date has been on single organisms in laboratory-based studies⁷. In general, these studies indicate that reduced pH will be detrimental to some organisms (some calcifying heterotrophs) but advantageous to others (non-calcifying algae), creating potential "winners" and "losers" in marine communities under future scenarios of higher atmospheric $CO_2^{3,4,8}$. Whether these single organism responses will translate to actual changes in natural communities, once all ecological interactions are included, is largely untested. If they do, it could result in altered ecosystem services, functioning and community structure⁹⁻¹¹. Therefore, community-scale OA experiments are an important gap in our current knowledge¹². Even though the effects of OA are predicted to affect higher latitude waters sooner than elsewhere, there has been little in-situ OA research on Antarctic benthic marine ecosystems¹³. The few community-scale OA studies that have utilised an in-situ field approach have been undertaken in tropical, deep sea or pelagic ecosystems^{7,14,15}, with no published polar research.

Microalgae and most autotrophic algae are generally expected to benefit under OA conditions in comparisons to heterotrophic organisms. This is likely the result of the increase in carbon sources (CO₂ and HCO₃-), used for photosynthesis, under OA conditions. However, different species of marine algae show different sensitivities to OA 16,17, with beneficial and detrimental effects of ocean acidification on photosynthesis reported for different algal species 17-21. Most studies attribute the different reponses to species specific differences in the efficiency of

¹Institute for Marine and Antarctic Studies, Castray Esplanade Hobart Tasmania, Hobart, Australia. ²Australia Antarctic Division, Antarctic Conservation and Management program, 203 Channel Hwy, Kingston, Tasmania, Australia. 3School of Life and Environmental Sciences, University of Sydney, 2006, Sydney, Australia. 4Aquation Pty Ltd, PO Box 3146, Umina Beach NSW, 2257, Sydney, Australia. ⁵College of Life Sciences, Ocean University of China, 5 YuShan Rd, Qingdao, China. Correspondence and requests for materials should be addressed to J.G.B. (email: jgblack@utas.edu.au)

carbon utilisation of CO_2 and HCO_3 —, via carbon concentrating mechanisms (CCM)^{17,22,23}. Different experimental protocols may also be a factor in some circumstances¹⁷. While these factors will likely contribute to responses to OA, attributing a generalised response of all algae to OA is confounded by the fact that marine algae (including microalgae) inhabit a diverse range of microenvironments.

Microalgae in the microphytobenthos (MPB) live in a microenvironment that is characterised by much steeper environmental gradients compared to their pelagic counterparts²⁴. Gradients of light, pH, O_2 and DIC change substantially in the MPB after the Diffusive Boundary Layer (DBL). The euphotic zone of MPB mats are typically much thinner than pelagic euphotic zones, yet the microbial activity is higher²⁵. This results in rapid and dynamic consumption and production of O_2 and CO_2 , influencing pH and DIC differentially down the sediment profile^{24–26}. These unique micro environmental factors could influence MPB responses to OA, yet there is a paucity of information on MPB OA responses²⁴. Additionally the ability of microalgae to deal with OA in the short term may be very different to longer term responses. A review of MPB responses to OA by Marques da Silva, *et al.*²⁴ stongly emphazies the need for a better understanding of MPB response to OA.

Microalgae play a vital role in ecosystem functioning ^{16,21}, especially in the MPB in Antarctic benthic ecosystems, which in low-light habitats under sea ice are the dominant autotroph in the absence of macro algae ^{27,28}. Microalgae form the basis of many marine food webs and perform key ecological roles in fixing carbon, recycling nutrients, stabilising sediments and contribute to modifying global climates via their role in the carbon cycle ^{28,29}. While the uptake of carbon by microalgae is essential for their photosynthesis, it is also intrinsically linked to ocean carbon cycling and will interact with OA ^{23,30–32}. Microalgae have a number of ecologically linked processes that may be altered by OA. For example, MPB photosynthetic yield, primary production and net production drive vital ecosystem services in soft sediment communities, which if altered will have flow on effects to the whole community³³. Changes to these ecosystem services will be positive or negative dependent largely on whether conditions are more or less preferential for microalgae³⁴.

Understanding the physiological preferences for microalgae in the context of carbonate chemistry is essential to understand the likely community shifts in response to future OA conditions⁶. Microalgal physiological preferences can be inferred from a range of quantitative measures such as the effective quantum yield of photosystem II light conversion efficiency (ϕ_{PSII}) and biomass changes^{21,35,36}. However, inferring short term (i.e. within hours) algal responses to OA would benefit from a behavioural indicator (i.e. avoidance/attraction behaviour). Active behaviour may seem unlikely for single celled marine algae, however many marine benthic diatoms have the capacity to migrate up and down in the sediment in response to external stimuli (i.e. sunlight, temperature, nutrients), to seek out their optimum or beneficial physiological conditions^{18,20,24,37}. Photo Active Radiation (PAR) is normally the primary driver of this vertical migration behaviour, which is referred to as the photo tactile response. This behaviour provides an additional measure of microalgal response to environmental conditions that is otherwise difficult to measure in most phytoplankton and macroalgae. Photo Tactile Response (PTR) in benthic diatoms has been used to measure responses to changed environmental conditions in other studies^{24,37}, however, it has not previously been used to examine responses to OA *in-situ*.

The purpose of this study was to identify eco-physiologically relevant trends in MPB responses to OA *in-situ*. The specific aims were to investigate if OA alters the already recognized PTR relationship to PAR; and to determine what effect changes in PTR may have on photosynthetic yield and net production. To test this we performed a series of *in-situ* mesocosm experiments in a sea-ice covered benthic habitat in Eastern Antarctica. Novel technology was used to provide insights into MPB community respiration, photosynthetic and behavioural responses to OA. This study provides the first measurements of *in-situ* effects of ocean acidification on MPB in a community context.

Results

Experimental treatments. A pH of 0.4 below ambient was maintained in incubation chambers for each deployment (Fig. 1). Other water parameters (i.e. aragonite (Ω) saturation, salinity, temperature) can be found in Table 1 or³⁸. Treatment water used in this experiment was pumped from the Antarctic Free Ocean Carbon Enrichment experiment (antFOCE) experimental chambers, acidified treatments maintained a pH offset of 0.3825 \pm 0.065, mean Ω_{ar} of 0.62 \pm 0.14 and mean fCO_2 values of 912.5 \pm 155 μ atm over the duration of their experiment. In contrast, water supplied to the control chambers maintained a pH of 8.061 \pm 0.044, mean Ω_{ar} of 1.39 \pm 0.11 and mean fCO_2 values of 354 \pm 42 μ atm. Treatment water pH, oxygen and salinity measurements were sampled from the antFOCE sensor system at the time water was pumped from antFOCE treatments into respiration chambers, (Fig. 1). The mean pH of acidified and control treatments, calculated over all deployments, was 7.687 \pm 0.01 and 8.082 \pm 0.018 respectively. xxx water parameters remained relatively stable between experiments (i.e salinity, temperature) (see Table 1 or³⁸).

Rapid light curves. *In-situ* rapid light curves indicated that the electron transport rate of MPB approached maximium relative electron transport rate (relETR_{max}) after $25\,\mu$ mol photons m⁻² s⁻¹ (Fig. 2).

Diatom species composition. Microscope counts of diatom cells indicated 15 dominant species (i.e. > 1% abundance) (see Table 2). *Trachyneis aspera, Cocconeis fasciolata, Pleurosigma elongatum* and *Cymbella* species were the most common species at the study site with 21, 17,14 and 9% relative abundance respectively. There were a total of 104 species found at this site within the study peroid. There were also low abundances in the MPB community of species commonly refered to as being associated with sea-ice (i.e. *Fragilariopsis cylindrus*).

Photo tactile response (PTR). Photo tactile response measured simultaneously in both treatment showed a significant relationship to the diel rate of change in PAR (i.e. first derivative of PTR changed with the first derivative of PAR) ($p = \langle 2e-16 \rangle$ (Fig. 3). Stated more simply, as light increased MPB biofilms migrated towards

	DIC (µmol kg ⁻¹)	TA (μmol kg ⁻¹)	$\Omega_{ m ar}$	fCO ₂ (µatm)
Control treatment				
Mean range (±s.d.)	2154 ± 35.72	2271 ± 28.48	1.39 ± 0.11	354±42
Acidified treatment				
Mean range (±s.d.)	2243.09 ± 54.66	2271 ± 28.15	0.62 ± 0.14	912.5 ± 154.5

Table 1. Averaged ranges of carbonate chemistry parameters of treatment water used in deployments A-D. DIC = calculated dissolved inorganic carbon; Ω_{ar} = saturation state of aragonite; fCO_2 = fugacity of CO_2 .(data reproduced from³⁸).

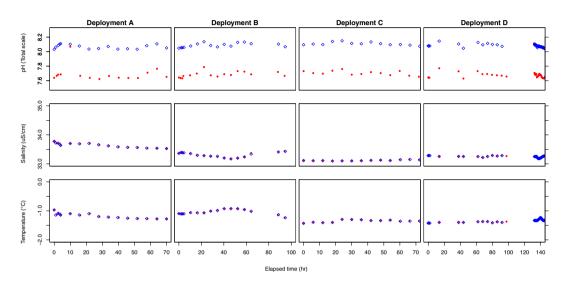


Figure 1. The pH (Total scale), salinity and temperature of control (blue diamonds) and acidified (lower red circles) in deployments (**A–D**) (from left to right respectively).

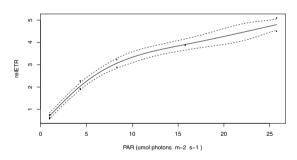


Figure 2. *In-situ* relETR of microphytobenthos under control conditions. A GAM model has been applied to estimate an average response (solid line) with 95% confidence intervals (broken lines).

the surface, as light decreased MPB migrated away from the surface. However, acidified PTR displayed a much stronger relationship to PAR $^{\rm d/dt}$ (p=2e-16, F=447.8), compared to the relationship in the control treatments (p=<2e-16, F=80) (see Fig. 3). The relationship between diel PAR and diel migration decreases in the control treatment at higher diel PAR rates of change (i.e. higher PAR values). By comparison the acidified diel migration rate maintains a more significant relationship to diel PAR rate even at higher PAR $^{\rm d/dt}$. Deployment C had a substantially different and irregular light regime, which resulted in a noncyclic and non-significant PTR pattern not seen in other deployments. This resulted in the MPB biomass staying at the surface for both treatments, with PTR to PAR being similar between treatments. Additionally as deployment time elapsed beyond 80 hours PTR was also non-significant in both treatments. Therefore, the effect of OA was explored on two MPB behavioural states, where MPB were either displaying normal PTR-PAR migration behaviour or not.

Diel rate of change of yield ϕ **PSII.** The effective quantum yield showed a diel rate of change that related to the diel change in PAR (i.e. first derivative of PAR was significantly related to the first derivative of diel yield ϕ_{PSII}) ($p = \langle 2e\text{-}16 \rangle$). Stated more simply, as PAR increased, the rate of photosynthesis (yield ϕ_{PSII}) also

Microphytobethos species	Relative Abundance (%)		
Trachyneis aspera	21		
Cocconeis fasciolata	17		
Pleurosigma elongatum	14		
Cymbella sp	9.3		
Achnanthes brevipes	4		
Amphora libyca	3.5		
Eucampia antarctica	3.5		
Pinnularia quadratarea	3.1		
Synedropsis recta	2.8		
Pleurosigma obscurum	2.3		
Navicula directa	1.8		
Fragilariopsis sublinearis	1.5		
Actinocyclus curvatulus	1.4		
Fragilariopsis cylindrus	1		
Diploneis splendida	0.7		
Actinocyclus actinochilus	0.7		
Odentella weissflogii	0.5		
Pinnularia quadratarea (variant constricta)	0.5		
Cocconeis pinnata	0.4		
Fragilariopsis kerguelensis	0.3		
Licmophora belgicea	0.2		
Thalassionema gelida	0.1		
Fragilariopsis ritscheri	0.1		
Manguinea fusiformis	0.1		
Auricula compacta	0.1		
Fragilariopsis obliquecostata	0.1		
Dicyocha sp	0.1		
Biddulphia areolata	0.1		
Fragilariopsis rhombica	0.05		
Melosira adeliae	0.05		
Pseudogomphonema kamtschatica	0.05		

Table 2. Microphytobenthic species at study site.

increased. The acidified treatment diel yield rate did show a more significant relationship to diel PAR (p = <2e-16, F-value = 545), compared to the control treatment (p = <2e-16, F-value = 10).

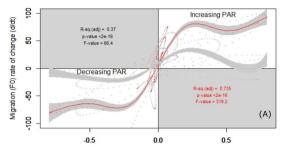
Yield ϕ **PSII maxima.** Yield (ϕ_{PSII}) increased as PAR increased until ~2 μ mol photons m⁻² s⁻¹ in both treatments when MPB were migrating normally. There was no substantial difference between treatments in yield maxima until PAR was higher than ~6 μ mol photons m⁻² s⁻¹, at which point control treatment yield then shows lower photosynthetic yield (Fig. 4). This is also observable at PAR maxima in Fig. 5(B and D). However, without PTR the acidified yield (ϕ_{PSII}) was trending lower than controls (Fig. 4), demonstrating that without PTR, there is a negative effect of OA on yield (ϕ_{PSII}). Under these conditions the PTR-PAR relationship had decreased and yield (ϕ_{PSII}) was found to be negatively affected in the acidified treatment compared to the control (Fig. 4).

Net production rate. Under low PAR conditions (i.e. $<10\,\mu$ mol photons m⁻² s⁻¹) a negative net production of oxygen was observed in all treatments/deployments. There did not appear to be a detectable pattern (from 20 cm above the sediment) of net production matching yield measurement under low PAR in deployments A, B and C (Fig. 4). However, under higher PAR conditions in deployment D (i.e. 8–10 μ mol photons m⁻² s⁻¹), a difference in the diel rate of net production was detected in the control, with an increase in the rate of net oxygen production observed in the acidified treatment (Fig. 4D).

Dark adapted yield (\phi PSII). There was no significant treatment effect on F_v/F_m . (see Table 3, Fig. 6).

Vertical sediment profile of Chl a. The vertical distribution of Chl a in the sediment decreased significantly below 8 mm (Fig. 7), with significant difference between treatments (p = 0.001, n = 4). However, these significant differences need to be considered in the context of starting heterogeneity of MPB biofilms, in addition to these samples being taken at the end of the deployment. The distribution of the Chl a at each depth relative the total Chl a in each sample core (Relative distribution of Chl a) is the more appropriate comparison. While there was no significant difference in relative Chl a profiles between treatments, Fig. 7 shows that Chl a in acidified

Generalised relationship between OA and migrating MPB



Generalised relationship between OA and non-migrating PTR

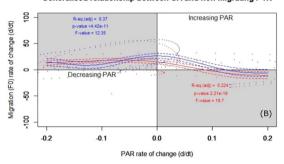


Figure 3. Rate of change of photo tactile response (PTR) in relation to the rate of change of PAR under normal migration behaviour (**A**) and non migrating PTR (**B**). The y-axis is the first derivative (F0^{d/dt}) of migration rate, with positive values indicating migration towards the surface, negative values migration away from the surface. The x-axis is the first derivative PAR (PAR^{d/dt}), with positive values when PAR increases, negative values for decreasing PAR. The colour of each test statistic (eg. *p*, F-value and Adjusted R-squared) represents it associated treatment. (**A**) Normal PTR in acidified treatment (Red line, with pink solid shaded 95% confidence intervals), normal PTR in control treatment (black line, with solid grey 95% confidence intervals). (**B**) Absent PTR in acidified treatment (red line, with red dashed 95% confidence intervals), absent PTR in control treatments (black line, with dashed black 95% confidence intervals).

Effect of OA on yield (PSII) in relation to PTR

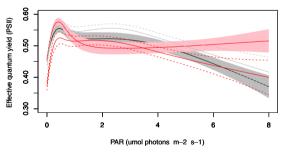


Figure 4. Effective quantum yield (ϕ_{PSII}) in relation PAR. Acidified treatment under normal PTR (Red line, with pink 95% confidence interval), control treatment under normal PTR (Black line, with grey 95% confidence interval). Non-PTR in acidified treatment (red line, with red dashed 95% confidence intervals), non-PTR in control treatment (grey line, with grey dashed 95% confidence intervals).

treatments have redistributed from deeper to shallower sediment depths (12-14 to 4-8 mm). There was no significant interaction between treatment and sediment depth for Chl a (i.e. profiles are similar).

Discussion

MPB microalgae demonstrated an ability to detect altered DIC states and actively changed their natural photo tactile behaviour to account for OA. Diel yield showed a strong relationship to diel PTR, with PTR and yield (ϕ_{PSII}) both having significant relationships with PAR. The changes in PTR under OA corresponded to increases in both diel yield rate and yield (ϕ_{PSII}) maxima. This was most pronounced under higher PAR. A number of external abiotic factors (such as irregular PAR cycles) can disrupt the normal PTR patterns. Under these circumstances, when the PTR-PAR relationship was not significant, yield (ϕ_{PSII}) was lower in the acidified treatments. This

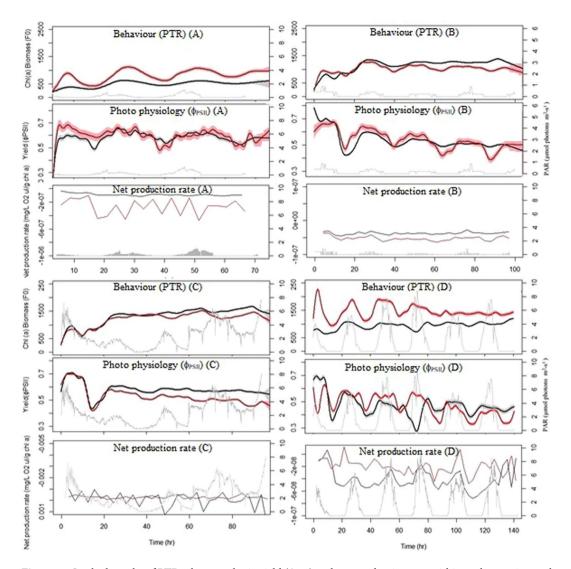


Figure 5. Stacked graphs of PTR, photosynthetic yield (ϕ_{PSII}) and net production rate within each experimental deployment (**A–D**). X-axis is elapsed time (hours) from start of deployment. Grey lines are Photosynthetic Active Radiation (PAR) with the scale on the right hand side of the graphs in μ mol photons m⁻² s⁻¹. Black lines (with grey 95% confidence intervals) are the control treatment and red lines (with pink 95% confidence intervals) are the acidified treatments.

Source	Df	Sum Sq	Mean Sq	F value	P value
In-vivo yield (F _v /F _m)					
Treatment	1	0.0236	0.023591	2.461	0.1227
Treatment:Deployment	6	0.1179	0.019644	2.049	0.0751
Residuals	53	0.5081	0.009586		
Chl a					
Treat	1	1831068	1831068	10.529	0.00143
Depth	9	14374551	1597172	9.184	3.7e-11
Treat:Depth	9	707018	78558	0.452	0.90459
Residuals	160	27826359	173915		

Table 3. Anova summaries for *in-vivo* yield (F_v/F_m) and sediment Chl a.

demonstrates the fundamental role of PTR in maintaining photo-physiology and how it interacts with OA. A minimal diel net production relationship (under high PAR only) was also observed.

Photo tactile response displayed a significant PAR driven diel pattern in both treatments ($p = \langle 2e-16 \rangle$) (Fig. 3), which has previously been reported in other ecosystems^{18,24,39,40}. However, the acidified treatment had a more significant relationship with PAR (with an F-value of 447.8 and 80 for acidified and control treatments respectively).

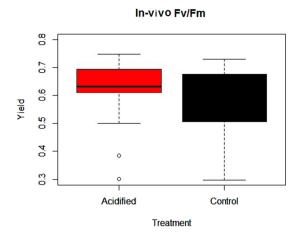


Figure 6. Fv/Fm (ϕ_{PSII}) of Control (black box) and Acidified (red box) treatments. Error bars+/- standard error.

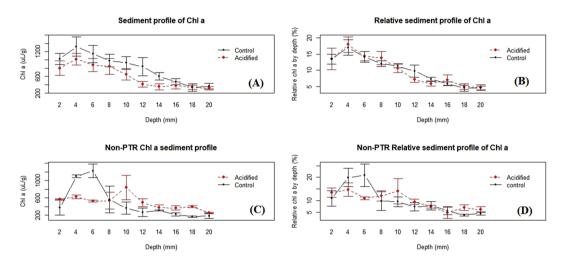


Figure 7. (A) Average chlorophyll a sediment profile of PTR deployments (**A,B** and **D**). **(B)** Relative Chl *a* concentration at each depth in relation to total core Chl *a* in PTR deployments. **(C)** Non-migrating PTR Chl *a* sediment profiles and **(D)** Non-migrating PTR relative Chl *a*. Control (black lines) and acidified (red lines) treatments in all plots. Error bars+/- standard error of subsamples.

Photo tactile responses in this study was not correlated with temperature or tide patterns, as found in other systems 40,41 . This is likely due to the very stable environmental conditions in Antarctica compared to other environments (i.e. temperature variation $0.1\,^{\circ}\text{C}^{-d}$, $<1\,\text{m}$ tidal change and a low flow rate $2-3\,\text{m}\,\text{s}^{-1}$). Microalgae were typically found down to $6-8\,\text{mm}$ into the sediment, with Chl a content decreasing substantially after 10 mm (Fig. 7).

Photosynthetic yield (ϕ_{PSII}) showed a significant diel yield relationship to PAR in both treatments (p = <2e-16 in both treatments). This was evidently intrinsically linked to PTR and demonstrated why behaviour of an MPB biofilm is a very important consideration for MPB primary producers under OA.

Yield (ϕ_{PSII}) of MPB is proportional to the amount of PAR, nutrients and carbon available for photosynthesis⁴². According to Hancke, *et al.*⁴³ the prior PAR regime (i.e. previous ~12 hours) will dictate the current microalgae PTR. Therefore MPB migration through the sediment enables moderation of exposure to PAR, nutrients and carbon in the overlying water (i.e. OA treatment)^{24,37} and will influence MPB primary production.

The majority of studies on microalgae have illustrated either positive or negligible effects of OA on their physiology 16 , with some reporting that many phytoplankton species are insensitive to OA 44,45 . However, our results clearly show that Antarctic MPB are very responsive to OA. The variable response among diatoms is reported to be due to the poor affinity of ribulose 1,5-bisphosphate carboxylase/oxygenase (RuBisCO) for CO₂ substrates used for photosynthesis, being only ~50% saturated under current CO₂ levels 16 . Therefore many diatoms operate carbon concentrating mechanisms (CCMs) which concentrate CO₂ around sites of photosynthesis. In future OA scenarios the DIC available for photosynthesis increases by ~190% (although only 1% of this is CO₂), and this is expected to saturate RuBisCO with CO₂ and reduce energy demands for concentrating CO₂ for photosynthesis. According to Cartaxana, *et al.* 18 , in a crowded MPB community, carbon may be a limiting resource even for organisms with highly efficient CCMs. Carbon limitation may explain why the results presented here indicate

an initial increased diel PTR and yield with elevated $f\text{CO}_2$ (i.e. OA conditions), followed by signs of acclimatisation. MPB biofilms can be up to ~1 mm thick, with an area of 1 m² representing a cell density equivalent to ~34,000 litres of high chlorophyll-a content open ocean water. Thus it is likely that carbon may be limited in these biofilms, and short term responses in PTR and yield (ϕ_{PSII}) are likely to be indicative of the relief of carbon limitation in these MPB biofilms.

The microalgal OA photo tactile response in the current study needs to be considered in the context of microalgae DIC/pH preferences and sediment pH gradients. Microalgae photosynthesis has the capacity to deplete DIC in the few millimetres either side of the DBL, with DIC depletion at the surface and increasing DIC with depth with strong vertical gradients²⁴. We would expect microalgae to seek out higher DIC/low pH regions when photosynthesising, yet avoid the energetic cost of pH homeostasis from low pH/high DIC when in other cellular cycles (i.e. Krebs or xanthophyll cycles).

A greater initial change in PTR and higher photosynthetic yields (ϕ_{PSII}) in the OA treatments occurred during periods of higher PTR^{d/dt} (Figs 3 and 5D). In this context it appears that initially, under higher PAR, microalgae are selecting environments with more available carbon and utilising this for increased photosynthetic rates. This corresponds with research by Cook and Røy⁴⁶ who found that increased rates of pore-water advection or addition of HCO₃ increased photosynthesis. Interestingly, experiments investigating the impact of increasing inorganic CO₂ on growth have produced a range of positive and negative results, which could be due to physiological variation between species or experimental protocols¹⁶. There are very few long term studies of OA on photosynthetic yield. However, increased diel yield was only apparent while the PTR-PAR relationship was maintained under acidified treatment conditions in the current study. After 80 hours, we see the acidified PTR decrease and the yield (ϕ_{PSII}) similar or lower than the control yield (ϕ_{PSII}) (Figs 4 and 5D). We suspect that MPB would be able to utilise extra DIC associated with OA for photosynthesis up until they have reached their maximum diel yield capacity. Any extra DIC not consumed by photosynthesis would then contribute to pH decreases (i.e. ocean acidification). Thus it is possible that negative diel yield after 80 hours is in part due to: (1) CO₂ saturated MPB biofilms not utilising extra DIC which then contributes to a lower pH in the diffusive boundary layer, which then requires them to increase proton pumps to maintain the intracellular pH homeostasis; or (2) an absence of PTR observed after 80 hours in the current study resulting in non-photochemical quenching (NPQ). If MPB are photo inhibited (NPQ) due to lack of PTR, then consequently they would not be using the extra DIC for photosynthesis and we would expect condition (1) above to influence yield. For example, deployment (C) had variable and non-cyclical PAR cycle events resulting in MPB staying at the sediment surface. This demonstrated that, when the diatoms remained on the surface with irregular PAR and maximised their exposure to OA and PAR, it resulted in negative photosynthetic yield (ϕ_{PSII}). It is possible that this is due to high non-photochemical quenching (NPQ) of the MPB as they receive more PAR than could be processed. Also, as stated above, if MPB are not actively using DIC for photosynthesis then it will contribute to decreasing pH in surrounding water bodies (i.e. in the Diffusive Boundary Layer). These findings are consistent with the findings of Hoppe, et al.⁴⁷ who illustrated that irregular PAR intensities strongly modulated the effects of OA on marine phytoplankton. They demonstrated that irregular/dynamic PAR reduced growth and strongly altered the effects of OA on primary production, being unaffected by elevated fCO_2 under constant PAR. Positive effects of OA on yield (ϕ_{PSII}) observed in the current study are assumed to be due to the ability to self-regulate exposure to OA and PAR through PTR. However, if this ability is impaired (i.e. no sediment migration) then we assume that MPB have to deal with lower pH without photosynthesis to buffer lower pH/high DIC (due to NPQ).

Previous studies on algae under OA scenarios have generally not found any differences in photosynthetic yield $^{48-50}$. However, the majority of studies investigating OA effects on yield (ϕ_{PSII}) have used Fv/Fm $^{49-52}$ rather than light-adapted samples or diel effective quantum yield (ϕ_{PSII}) measurements. There is a fundamental physiological difference in what yields signify under dark or light-adapted methods. Dark adapted yield (Fv/Fm) represents the maximum capacity of the photosystem to convert PAR energy into charged states and direct this into the photochemical pathway. This can be used to test the ability of photosystem (II) to function (i.e. if (ϕ_{PSII}) is damaged by a stimulus/toxin), with negative dark adapted yield (Fv/Fm) responses typically seen under toxin and nutrient stress 53 . In comparison, diel light-adapted yield is testing photosystem (II) rate of functioning at natural irradiances and is ideal for comparative OA field studies 53 . Therefore we would emphasize that in the context of OA and ecosystem responses, it is more appropriate to be examining diel light-adapted yield (ϕ_{PSII}) , especially in the field. In the current study we examined both light and dark-adapted yield (ϕ_{PSII}) . In contrast to what we observed with light-adapted-yield (ϕ_{PSII}) , when we examined dark adapted Fv/Fm (20 minutes) we observed no significant difference between treatments (Fig. 6).

Photosynthetic organisms have a high capacity to modify pH, with evidence that pH is regulated at the cell/water interface^{25,54}. Therefore we would not expect to see damage to photosystem (II) from more acidic conditions, provided they have the ability to buffer this at the cell/water interface using extra DIC for photosynthesis. So it is not surprising that there was no difference in dark adapted yield (ϕ_{PSII}) in this study or previous studies.

Photosynthetic yield (ϕ_{PSII}) correlates with oxygen production in most algae^{27,55}. In the current study, natural communities of heterotrophs and autotrophs were enclosed in the same respiration chambers in an ecological setting, therefore we can only report net community production rather than individual contributions of MPB to net production in relation to increased yield (ϕ_{PSII}). Furthermore, the DO sensors were 20 cm above the sediment in a 27 litre volume of water. It is expected that this reduced our detection limit substantially. Repeated 6 hour respiration incubations were always negative in net production. This indicates two potential situations: (1) a higher biomass of heterotrophs relative to microalgae; or, (2) the irradiance was not sufficient to reach the E_c (E^k), where oxygen production exceeds consumption. The present study observed a corresponding diel pattern in the net production rate and yield under higher PAR conditions in controls (Fig. 5D). This indicates that at higher PAR levels our DO detection limit was adequate. In this same deployment the acidified treatment had a higher net

production rate. Therefore we are confident that the higher net production rates observed in acidified treatments in deployment D are related to corresponding higher yield values.

Heterotrophic organism metabolic functioning (i.e. respiration) will reduce net production and needs to be considered in the context of our results. Metabolic up-regulation or depression in response to OA has been documented in many marine species, however some species show no change in metabolism 56,57 . Therefore, increases in net production under OA might be partially attributed to MPB increasing O_2 production, or a change in heterotrophic O_2 consumption.

Coastal pH often displays a strong short term diel rather than a static pH pattern. The short term responses in the current study are a valuable insight into MPB's capacity to deal with short term changes in carbonate chemistry conditions. The results from this study demonstrate that MPB communities are not insensitive to higher DIC and lower pH associated with OA. Any changes in net production and photosynthetic yield (ϕ_{PSII}) will have flow on effects via changes to dissolved oxygen levels and primary production as increased biomass at the base of the food web^{4,58}. With the global importance of MPB in supplying ecosystem services and their role in carbon cycling, any change in their physiology or ecology due to environmental changes such as ocean acidification could have potentially widespread consequences. Longer term ocean acidification induced changes in MPB physiology will be dependent on the following factors; (1) ability to migrate effectively through the sediment; and (2) ability to utilise extra DIC associated with OA for photosynthesis.

Materials and Methods

Study site. Field experiments were conducted at O'Brien Bay, near Casey Station, East Antarctica (66.311500° S, 110.514216° E) between 1st December 2014 and 1st March 2015. The site was characterised by 2.5 m thick multiyear sea ice and the depth ranged from 12 to 14 m. The dominant primary producers at the study site were microphytobenthos (MPB) and sea ice algae, with a complete absence of macroalgae due to the low PAR conditions at the site, which rarely exceeded $10\,\mu$ mol photons m^{-2} s⁻¹ at the seafloor (see³⁸). The MPB at the site formed a dense mat (~0.5–1 mm) on top of the sediment. The benthic marine communities at the site include a range of mobile macrofauna including burrowing anemones, asteroids, holothurians, and filter feeding invertebrates (sponges, ascidians, polychaetes). A previous study indicated significant spatial variation in infaunal communities at scales as small as 10 m, but such differences were small compared to those evident at larger spatial scales⁵⁹. The treatments in the current study were side by side (i.e. <75 cm apart), therefore we make the assumption that heterotrophic and autotrophic communities were relatively similar between treatments, within deployments. A general description of the diatom communities present at the site can be found in Polmear, *et al.*⁶⁰ or supporting species list (Table 2). Snow cover and thickness on the sea ice varied over the experimental period, resulting in different PAR exposures for each experiment. Water flow rate at the site was 2–3 cm s⁻¹. Temperature, alkalinity, pH, salinity, dissolved oxygen, were recorded continuously and were considered stable during the study period (see³⁸).

Experimental design. Experiments were conducted in two sealed respiration incubation chambers with a volume of 27 litres (Submersible Photosynthesis-Respiration System, Aquation Pty Ltd, Umina Beach, Australia), deployed on the seabed over patches of MPB. The chambers were made of UV-transparent acrylic with 25 cm diameter (chamber diameter) stainless steel sleeves on the base for insertion into the sediment. The chambers pumped acidified or control water from a larger long-term CO₂ enrichment experiment (Antarctic Free Ocean Carbon Enrichment antFOCE) via automated pumps (see³⁸). A five hour acclimation period was used to introduce water from the antFOCE system gradually, with an increasing ratio of treatment: ambient water pumped into chambers every 5 minutes.

Chamber deployment sites were within 15 m of the antFOCE experimental site. Treatments were placed side by side to reduce spatial differences in species abundance and irradiance. All experiments were conducted at a depth of 12-14 m. Acidified treatments maintained a pH reduction of 0.4 from ambient, in line with IPCC predictions under business as usual emissions scenarios for the year 2100. This reduction in pH, which equates to atmospheric CO₂ levels of approximately 936 ppm², is hereafter referred to as the acidified treatment. Control treatments maintained an ambient pH level of 8.019 to 8.130, hereafter referred to as the control treatment. These experiments were repeated 4 times (deployments A-D), for periods between 72-144 hours, in the same area between January and February 2015. Figure 1 shows the pH offset for each deployment, as well as other water quality parameters (i.e. salinity and temperature), which did not change between deployments. The Submersible Photosynthesis-Respiration System comprised two chambers, which each included a Shutter Fluorometer fluorescence sensor (Aquation Pty Ltd, Umina Beach, Australia) to record fluorescence parameters and a dissolved oxygen (DO) probe (In-Situ Inc. Fort Collins, Colorado, USA) and two LI-192SA planar PAR sensors (Li-COR Inc., Lincoln, Nebraska, USA) to record environmental variables. These sensors were coupled to an underwater programmable data logger (Submersible Datalogger, Aquation Pty Ltd, Umina Beach, Australia) which recorded measurements over the following time scales: ϕ_{PSII} (F0 and Fm) every 30 minutes, Dissolved oxygen (DO) every 5 minutes, PAR every 5 minutes, temperature every 5 minutes. The chambers were flushed (for six minutes) every 6 hours with appropriate treatment or control water.

Chlorophyll fluorescence measurements. All *in-situ* Pulse Amplitude Modulated (PAM) fluorescence measurements (ϕ_{PSII}) are based on samples that had been exposed to natural irradiances, followed by 2–4 seconds of shading as the sensor closes into position prior to measurement. This brief interval of low to zero PAR had no detectable influence on photo acclimation, hence F_0 and F'_0 are reported as steady state minimum (F_0) and maximum (F_m) fluorescence (see⁶¹ for PAM fluorescence terminology). Yield ϕ_{PSII} was determined by the equation: yield $\phi_{PSII} = (F_m' - F')/F_m'$. Maximum photochemical quantum yield of PSII dark-adapted yield (F_v/F_m) was determined after 20 minutes of dark adaption and was calculated as $\phi_{PSII} = (F_m - F_0)/F_m$.

The minimum fluorescence value (F_0) was used as a proxy for biomass of MPB biofilms at the surface. This technique was validated by Serôdio *et al.* (2006) and was used to determine MPB sediment migration patterns. Concurrent time lapse photography (Cannon EOS600D and using a digisnap 2000) visually confirmed that the diel migration of MPB corresponded to a measurable F_0 diel pattern (unpublished data). Changes in F_0 under different PAR environments are expected to be minimal and consistent across both treatments. At the start of each deployment the Aquation software (Aquation Direct V2.0, Umina Beach, Australia) automatically adjusted the gain and auto zero. Rapid light curves (RLC) were used to calculate the relative Electron Transport Rate (relETR), which was calculated as: relETR = ($\Phi_{PSII} \times PAR$) (where values of absorbance and PSI:PSII are assumed to be in unity). Rapid light curves (RLCs) consist of eight steps of increasing actinic PAR for 10 seconds, with a saturating pulse measurement at the end of each actinic step. In all figures and tables the time of day is presented as Coordinated Universal Time (UTC + 8 hours) or elapsed time since beginning of deployment.

Net production. Dissolved oxygen sensors recorded every 5 minutes in each chamber. Linear regressions were applied to discrete blocks of data (180 minutes) with the slope coefficient used to measure the rate of change of Net production. These values were standardised by total microalgal sediment Chl *a* content.

Sample Collection. At the end of each experiment, five core samples for species identification and biomass and one core sample for grain size analyses were taken using PAR-proof syringes $(3 \text{ cm} \times 10 \text{ cm})$ and carefully extracted from the sediment, placed in a dark container and returned to the surface. Dark acclimated yield measurements (Fv/Fm) were taken for surface MPB after 20 minutes and then the cores were sliced every 2 mm down to 20 mm in the field. Each 2 mm section was immediately placed in the dark in cryotubes at -20 and divided in half for species identification and chlorophyll a (Chl a). These sample were then transported back to a -80 °C freezer (within 3 hours). Samples used for species identification were preserved in glutaraldehyde (4%) and stored at ~ 10 °C for later analyses. The remaining section of each core was placed on ice and transported back to a -80 °C freezer. All analyses were conducted at the Australian Antarctic Division (Hobart, Australia) or the Institute for Marine and Antarctic Studies laboratories (University of Tasmania, Hobart).

Environmental water and carbonate chemistry analysis. Sea water pH, salinity and temperature were recorded every 30 minutes via the antFOCE system³⁸. Dissolved inorganic carbon and total alkalinity samples were taken regularly to quantify the carbonate chemistry of treatments. The CO2SYS function in MATLAB (version 9.4.0) was used to calculate pH (from DIC, TA data). The pH (total scale) and salinity in each antFOCE chamber was determined at the time when water was pumped into the deployment chambers (Fig. 1).

Chlorophyll *a* **analysis.** Sediment pigment samples were extracted in methanol (99% HPLC grade sigma alrich) for 18 hours then analysed by fluorometer (Turner 10 AU, San Jose, California, USA). The fluorometer was calibrated with a Chl *a* standard (Sigma Alrich Chl *a*) and a solid standard (Turner) was used between each measurement to determine drift in signal. Hydrochloric acid (0.1%) was added to determine the phaeophytin component of each sample. These values were standardised by sediment weight.

Species identification. Samples for MPB species identification were centrifuged at low revolutions (300 rpm) for 15 minutes (1 gram in 10 ml of filtered seawater) to remove the larger sediment particles. The MPB cells were then diluted in 50 ml of filtered seawater. Subsequent aliquots (1 ml) of this dilution was then analysed on a Flowcam (Benchtop B3 series) with associated software (Visual spreadsheet, version 3.4). Only cells that had complete unbroken frustules with intact cell contents were counted. Species were identified according to taxonomic species keys in Scott and Marchant⁶². Total abundance was calculated from the cell counts of each species as a percentage of total cells counted.

Data analysis. Extracting first derivatives from individual deployment models. Separate generalized additive model (GAMs) were fitted to each deployment response (i.e. diel F_0 , ϕ_{PSII} and PAR) using elapsed time since the start of the deployment as the independent variable. Models assumed a Gaussian family for the response, with an identity link function, which proved adequate for these data. Each response variable was regressed against elapsed time of deployment using an adaptive P-spline with 20 basis functions, and 5 penalty basis functions determining the degree to which that flexibility was allocated across the covariate space⁶³. This type of adaptive smooth allows the degree of smoothing to vary along the covariate range, as dictated by the variability in the response. An optimal degree of smoothing was determined via cross-validation with final model fit and residuals assessed as suggested by Wood⁶³ (via k-index, adjusted R², and standard residual diagnostics). A temporal influence identified from these GAM model residuals directed us to allow effective degrees of freedom to be set at individual time points (using bs = "ad"). The final models had R-squared values between 0.87 + /-0.09 (1 sd), unless otherwise stated in figures. The first derivative (d/dt, i.e. the slope) for each response variable (diel F_0 , ϕ_{PSII} and PAR) in each deployment model above was extracted for 200 equally spaced segments on the predictor axis (elapsed time). This was carried out using the package Deriv R⁶⁴ in R. The first derivative, denoted here as didt represents the rate of change of each variable over the same elapsed time segment for each deployment. The function 'SignifD' in R was used to identify significant rates of change in PTRd/dt over time, with no PTR within a 12 hour period classed as non-migratory MPB. The data was then grouped into two separate categories (migrating MPB and non-migrating MPB). The first derivative data from the migrating deployments (A, B and D) were merged into a single dataset for analysis. The non-migrating MPB (deployment C) was examined separately.

Examining trends of diel F_0 and ϕ_{PSII} rates of change. To examine how the rate of change of PAR and the rate of change of MPB migration co-varied, we regressed the first derivatives of these variables in a combined dataset model. These GAMs were then used to identify the relationship between the rate of change in PAR and the rate of change in photo tactile behaviour or yield (ϕ_{PSII}) between treatments. Models assumed a Gaussian family for the response, with an identity link function, which proved adequate for these data. Each response variable was regressed against the corresponding first derivative of PAR using an adaptive P-spline with 8 basis functions, and 5 penalty basis functions⁶³. Both treatments were allocated the same basis function and model parameters. The model fit, residuals and effective degree of freedom were checked as above and as suggested by Wood⁶³.

To determine the rate of change of net production, coefficients of the regression of oxygen consumption in discrete $180 \, \text{minute}$ data blocks were determined. R squared values for fitted regressions of these coefficients were consistently high (>0.9). We make two assumptions with calculation of net production rates: the first is that heterotrophic community members do not change their metabolism under OA; and the second is that the heterotrophic communities are similar between treatments.

Fv/Fm was analysed using a nested ANOVA with deployment nested within treatment (n = 4).

Vertical sediment Chl a profiles were analysed on raw Chl a, and the percentage of Chl a present at each depth relative to the total Chl a in each core (relative Chl a). The data was grouped in migrating and non-migrating MPB biofilms. Core Chl a data was analysed by a two-way ANOVA with depth, treatment (n = 4).

All statistical analyses were carried out in R (version 3.2).

Data Availibilty

https://doi.org/10.4225/15/5a28b5eb04f71 and https://doi.org/10.26179/5c1827d5d6711

References

- Caldeira, K. & Wickett, M. E. Oceanography: anthropogenic carbon and ocean pH. Nature 425, 365-365, https://doi.org/10.1038/425365a (2003).
- 2. IPCC. Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change. 1535 (Cambridge, United Kingdom and New York, NY, USA, 2013).
- Brierley, A. S. & Kingsford, M. J. Impacts of climate change on marine organisms and ecosystems. Current Biology 19, R602–R614, https://doi.org/10.1016/j.cub.2009.05.046 (2009).
- Guinotte, J. M. & Fabry, V. J. Ocean acidification and its potential effects on marine ecosystems. Annals of the New York Academy of Sciences 1134, 320–342, https://doi.org/10.1196/annals.1439.013 (2008).
- Kroeker, K. J. et al. Impacts of ocean acidification on marine organisms: quantifying sensitivities and interaction with warming. Global Change Biology 19, 1884–1896, https://doi.org/10.1111/gcb.12179 (2013).
- Kroeker, K. J., Kordas, R. L., Crim, R. N. & Singh, G. G. Meta-analysis reveals negative yet variable effects of ocean acidification on marine organisms. Ecology Letters 13, 1419–1434, https://doi.org/10.1111/j.1461-0248.2010.01518.x (2010).
- 7. Kline, D. I. et al. A short-term in situ CO2 enrichment experiment on Heron Island (GBR). Scientific reports 2, https://doi.org/10.1038/srep00413 (2012).
- 8. Hoegh-Guldberg, O. & Bruno, J. F. The impact of climate change on the world's marine ecosystems. Science 328, 1523-1528 (2010).
- 9. Bellerby, R. *et al.* Marine ecosystem community carbon and nutrient uptake stoichiometry under varying ocean acidification during the PeECE III experiment. *Biogeosciences (BG)* 5, 1517–1527, https://doi.org/10.5194/bg-5-1517-2008 (2008).
- 10. Jokiel, P. L. et al. Ocean acidification and calcifying reef organisms: a mesocosm investigation. Coral Reefs 27, 473–483, https://doi.org/10.1007/s00338-008-0380-9 (2008).
- 11. Langdon, C. et al. Effect of elevated CO2 on the community metabolism of an experimental coral reef. Global Biogeochemical Cycles 17, 1011, https://doi.org/10.1029/2002GB001941 (2003).
- 12. Riebesell, Ua. G. J.-P. Lessons learned from ocean acidification research. *Nature Climate Change* 5, 12–14, https://doi.org/10.1038/nclimate2456 (2015).
- 13. McNeil, B. I. & Matear, R. J. Southern Ocean acidification: A tipping point at 450-ppm atmospheric CO2. *Proceedings of the National Academy of Sciences* 105, 18860–18864, https://doi.org/10.1073/pnas.0806318105 (2008).
- Barry, J. P. et al. Use of a Free Ocean CO2 Enrichment (FOCE) System to Evaluate the Effects of Ocean Acidification on the Foraging Behavior of a Deep-Sea Urchin. Environmental science & technology 48, 9890–9897, https://doi.org/10.1021/es501603r (2014).
- 15. Brewer, P. G., Gattuso, J.-P. & Hendriks, I. E. Free-ocean CO2 enrichment (FOCE) systems: Present status and future developments. Biogeosciences 11, 4057–4075, https://doi.org/10.5194/bg-11-4057-2014 (2014).
- Finkel, Z. V. et al. Phytoplankton in a changing world: cell size and elemental stoichiometry. Journal of Plankton Research 32, 119–137, https://doi.org/10.1093/plankt/fbp098 (2009).
- 17. Hurd, C. L., Hepburn, C. D., Currie, K. I., Raven, J. A. & Hunter, K. A. Testing the effects of ocean acidification on algal metabolism: considerations for experimental designs. *Journal of Phycology* 45, 1236–1251, https://doi.org/10.1111/j.1529-8817.2009.00768.x (2009).
- 18. Cartaxana, P. et al. Effects of elevated temperature and CO2 on intertidal microphytobenthos. BMC ecology 15, 10, https://doi.org/10.1186/s12898-015-0043-y (2015).
- 19. Vieira, S., Ribeiro, L., da Silva, J. M. & Cartaxana, P. Effects of short-term changes in sediment temperature on the photosynthesis of two intertidal microphytobenthos communities. *Estuarine, Coastal and Shelf Science* 119, 112–118, https://doi.org/10.1016/j.ecss.2013.01.001 (2013).
- Cartaxana, P. et al. Physiological versus behavioral photoprotection in intertidal epipelic and epipsammic benthic diatom communities. Journal of Experimental Marine Biology and Ecology 405, 120–127, https://doi.org/10.1016/j.jembe.2011.05.027 (2011)
- 21. Johnson, V. et al. Responses of marine benthic microalgae to elevated CO2. Marine Biology 160, 1813–1824, https://doi.org/10.1007/s00227-011-1840-2 (2013).
- Doney, S. C., Fabry, V. J., Feely, R. A. & Kleypas, J. A. Ocean acidification: the other CO2 problem. Marine Science 1, https://doi. org/10.1146/annurev.marine.010908.163834 (2009).
- 23. Mercado, J. M. & Gordillo, F. Inorganic carbon acquisition in algal communities: are the laboratory data relevant to the natural ecosystems? *Photosynthesis research* **109**, 257–267, https://doi.org/10.1007/s11120-011-9646-0 (2011).
- Marques da Silva, J., Cruz, S. & Cartaxana, P. Inorganic carbon availability in bethic diatom communities:photosynthesis and migration. *Philosophical Transactions of the Royal Society B: Biological Sciences* 372, https://doi.org/10.1098/rstb.2016.0398 (2017).
- 25. Jørgensen, B. B. In *Microbial Mats: Structure, Development and Environmental Significance* (eds Lucas J. Stal & Pierre Caumette) 243-253 (Springer Berlin Heidelberg, 1994).
- 26. Wang, J., Zhao, L. & Wei, H. Variable diffusion boundary layer and diffusion flux at sediment-water interface in response to dynamic forcing over an intertidal mudflat. Chinese Science Bulletin 57, 1568–1577, https://doi.org/10.1007/s11434-012-4988-3 (2012).

- 27. McMinn, A. & Hegseth, E. Quantum yield and photosynthetic parameters of marine microalgae from the southern Arctic Ocean, Svalbard. *Journal of the Marine Biological Association of the UK* 84, 865–871, https://doi.org/10.1017/S0025315404010112h (2004).
- 28. Mcminn, A., Runcie, J. W. & Riddle, M. Effect of seasonal sea ice breakout on the photosynthesis of benthic diatom mats at Casey, Antarctica. *Journal of phycology* 40, 62–69, https://doi.org/10.1046/j.1529-8817.2004.02127.x (2004).
- 29. McMinn, A., Ashworth, C. & Ryan, K. *In situ* net primary productivity of an Antarctic fast ice bottom algal community. *Aquatic Microbial Ecology* 21, 177–185, https://doi.org/10.3354/ame021177 (2000).
- 30. Moerdijk-Poortvliet, T. C. W. Carbon cycling in benthic diatom mats: Novel applications of LC/IRMS PhD thesis, Institute for Biodiversity and Ecosystem Dynamics (IBED) (2016).
- Raven, J. A. & Geider, R. J. Adaptation, acclimation and regulation in algal photosynthesis. Photosynthesis in algae 14, 385–412, https://doi.org/10.1007/978-94-0071038-2 17 (2003).
- 32. Roberts, K., Granum, E., Leegood, R. C. & Raven, J. A. Carbon acquisition by diatoms. *Photosynthesis Research* **93**, 79–88, https://doi.org/10.1007/s11120-007-9172-2 (2007).
- 33. Barranguet, C. & Kromkamp, J. Estimating primary production rates from photosynthetic electron transport in estuarine microphytobenthos. *Marine Ecology Progress Series* **204** (2000).
- 34. Wynne, D. & Rhee, G. Effects of light intensity and quality on the relative N and P requirement (the optimum N: P ratio) of marine planktonic algae. *Journal of Plankton Research* 8, 91–103, https://doi.org/10.1093/plankt/8.1.91 (1986).
- 35. Boyd, P. et al. Physiological responses of a Southern Ocean diatom to complex future ocean conditions. Nature Climate Change 6, 207–213, https://doi.org/10.1038/nclmate2811 (2015).
- J. Geider, R. & Osborne, B. A. Respiration and microalgal growth: a review of the quantitative relationship between dark respiration and growth. New Phytologist 112, 327–341, https://doi.org/10.1111/j.1469-8137.1989.tb00321.x (1989).
- 37. Serôdio, J., Coelho, H., Vieira, S. & Cruz, S. Microphytobenthos vertical migratory photoresponse as characterised by light-response curves of surface biomass. *Estuarine, Coastal and Shelf Science* **68**, 547–556, https://doi.org/10.1016/j.ecss.2006.03.005 (2006).
- 38. Stark, J. et al. Carbonate chemistry of an in-situ free-ocean CO2 enrichment experiment (antFOCE) in comparison to short term variation in Antarctic coastal waters. Nature Scientific Reports, https://doi.org/10.1038/s41598-018-21029-1 (2018).
- Ezequiel, J., Laviale, M., Frankenbach, S., Cartaxana, P. & Serôdio, J. Photoacclimation state determines the photobehaviour of motile microalgae: The case of a benthic diatom. *Journal of Experimental Marine Biology and Ecology* 468, 11–20, https://doi. org/10.1016/j.jembe.2015.03.004 (2015).
- 40. Jesus, B. et al. Adaptations of microphytobenthos assemblages to sediment type and tidal position. Continental Shelf Research 29, 1624–1634, https://doi.org/10.1016/j.csr.2009.05.006 (2009).
- 41. Perkins, R. et al. Responses of microphytobenthos to light: primary production and carbohydrate allocation over an emersion period. Marine Ecology Progress Series 223, 101–112 (2001).
- 42. Geider, R., MacIntyre, H. & Kana, T. Dynamic model of phytoplankton growth and acclimation: responses of the balanced growth rate and the chlorophyll a: carbon ratio to light, nutrient-limitation and temperature. *Marine Ecology Progress Series* 148, 187–200 (1997)
- 43. Hancke, K. et al. Effects of temperature and irradiance on a benthic microalgal community: A combined two-dimensional oxygen and fluorescence imaging approach. Limnology and Oceanography 59, 1599–1611, https://doi.org/10.4319/lo.2014.59.5.1599 (2014).
- 44. Rost, B., Zondervan, I. & Wolf-Gladrow, D. Sensitivity of phytoplankton to future changes in ocean carbonate chemistry: current knowledge, contradictions and research directions. *Marine ecology progress series*, 373237. 227, 227–237 (2008).
- 45. Trimborn, S., Brenneis, T., Sweet, E. & Rost, B. Sensitivity of Antarctic phytoplankton species to ocean acidification: Growth, carbon acquisition, and species interaction. *Limnology and Oceanography* 58, 997–1007, https://doi.org/10.4319/lo.2013.58.3.0997 (2013).
- Cook, P. L. & Røy, H. Advective relief of CO2 limitation in microphytobenthos in highly productive sandy sediments. *Limnology and Oceanography* 51, 1594–1601, https://doi.org/10.4319/lo.2006.51.4.1594 (2006).
- 47. Hoppe, C. J. M., Holtz, L.-M., Trimborn, S. & Rost, B. Ocean acidification decreases the light-use efficiency in an Antarctic diatom under dynamic but not constant light. *New Phytologist* 207, 159–171, https://doi.org/10.1111/nph.13334 (2015).
- 48. Johnson, M. D., Price, N. N. & Smith, J. E. Contrasting effects of ocean acidification on tropical fleshy and calcareous algae. *PeerJ* 2, e411, https://doi.org/10.7717/peerj.411 (2014).
- 49. Price, N. N., Hamilton, S. L., Tootell, J. S. & Smith, J. E. Species-specific consequences of ocean acidification for the calcareous tropical green algae Halimeda. *Marine Ecology Progress Series* 440, 67–78 (2011).
- Xu, J. & Gao, K. Future CO2-induced ocean acidification mediates the physiological performance of a green tide alga. Plant physiology 160, 1762–1769, https://doi.org/10.1104/pp.112.206961 (2012).
- 51. Olischläger, M., Bartsch, I., Gutow, L. & Wiencke, C. Effects of ocean acidification on different life-cycle stages of the kelp Laminaria hyperborea (Phaeophyceae). *Botanica Marina* 55, 511–525, https://doi.org/10.1515/bot-2012-0163 (2012).
- Olischläger, M. & Wiencke, C. Ocean acidification alleviates low-temperature effects on growth and photosynthesis of the red alga Neosiphonia harveyi (Rhodophyta). Journal of experimental botany 64, 5587–5597, https://doi.org/10.1093/jxb/ert329 (2013).
- Roháček, K. & Barták, M. Technique of the Modulated Chlorophyll Fluorescence: Basic Concepts, Useful Parameters, and Some Applications. Photosynthetica 37, 339–363, https://doi.org/10.1023/a:1007172424619 (1999).
- 54. Cornwall, C. E. et al. Diffusion boundary layers ameliorate the negative effects of ocean acidification on the temperate coralline macroalga Arthrocardia corymbosa. PloS one 9, https://doi.org/10.1371/journal.pone.0109468 (2014).
- 55. Longstaff, B. J. et al. An in situ study of photosynthetic oxygen exchange and electron transport rate in the marine macroalga Ulva lactuca (Chlorophyta). Photosynthesis Research 74, 281–293, https://doi.org/10.1023/a:1021279627409 (2002).
- 56. Pörtner, H. O., Langenbuch, M. & Michaelidis, B. Synergistic effects of temperature extremes, hypoxia, and increases in CO2 on marine animals: From Earth history to global change. *Journal of Geophysical Research: Oceans* 110, C09S10, https://doi.org/10.1029/2004JC002561 (2005).
- 57. Widdicombe, S. & Spicer, J. I. Predicting the impact of ocean acidification on benthic biodiversity: what can animal physiology tell us? *Journal of Experimental Marine Biology and Ecology* **366**, 187–197, https://doi.org/10.1016/j.jembe.2008.07.024 (2008).
- is. Cooley, S. R., Kite-Powell, H. L. & Doney, S. C. Ocean acidification's potential to alter global marine ecosystem services (2009)
- 59. Stark, J. S., Riddle, M. J. & Simpson, R. D. Human impacts in soft-sediment assemblages at Casey Station, East Antarctica: spatial variation, taxonomic resolution and data transformation. *Austral Ecology* 28, 287–304, https://doi.org/10.1046/j.1442-9993.2003.01289.x (2003).
- 60. Polmear, R., Stark, J., Roberts, D. & McMinn, A. The effects of oil pollution on Antarctic benthic diatom communities over 5 years. Marine pollution bulletin 90, 33–40, https://doi.org/10.1016/j.marpolbul.2014.11.035 (2015).
- 61. Consalvey, M., Perkins, R. G., Paterson, D. M. & Underwood, G. J. PAM fluorescence: a beginners guide for benthic diatomists. Diatom Research 20, 1–22, https://doi.org/10.1080/0269249X.2005.9705619 (2005).
- 62. Scott, F. J. & Marchant, H. J. Antarctic marine protists. (Csiro, 2005)
- 63. Wood, S. Generalized Additive Models: An introduction with R. (CRC Press, 2006).
- 64. Simpson, G. L. Modelling palaeoecological time series using generalized additive models. bioRxiv, https://doi.org/10.1101/322248 (2018).

Acknowledgements

This project was supported by the Australian Antarctic Division (AAS 4127), and under the Australian Research Council's Special Research Initiative for Antarctic Gateway Partnership (Project ID SR140300001), an Australian Research Council Linkage Infrastructure, Equipment and Facilities grant (LE130100220) and supported by the University of Tasmania Institute for Marine and Antarctic Studies, CSIRO Oceans and Atmosphere, and the Monterey Bay Aquarium Research Institute (USA). We are grateful to the following people for their support: Mark Milnes, Steven Whiteside, William Kirkwood, Edward Peltzer, Kym Newberry, Jono Reeve, Nick Roden, Matthew Wright, Helena Baird, Cathryn Wynn-Edwards, Martin Riddle, Ian Snape, Andrew Cawthorn, Clint Berry, Mike Sparg, Glenn Dunshea, Tony Press, Nick Gales, David Lyons and Kate Berry, as well as the personnel at Casey station for their logistical support throughout the field season. Permits from the Department of the Environment, Australian Government, were provided to conduct this research.

Author Contributions

J. Black and J Stark planned and designed the experiments, J. Black conducted the field experiments, prepared graphical figures, and performed statistical analyses. J. McKinlay, S. Wootherspoon and J. Black designed statistical analysis. J. Stark, G. Johnstone and J.Runcie assisted with fieldwork. All authors contributed to this manuscript.

Additional Information

Competing Interests: The authors declare no competing interests.

Publisher's note: Springer Nature remains neutral with regard to jurisdictional claims in published maps and institutional affiliations.

Open Access This article is licensed under a Creative Commons Attribution 4.0 International License, which permits use, sharing, adaptation, distribution and reproduction in any medium or format, as long as you give appropriate credit to the original author(s) and the source, provide a link to the Creative Commons license, and indicate if changes were made. The images or other third party material in this article are included in the article's Creative Commons license, unless indicated otherwise in a credit line to the material. If material is not included in the article's Creative Commons license and your intended use is not permitted by statutory regulation or exceeds the permitted use, you will need to obtain permission directly from the copyright holder. To view a copy of this license, visit https://creativecommons.org/licenses/by/4.0/.

© The Author(s) 2019